

# 地震活動異常と地殻変動異常 と前駆的非地震性すべり

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## Anomalies in seismic activity and transient crustal deformation

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# ETAS model:

$$\lambda(t) = \mu + \sum_{\{j:t_j < t\}} e^{\alpha\{M_j - M_c\}} \nu(t - t_j)$$

where

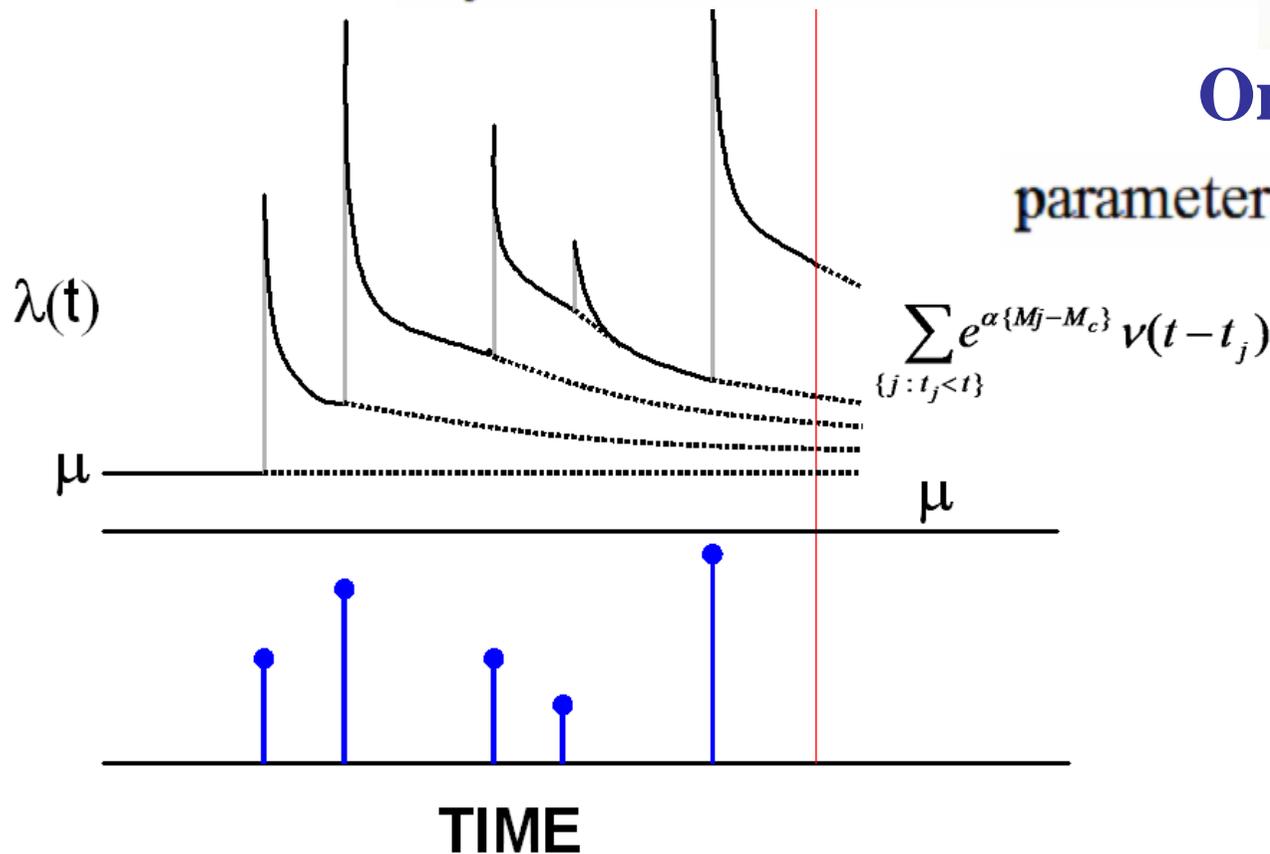
$t_j$  is occurrence time of  $j$ th event:

$M_j$  is magnitude of  $j$ th event;

$$\nu(t) = \frac{K}{(t + c)^p}$$

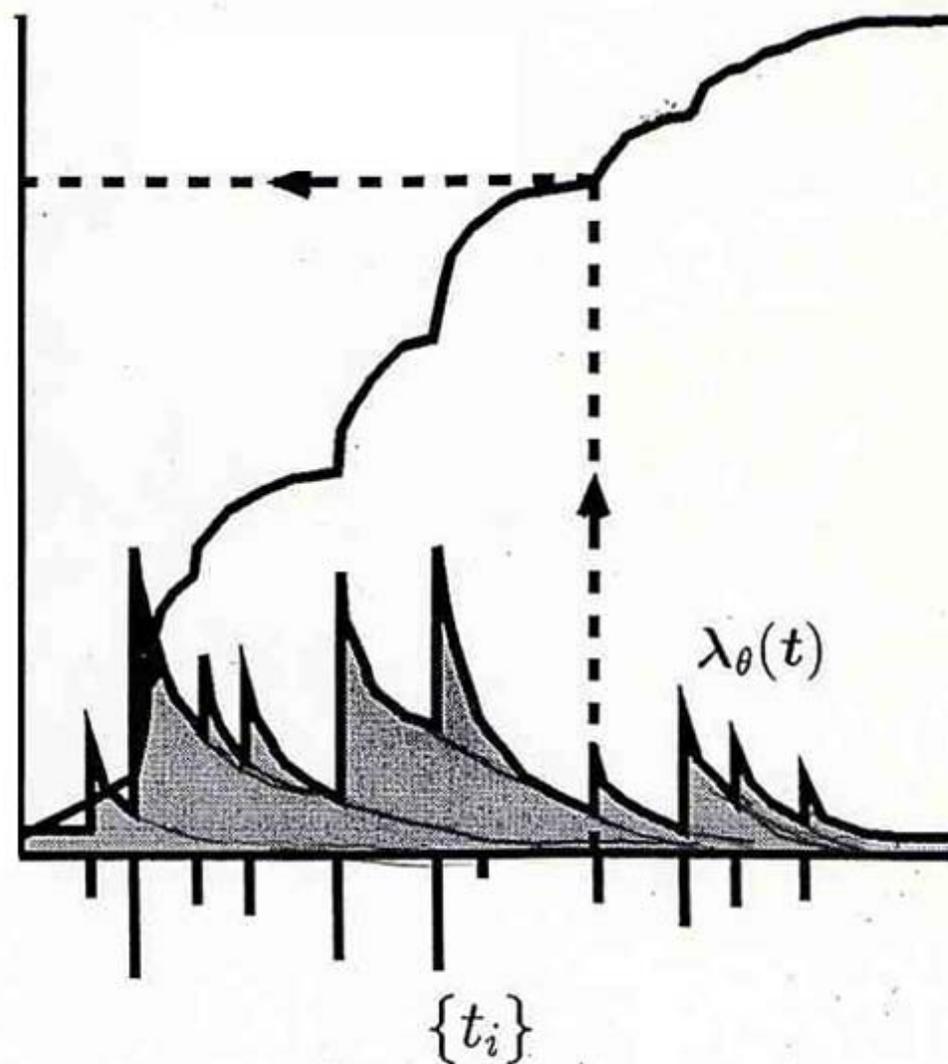
## Omori-Utsu formula

parameters are  $(\mu, K, c, \alpha, p)$ .

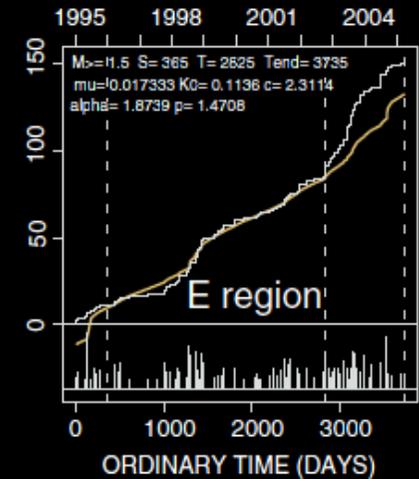
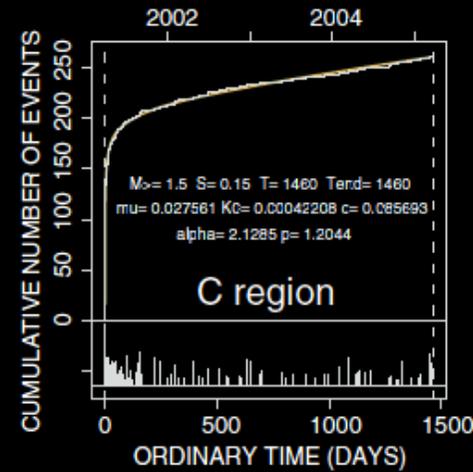
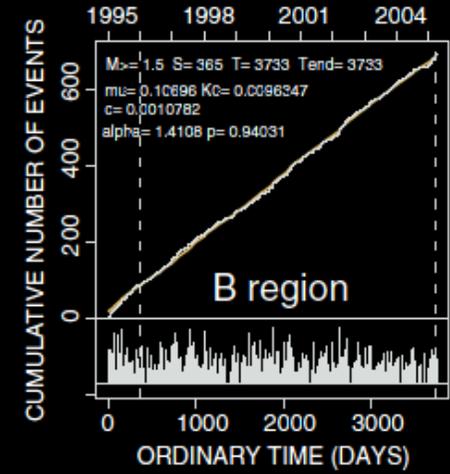
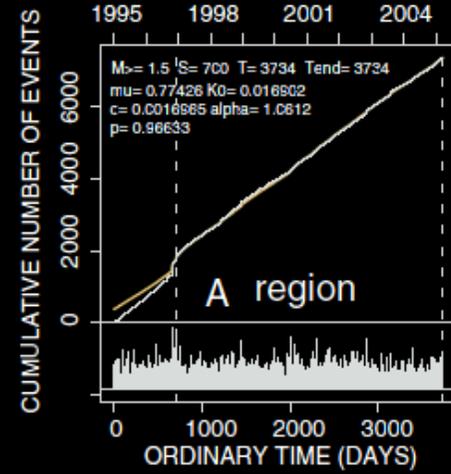
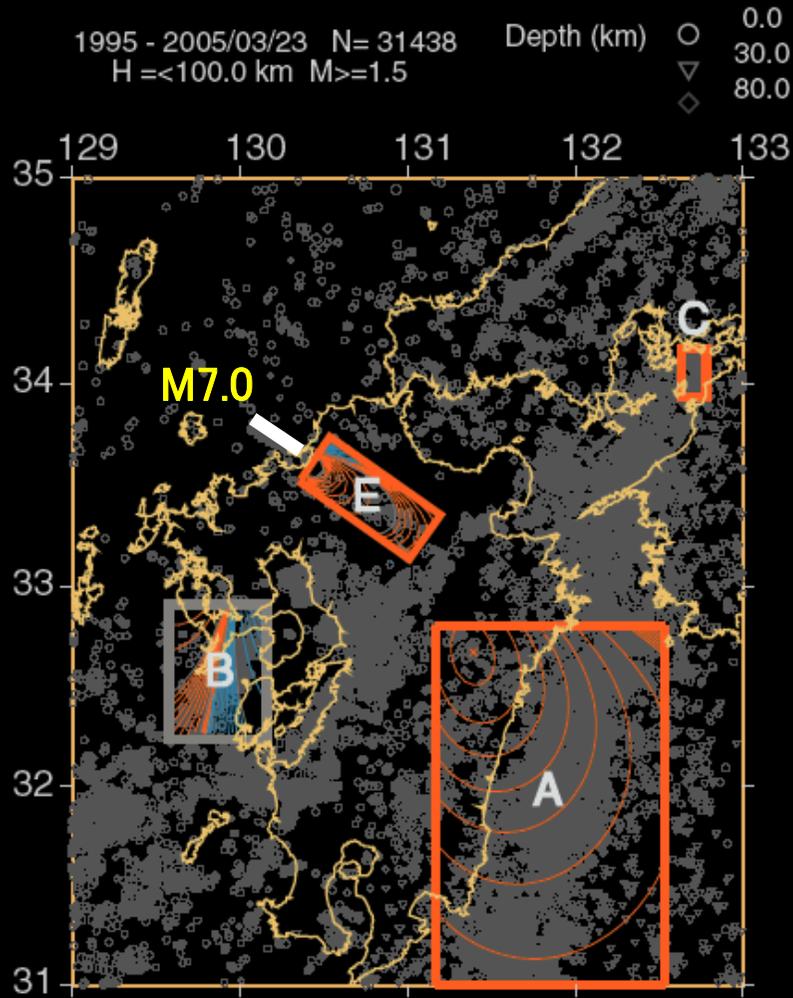


Theoretical cumulative  
number of the events:

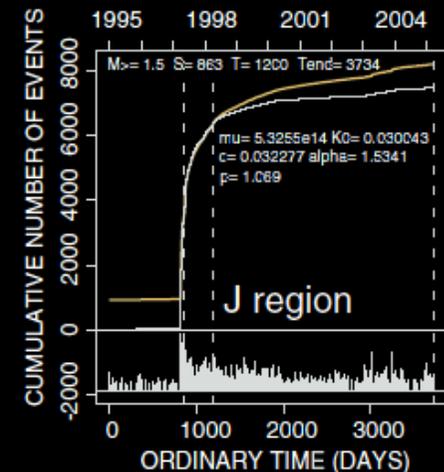
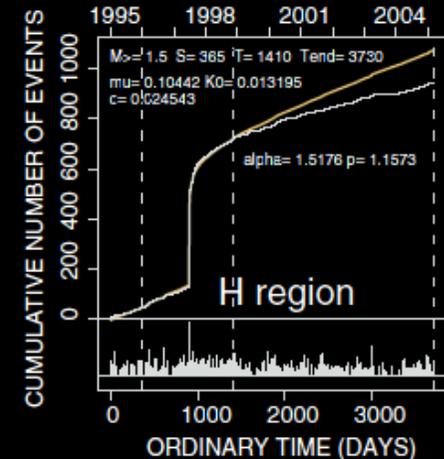
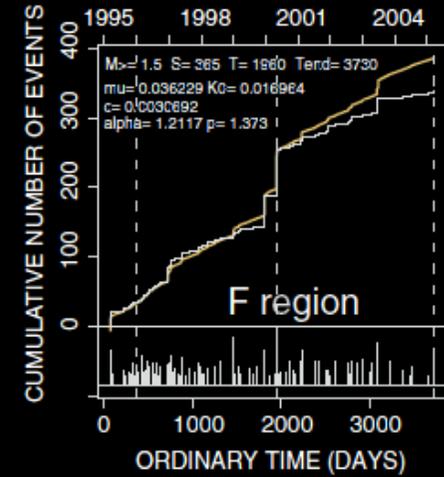
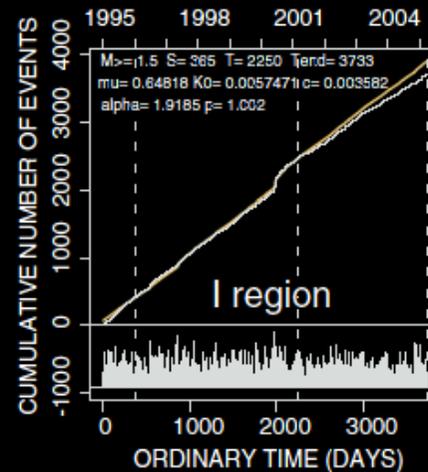
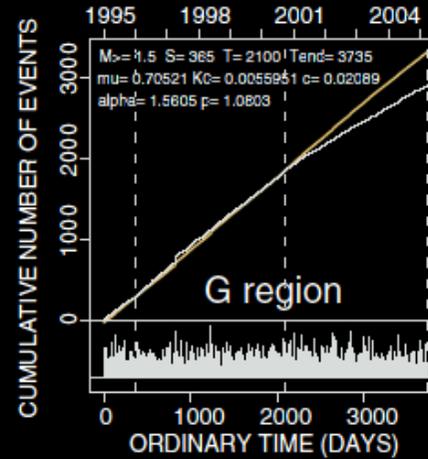
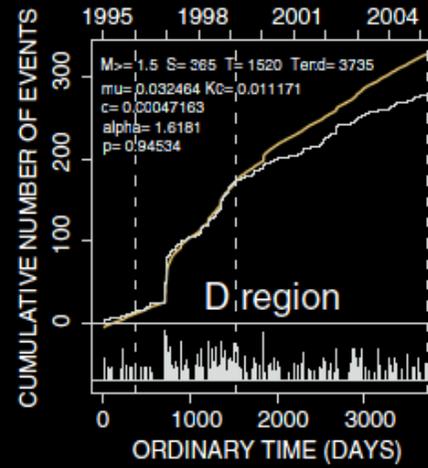
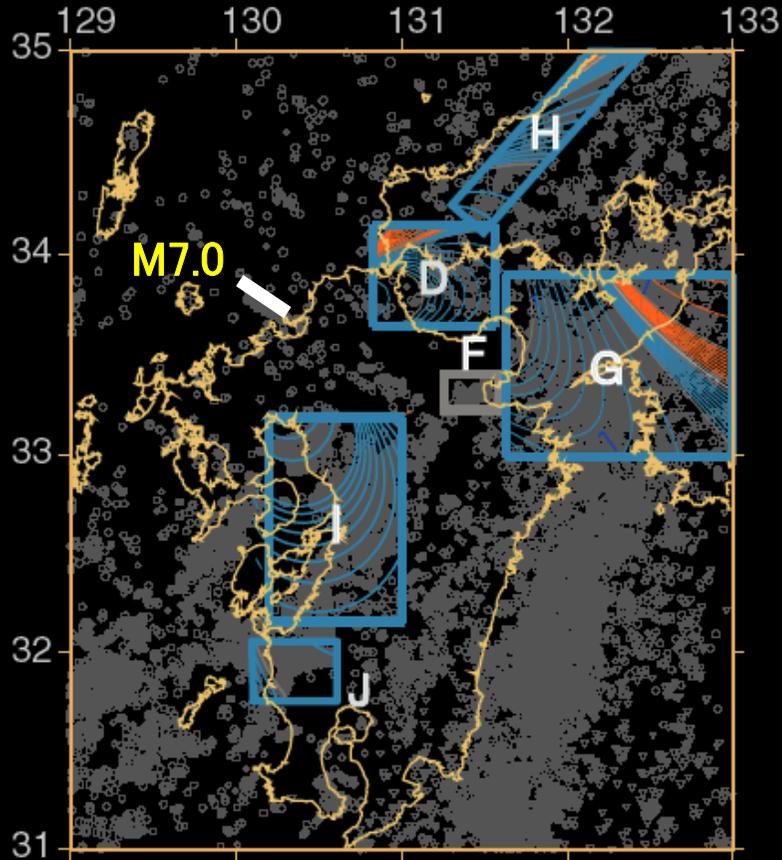
$$\Lambda(t) = \int_0^t \lambda(s) ds$$



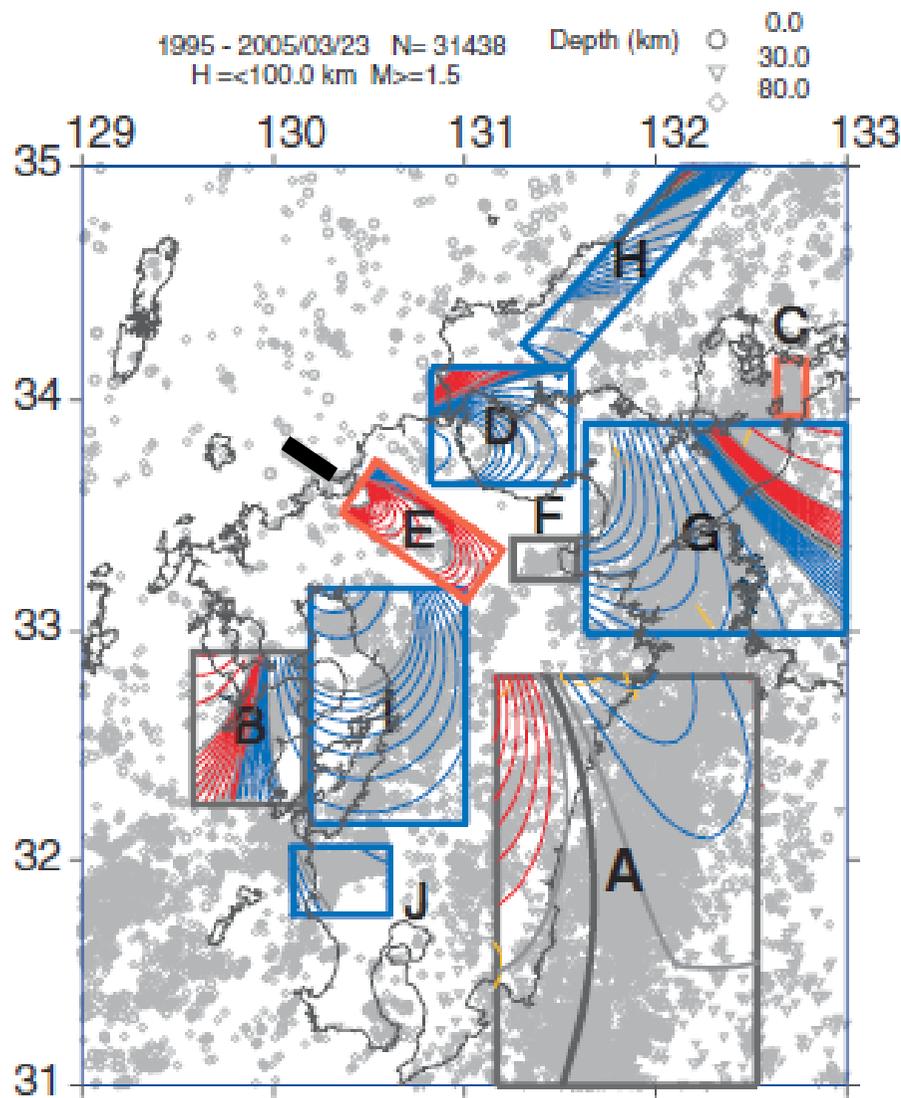
Ordinary Time



1995 - 2005/03/23 N= 31438 Depth (km) ○ 0.0  
 H =<100.0 km M>=1.5 ▽ 30.0  
 ◇ 80.0



# 福岡県西方沖のすべり



# 豊後水道のすべり

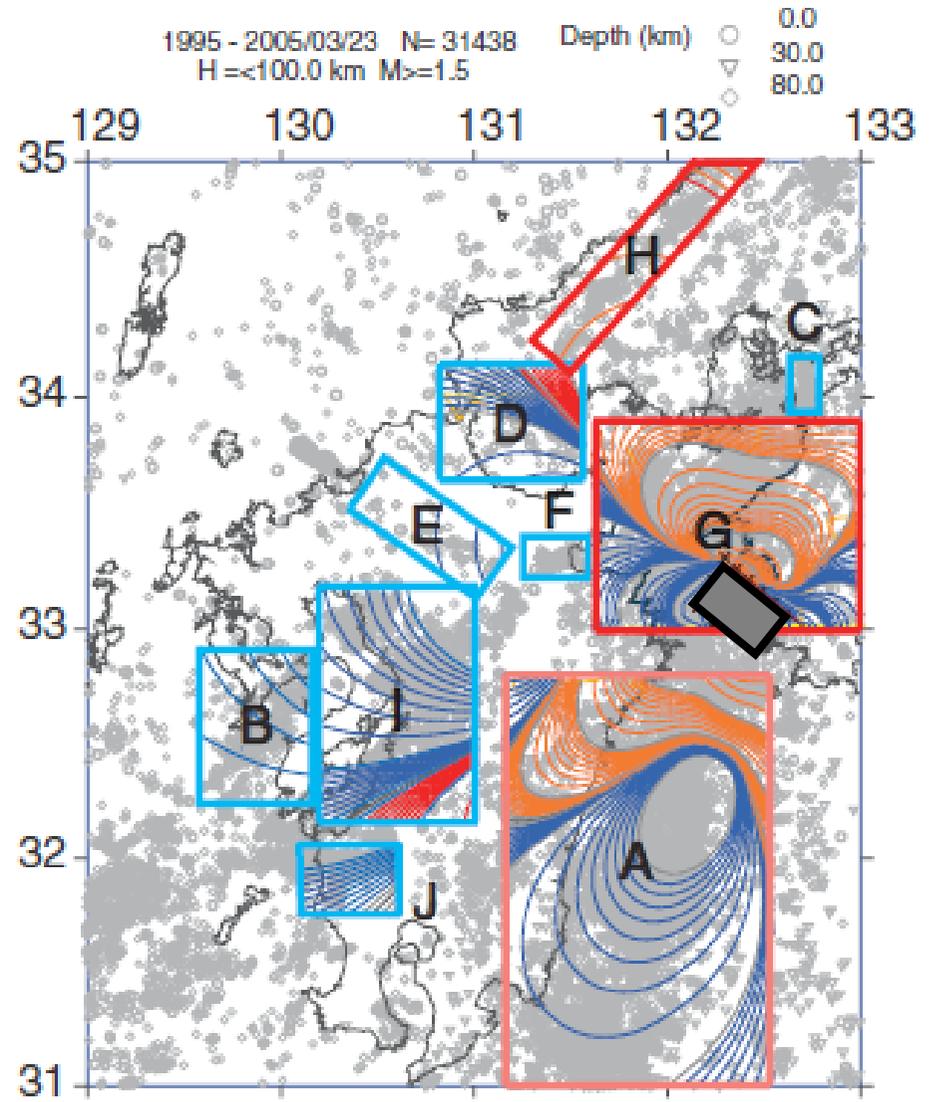


表 1 受け手の断層パラメタと $\Delta$ CFS 値

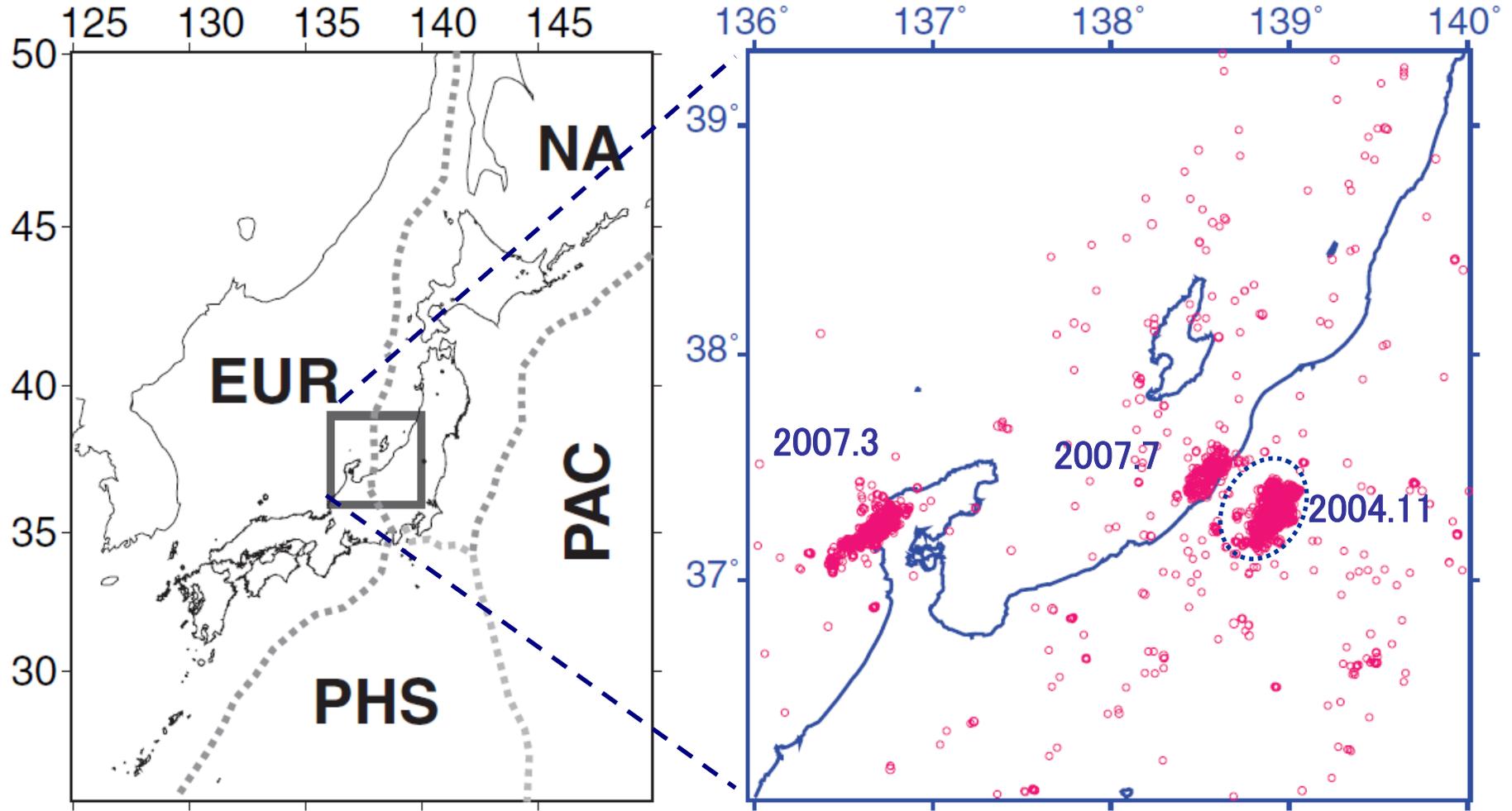
Table 1. Assumed receiver fault configurations and  $\Delta$ CFS values

領域	Strike (deg.)	Dip (deg.)	Rake (deg.)	Depth (km)	$\Delta$ CFS <sup>*</sup> (milli-bars)		Seismicity change ( $\Delta$ AIC <sup>**</sup> )
					福岡県西方沖	豊後水道	
A	210	30	90	--	0.	-50. ~ +150.	Normal
B	45	90	180	10	0.	-2.	Normal
	90	45	-90		0.	-1.	
C	179	55	-82	45	+4.	-8.	Normal
D	135	90	0	10	-2.	-7.	Quiet (-7.8)
E	135	90	0	10	+5. ~ +50.	-4.	Activate (+3.1)
F	90	90	180	10	-1.	-20.	Quiet (-1.8)
	90	45	-90		+1.	-8.	
G	170	75	-90	--	0.	-1. ~ +1. (bar)	Quiet (-75.2)
	330	35	-110		0.	-1. ~ +1. (bar)	
H	45	90	180	10	-1.	+8.	Quiet (-65.6)
	225	45	180		-4.	-2.	
I	45	90	180	10	-4.	-4.	Quiet (-29.2)
	90	80	-50		-4.	-2.	
	90	45	-90		-3.	-1.	
J	280	90	0	10	-1.	-0.3	Quiet (-194.2)

(\*) 福岡県西方沖の場合は本震の10%分の前駆すべり量を仮定している。

(\*\*) 正常な場合の ETAS と変化がある場合の 2 つ分の ETAS モデルの変化点補正済みの AIC の差 (変化点パラメタの調節に対して約 3.0 のペナルティを課している<sup>2,3)</sup>。

# 2004 Chuetsu M6.8

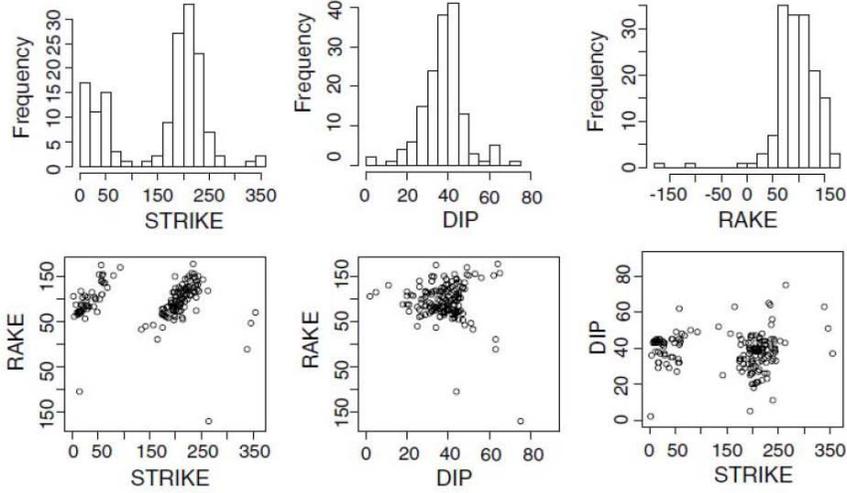


1997/01/01 - 2004/10/30 Number of events = 4060

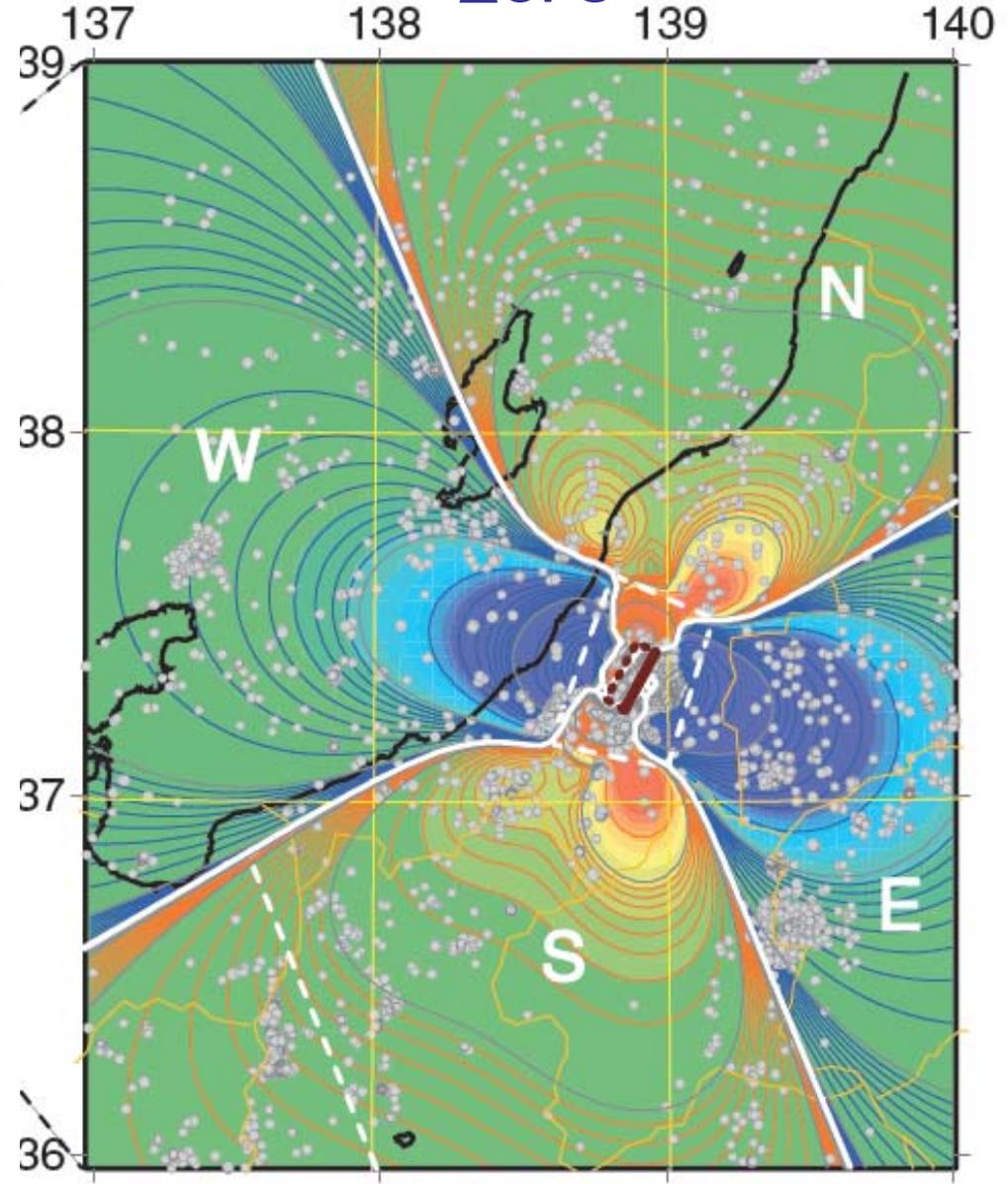
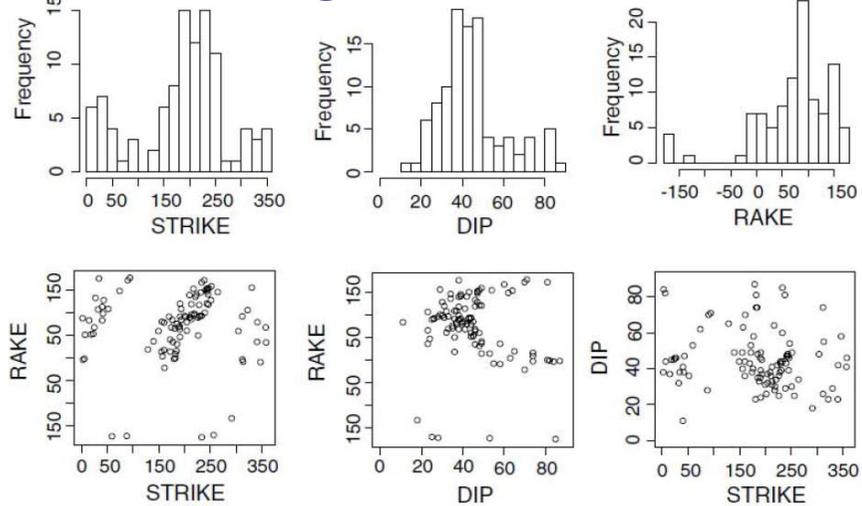
Depth =< 25.0 km, M >= 2.0

$\Delta$ CFS

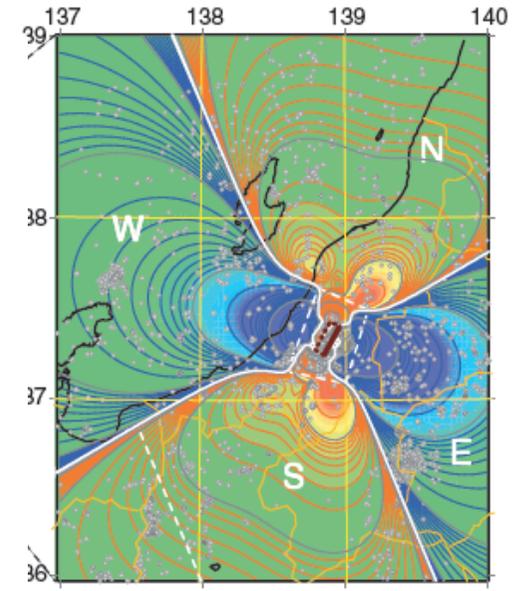
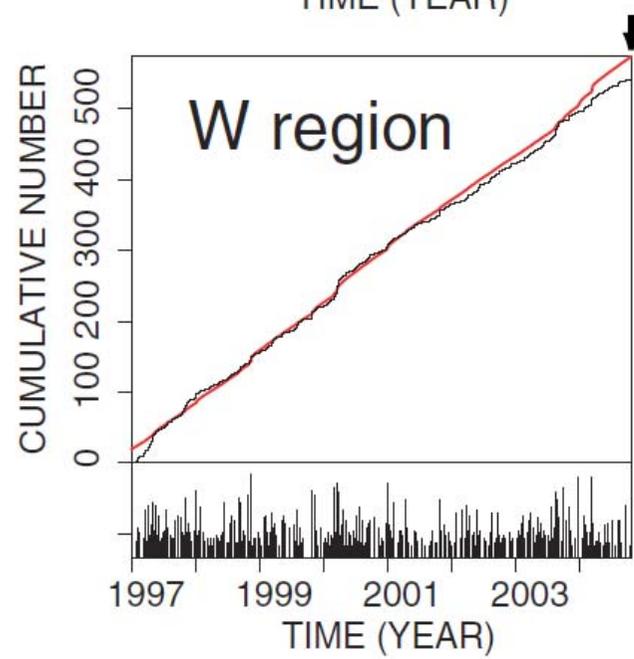
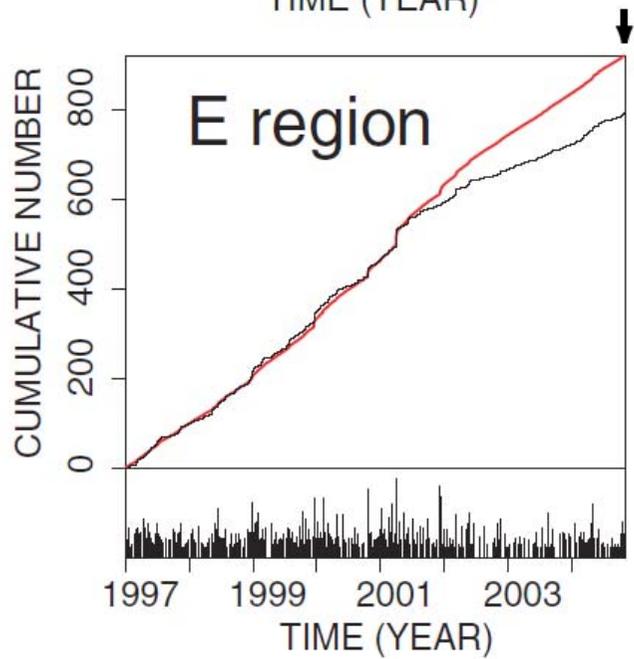
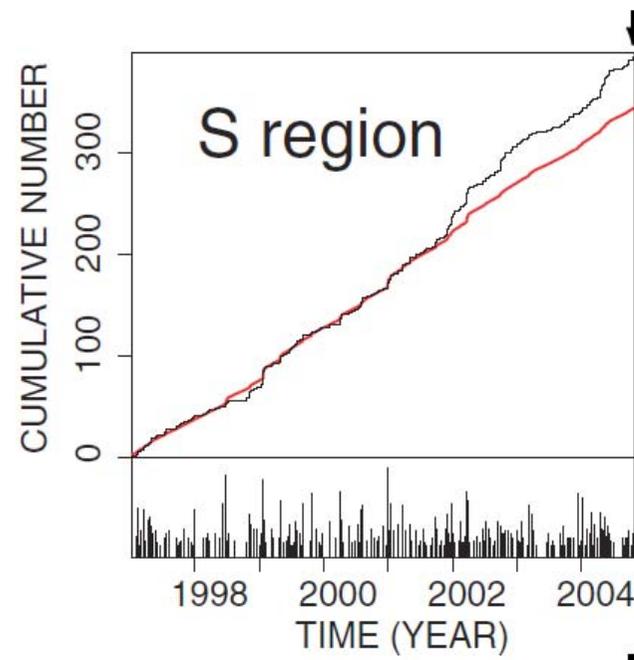
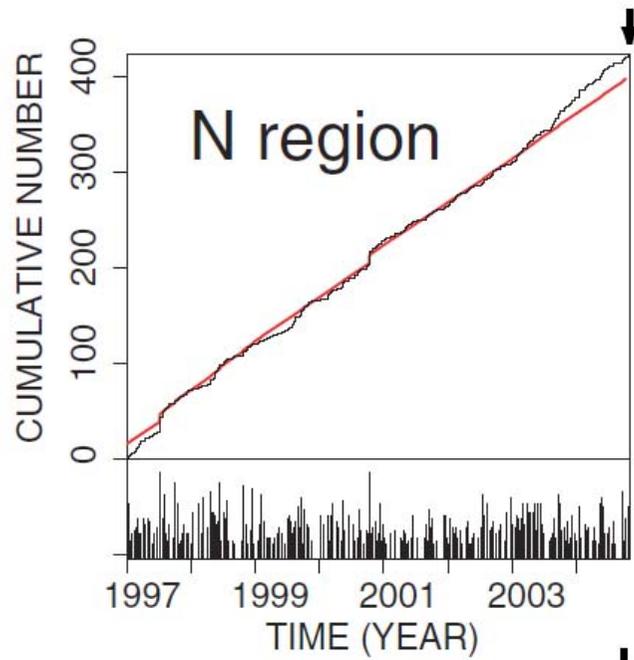
(a) aftershocks



(b) background



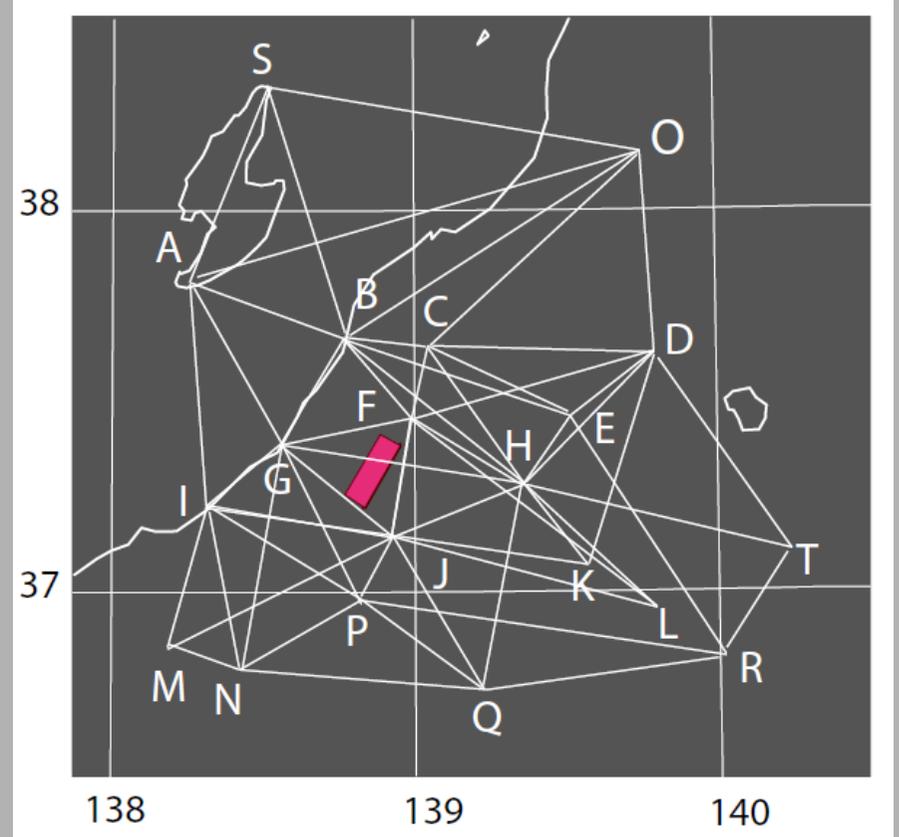
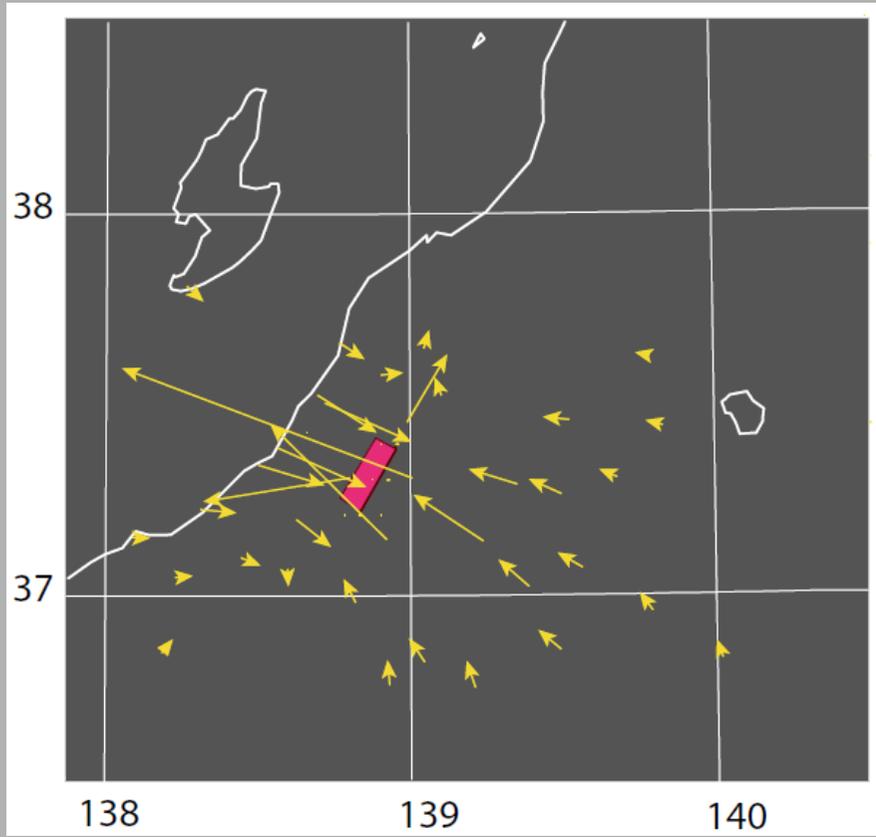
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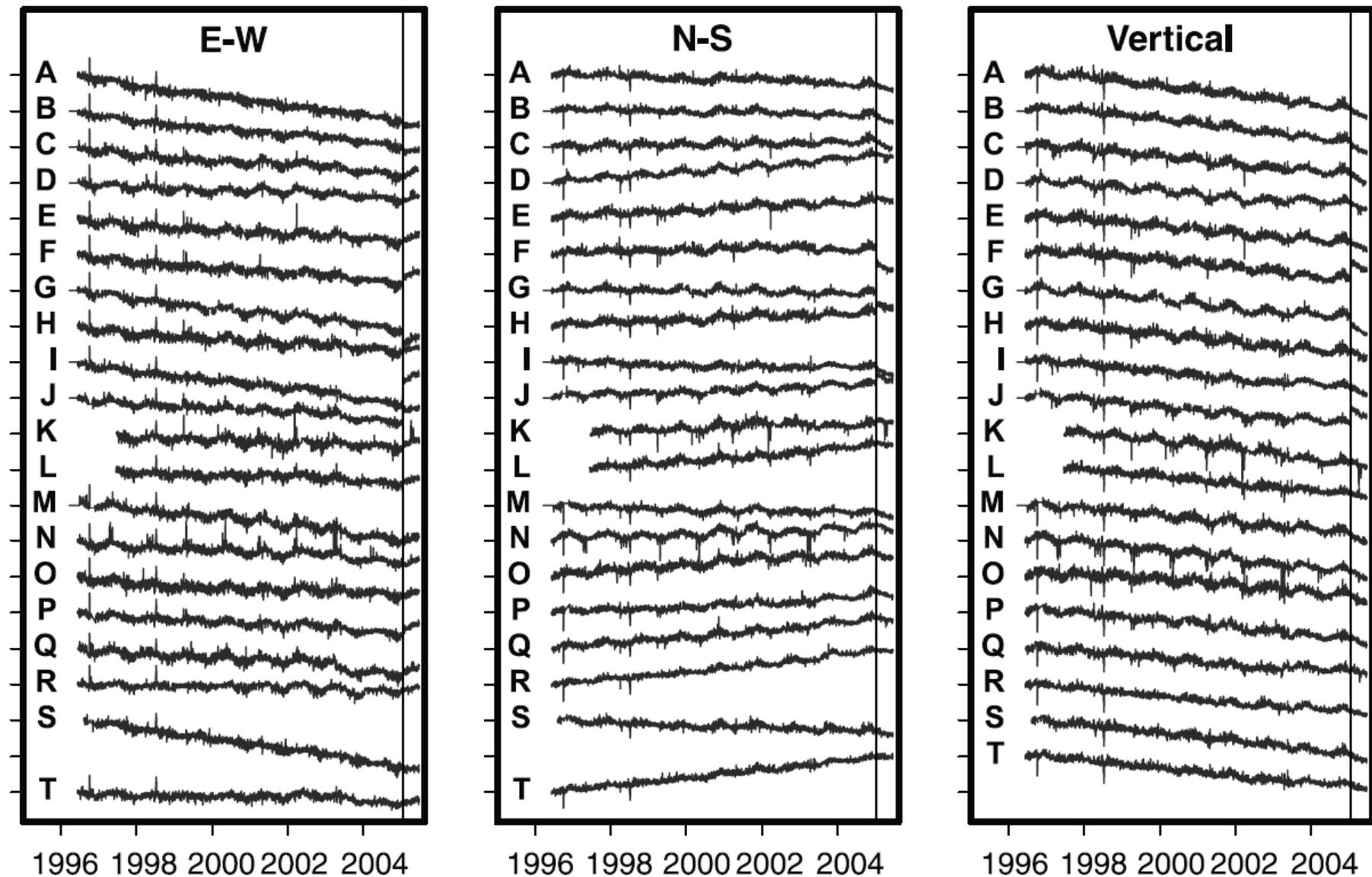


$\Delta$ CFS

**ETAS cumulative curve**



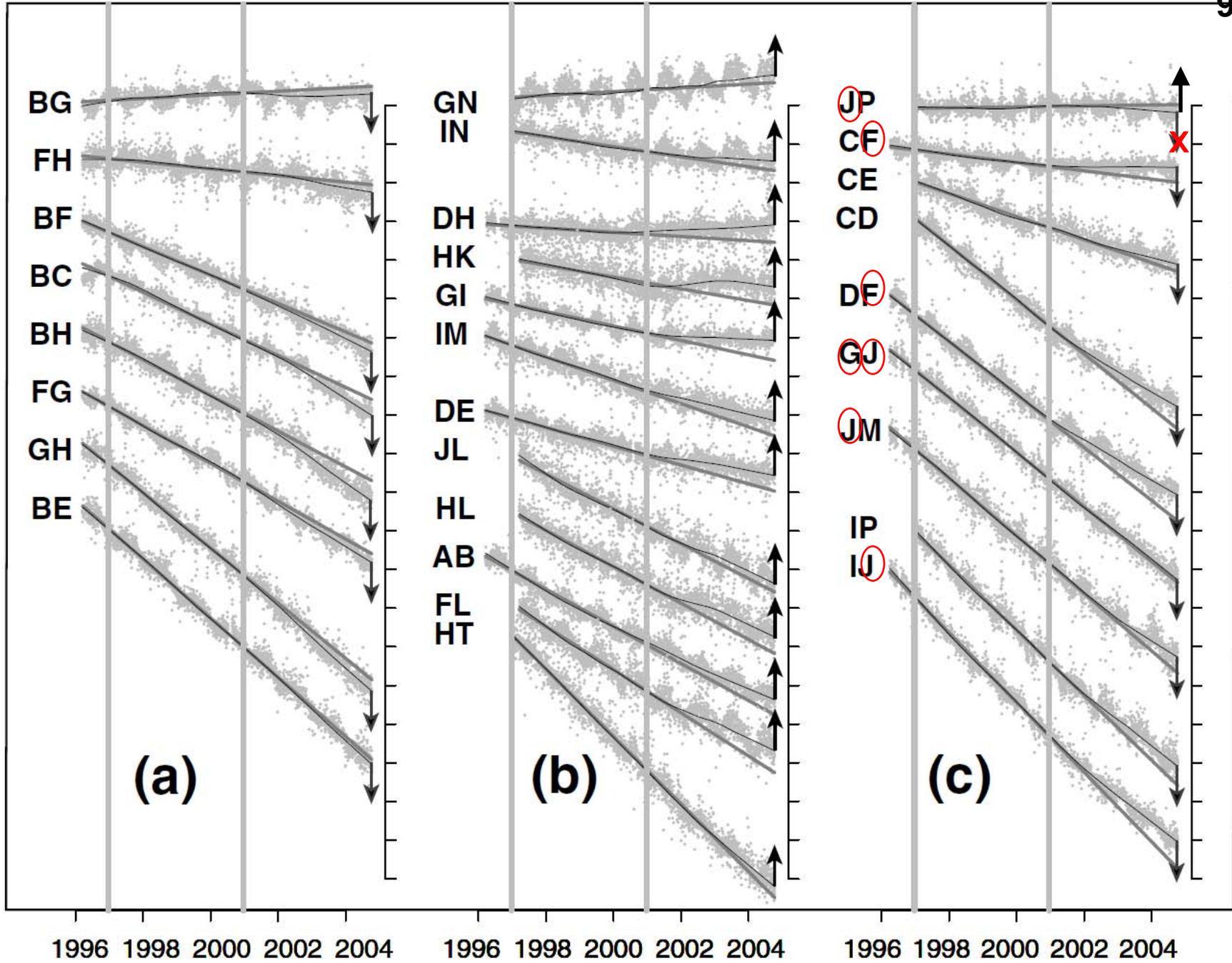




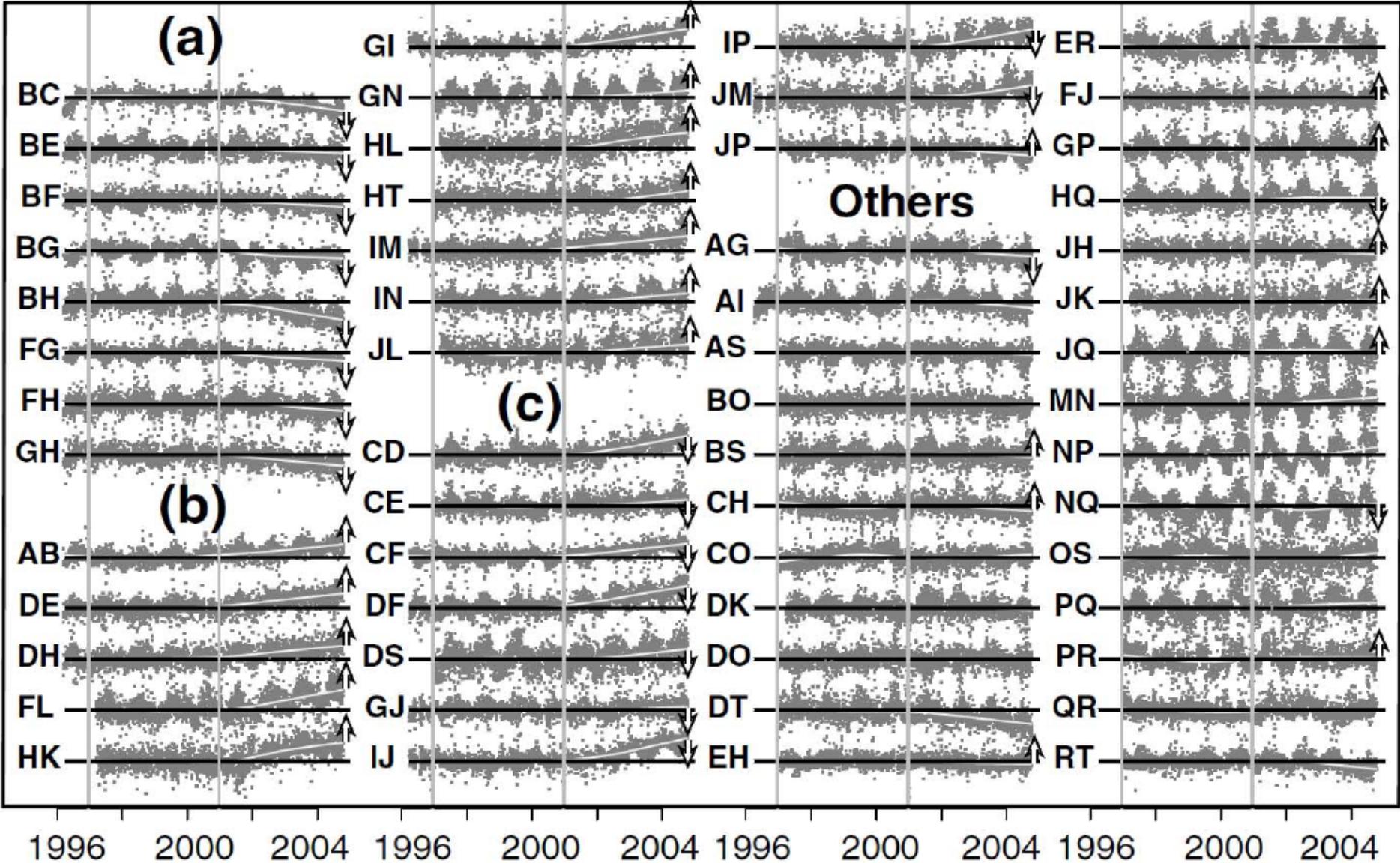
**Figure 7.** Daily position estimates for the (left) horizontal west-east, (middle) horizontal north-south, and (right) vertical coordinates of the GEONET stations A–T relative to their medians. The marked unit in vertical axis indicates 10 cm in distance. The thin vertical lines in 2004 indicate the occurrence time of the Chuetsu earthquake.

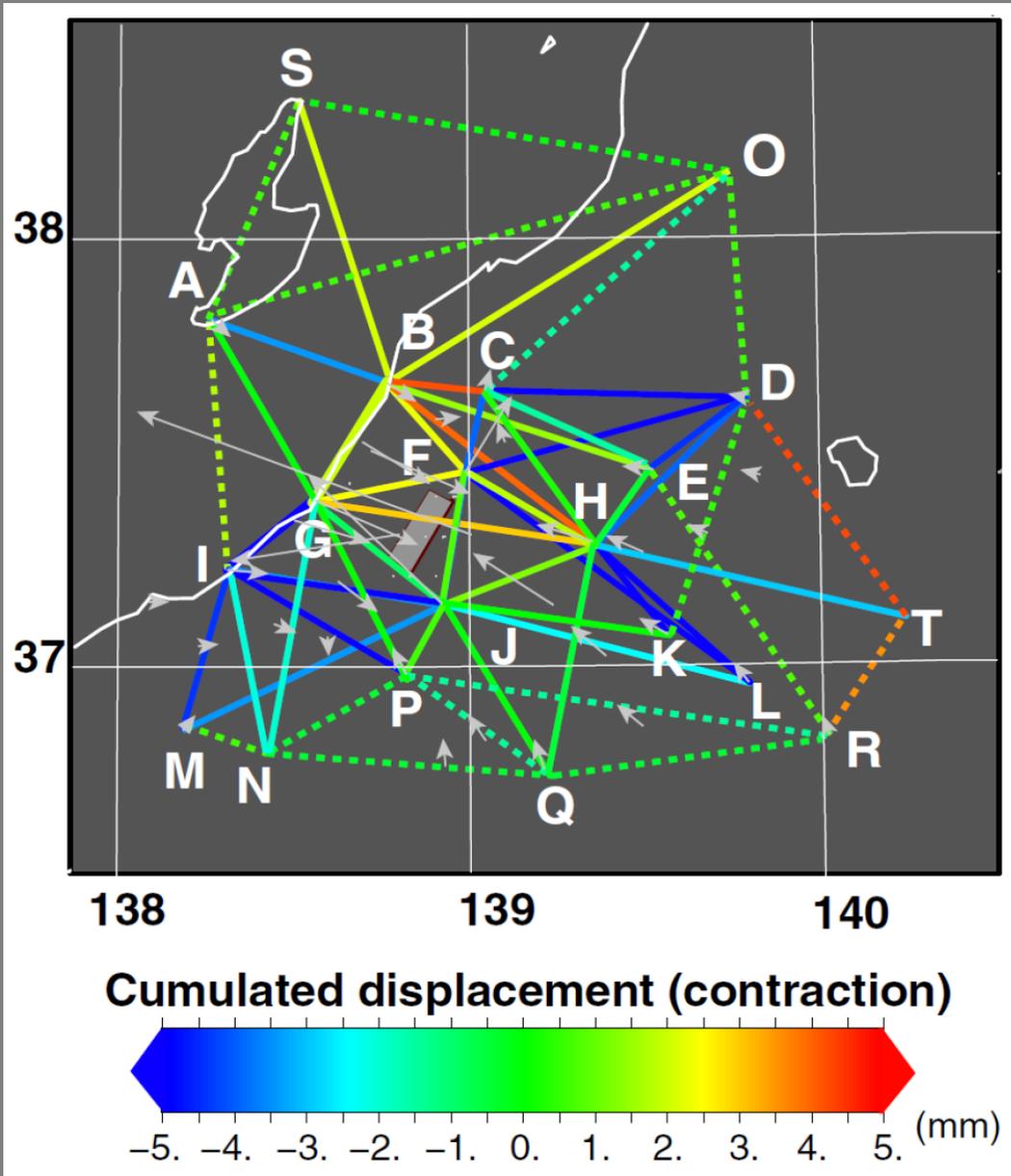
# Daily time series of distance between stations

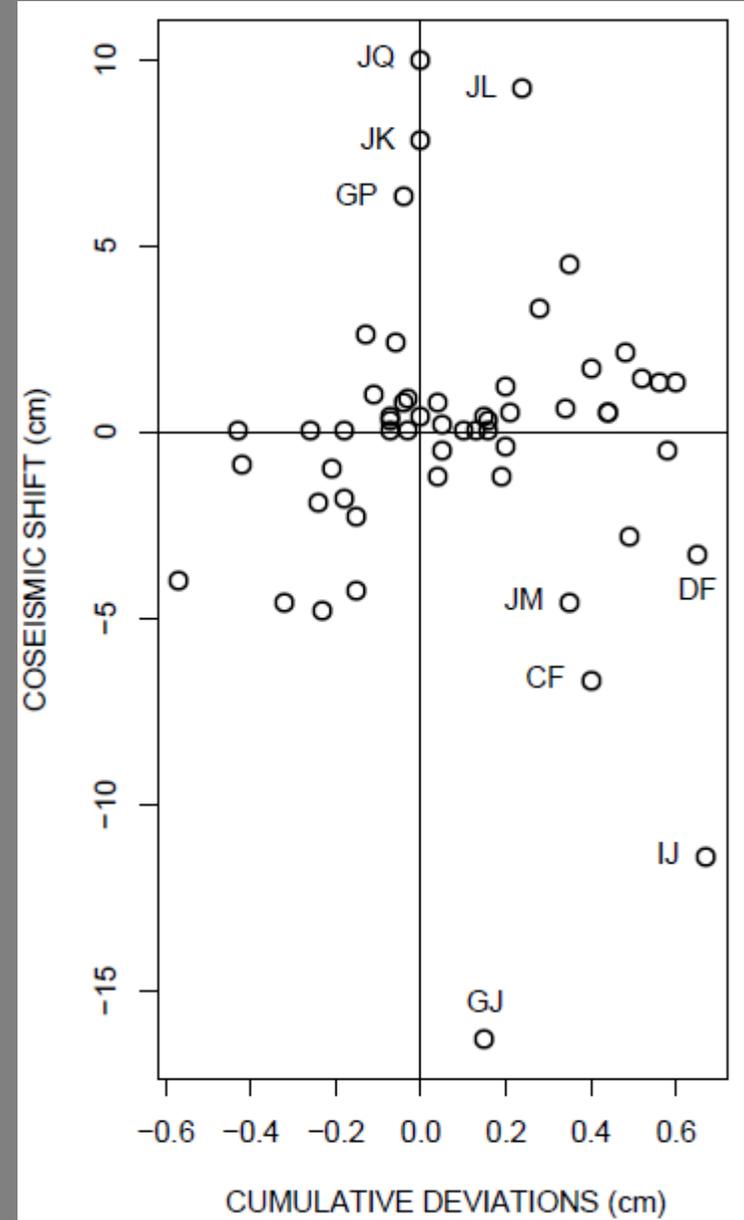
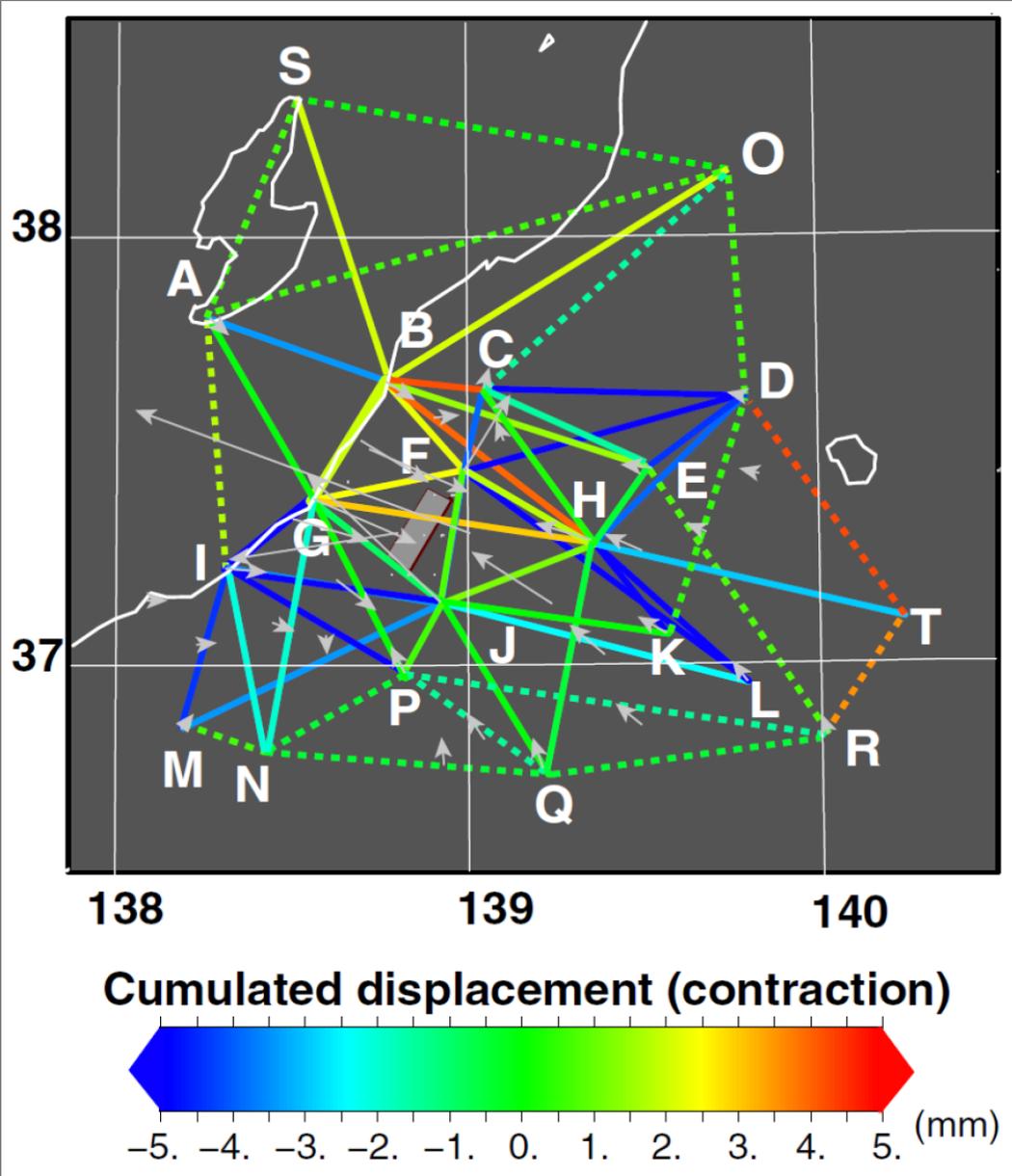
Extension ↑  
Contraction ↓



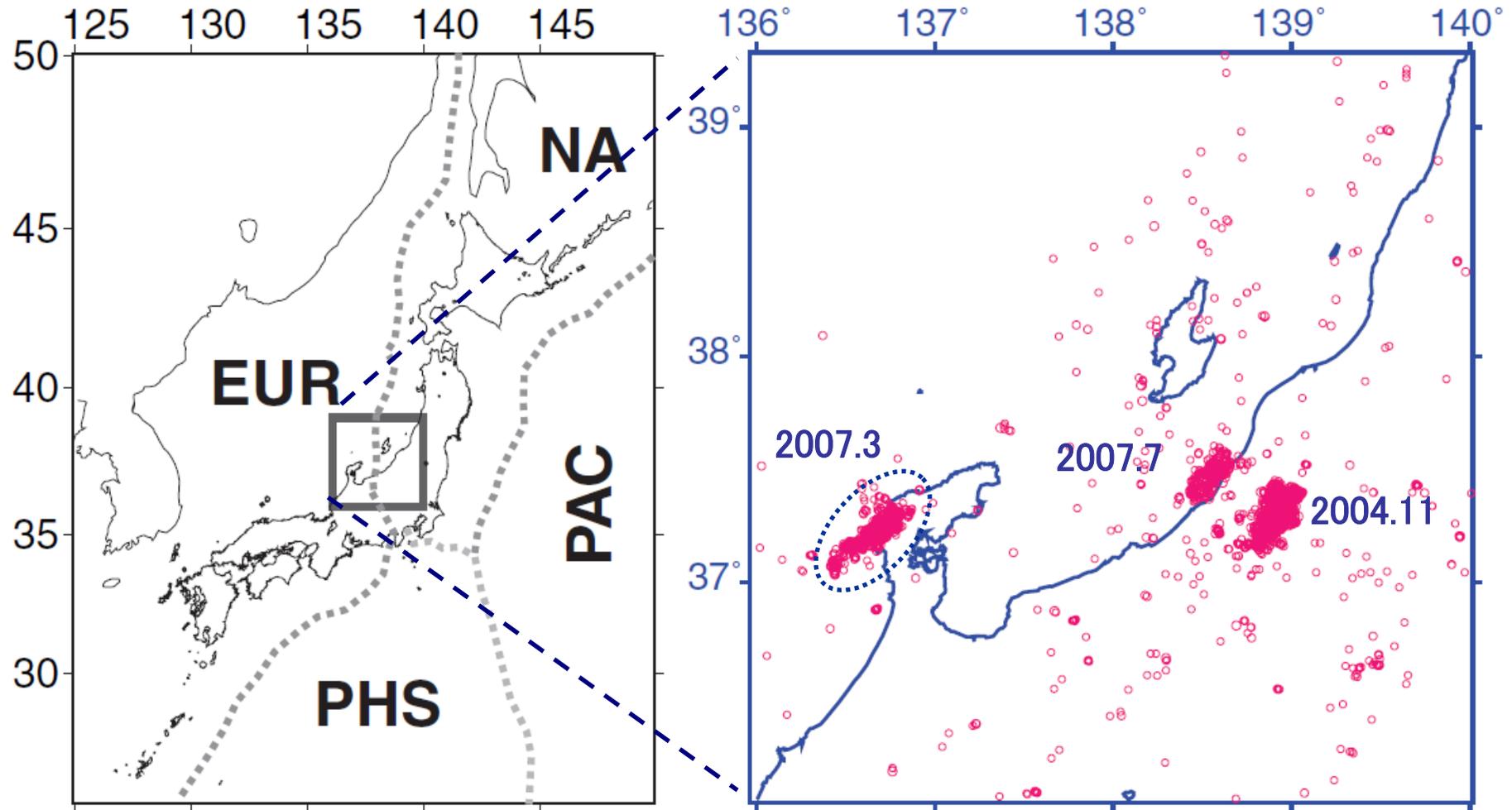
# Residual time series of distances

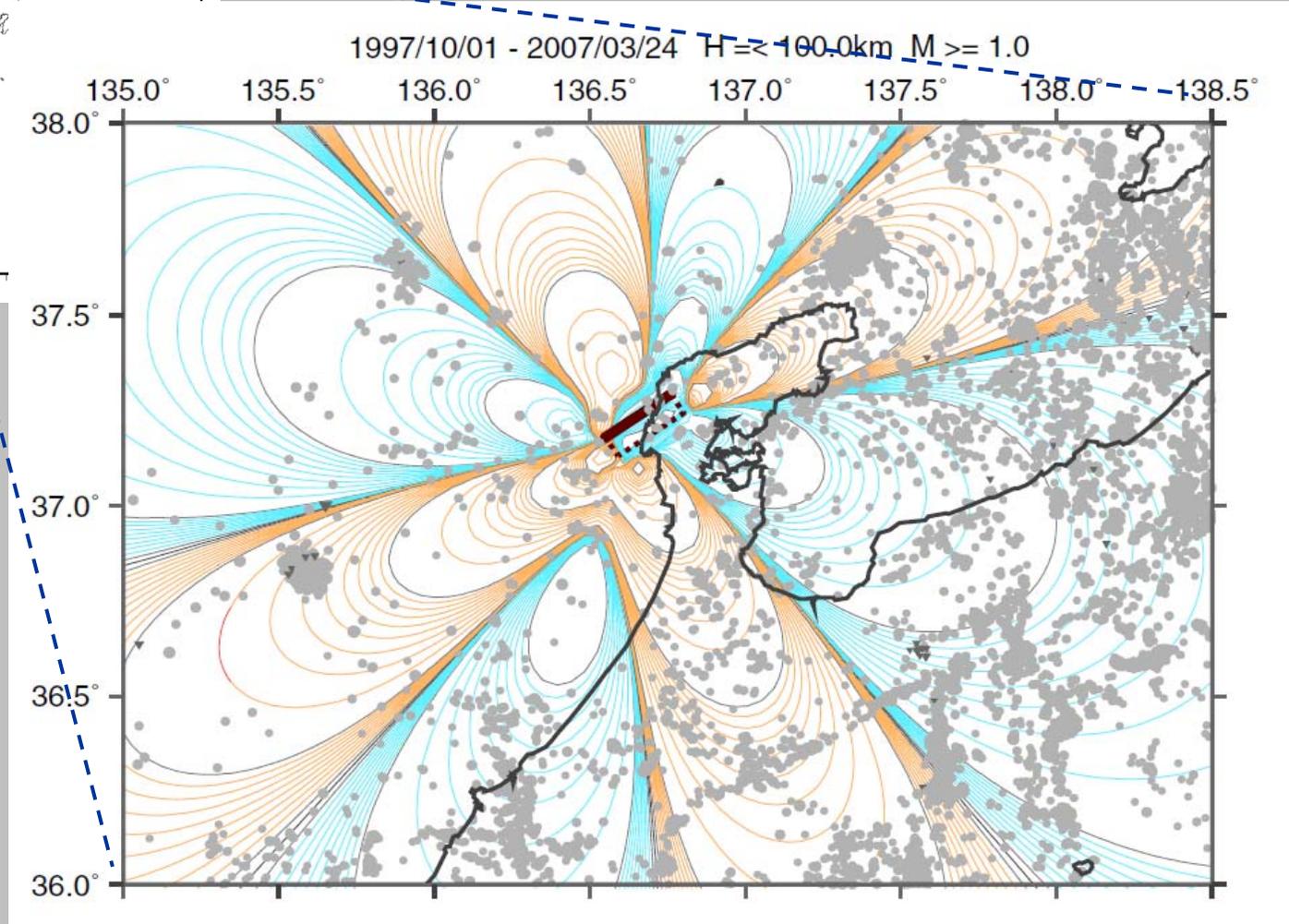
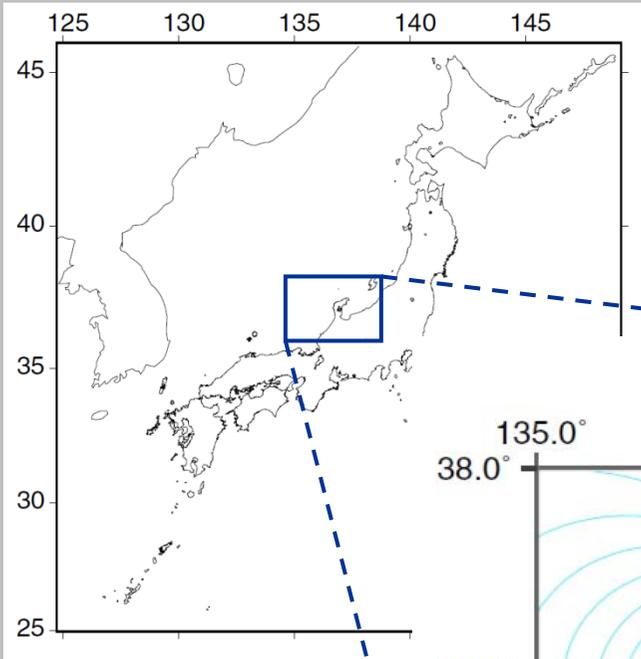


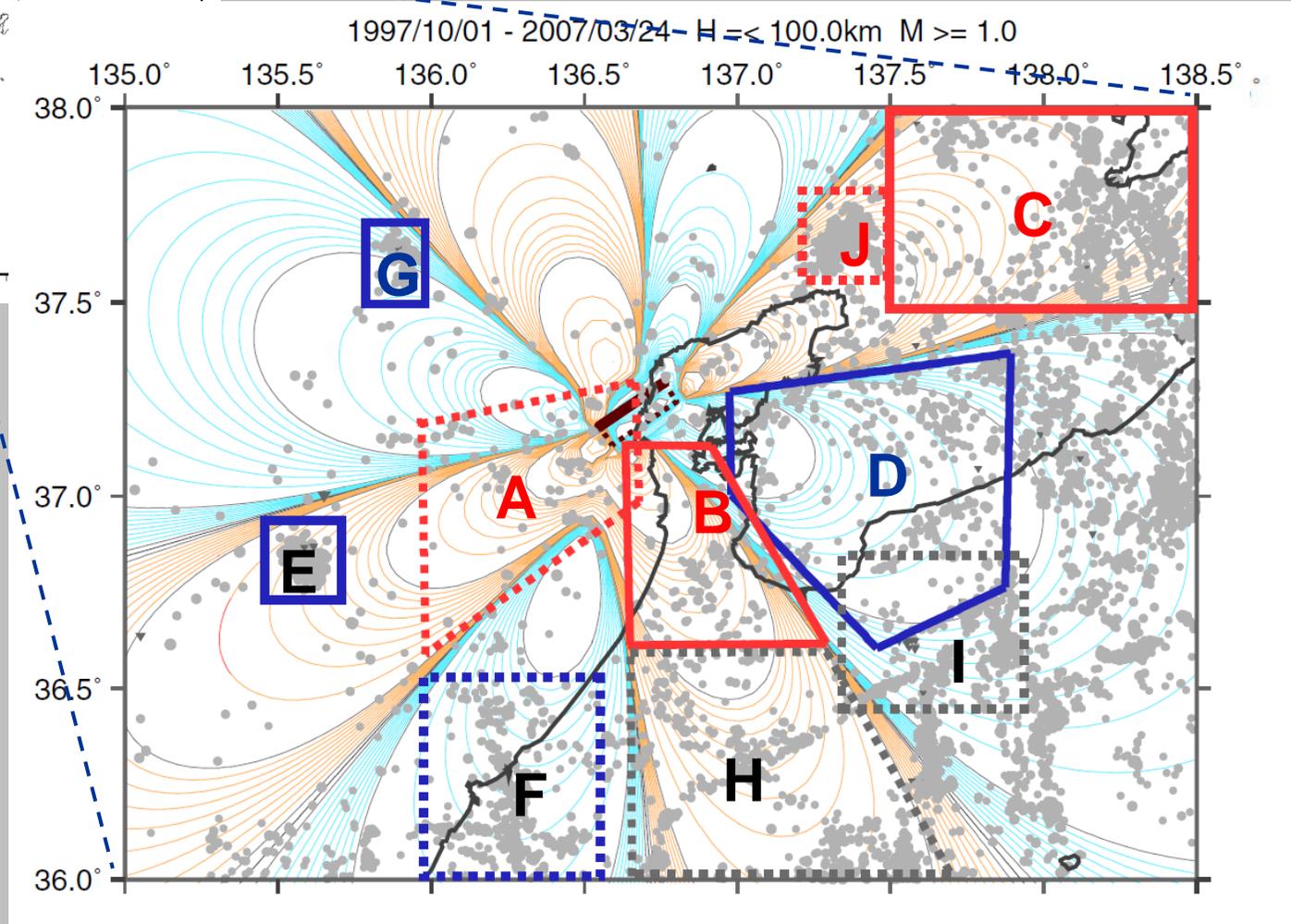
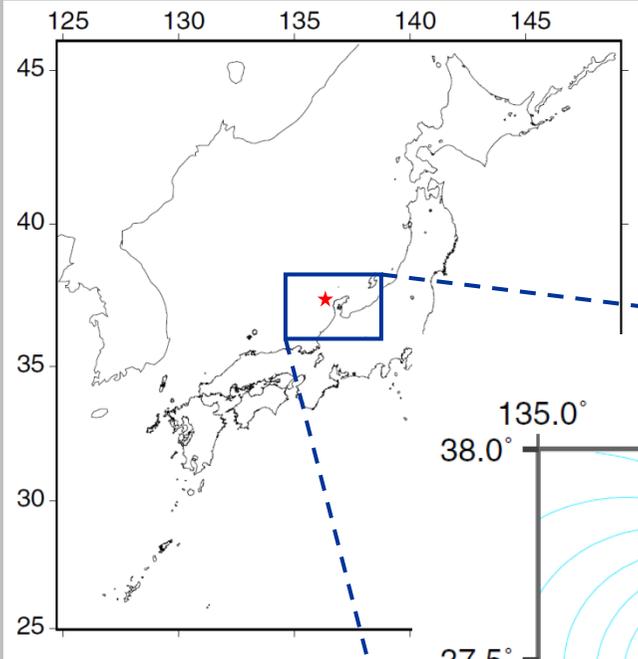


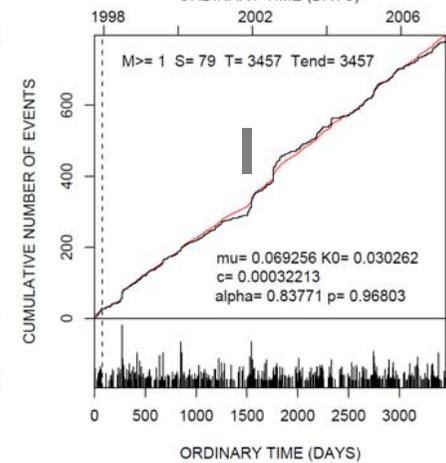
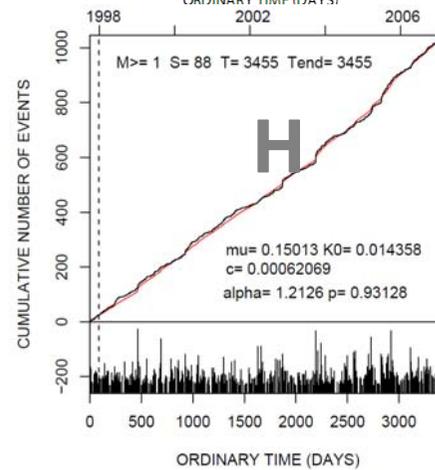
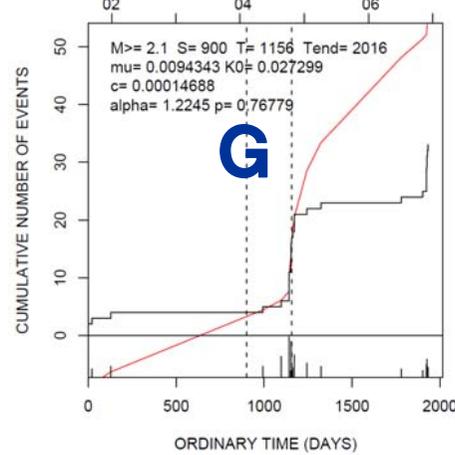
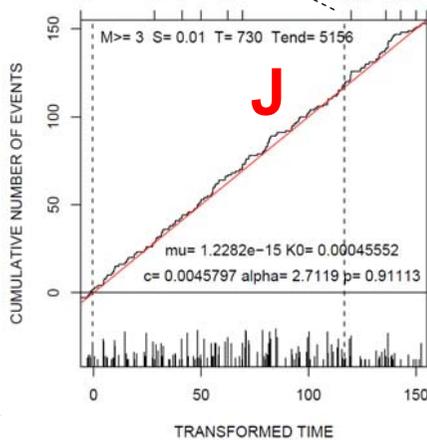
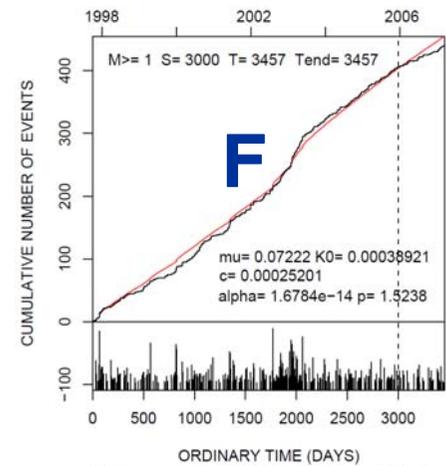
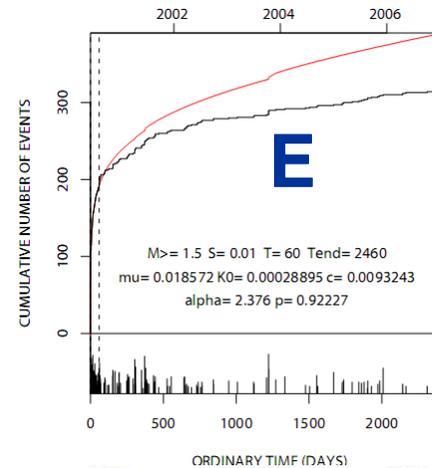
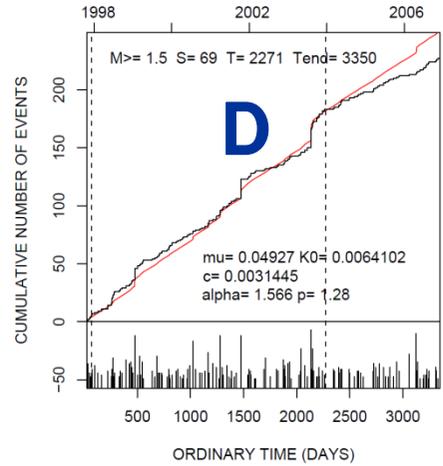
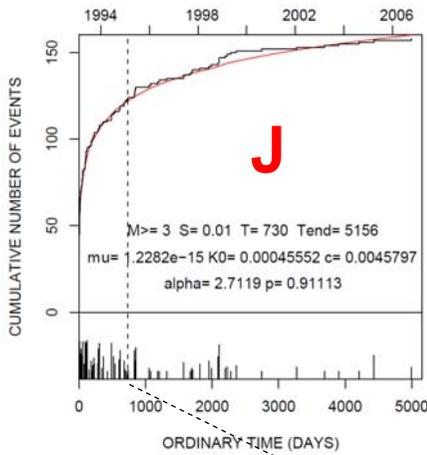
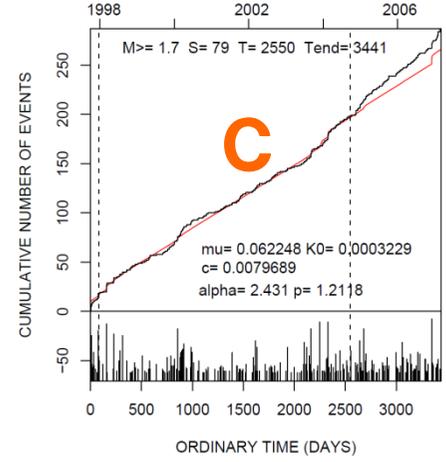
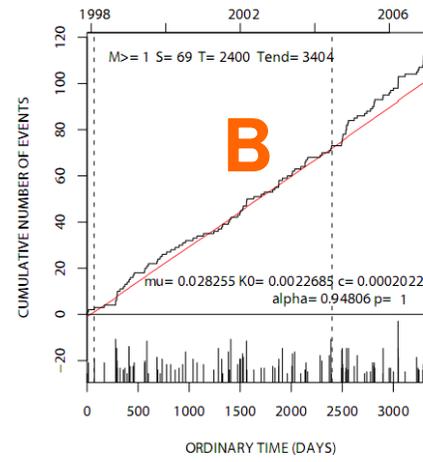
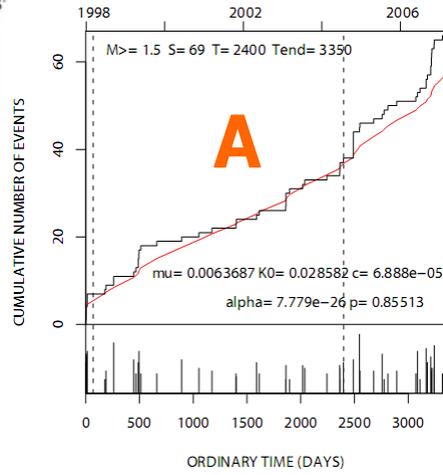
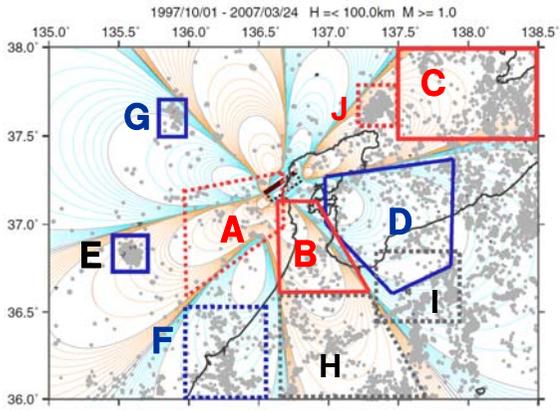


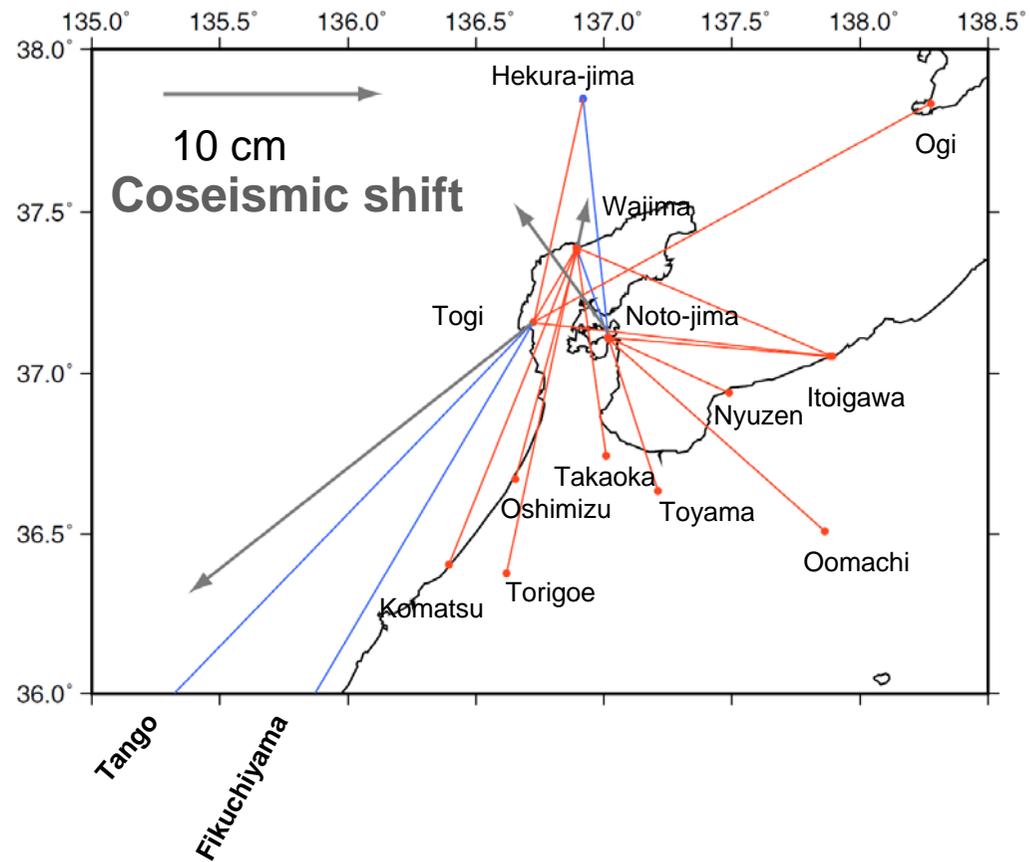
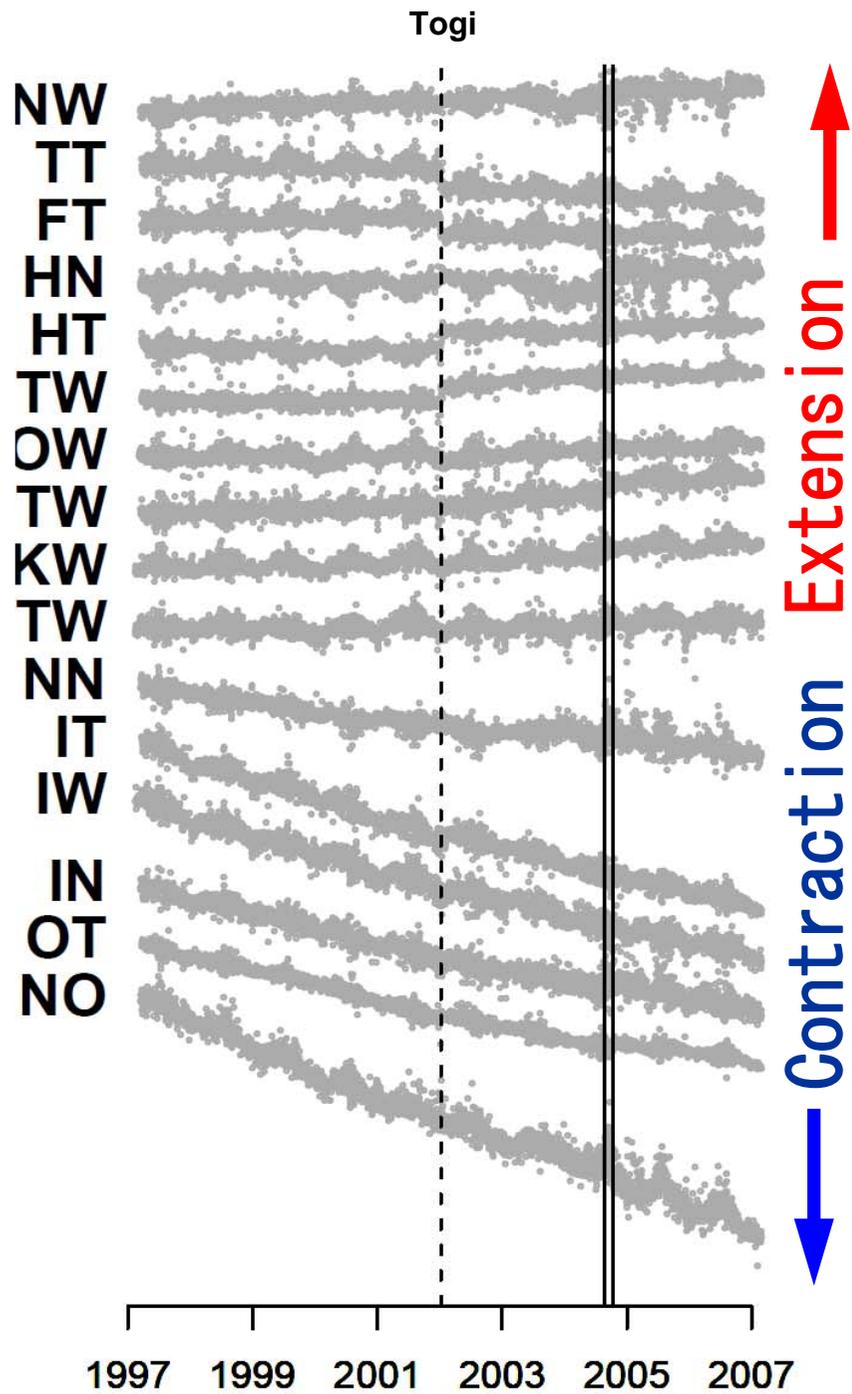
# 2007 Noto Peninsula M6.9

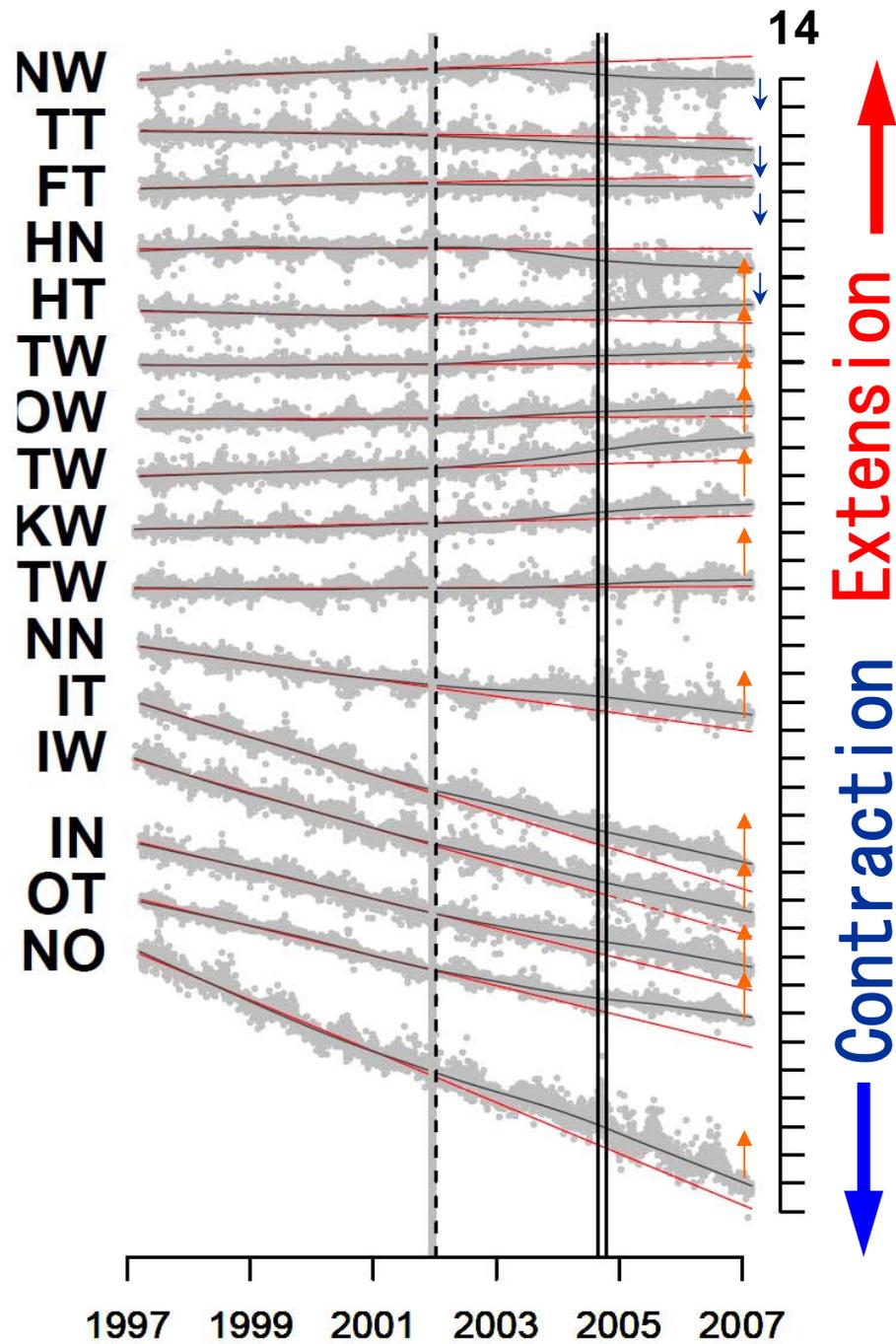
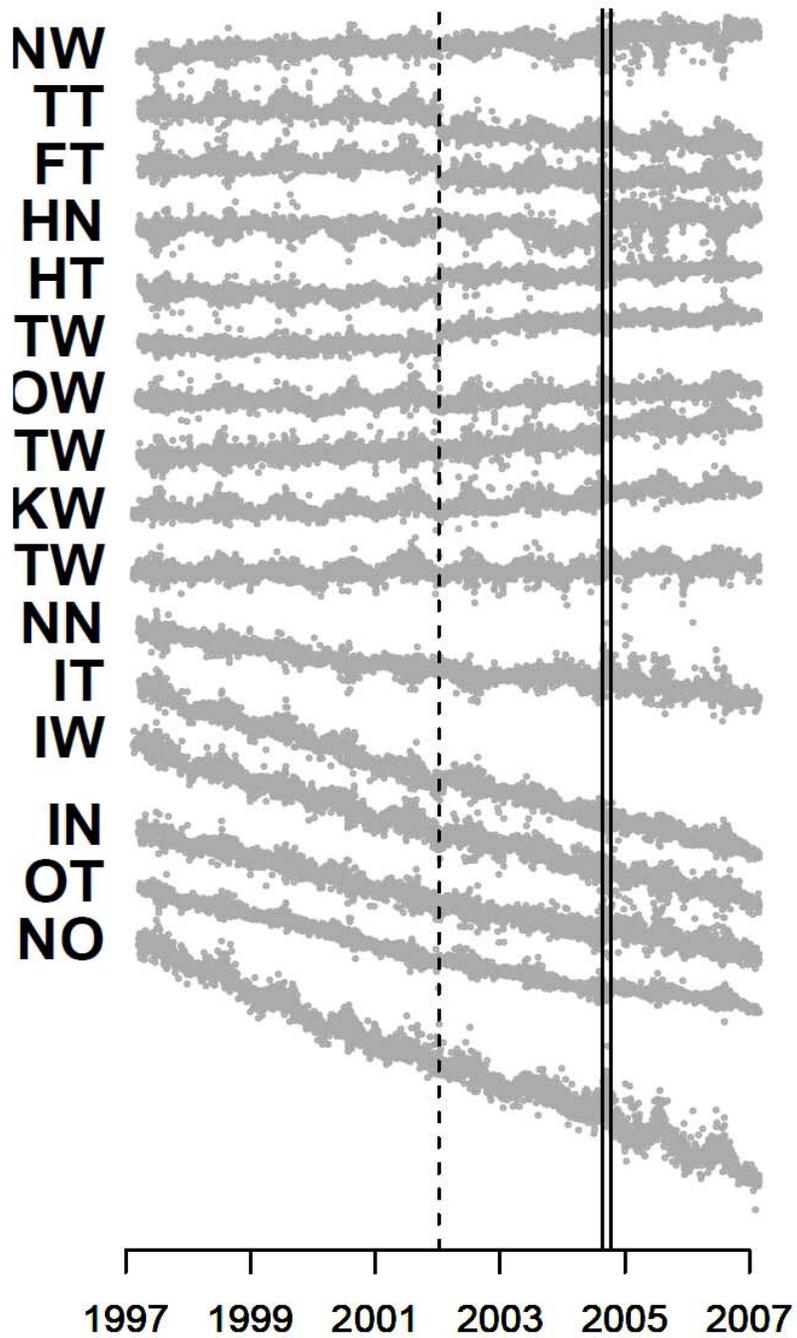


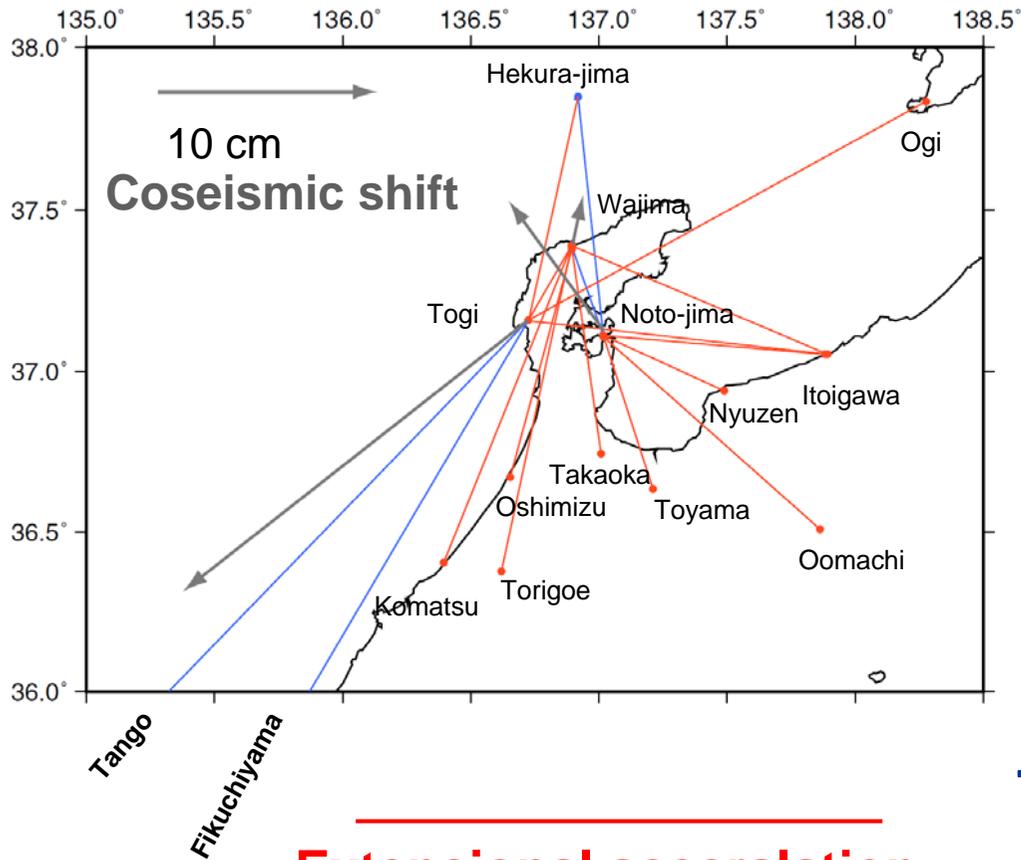








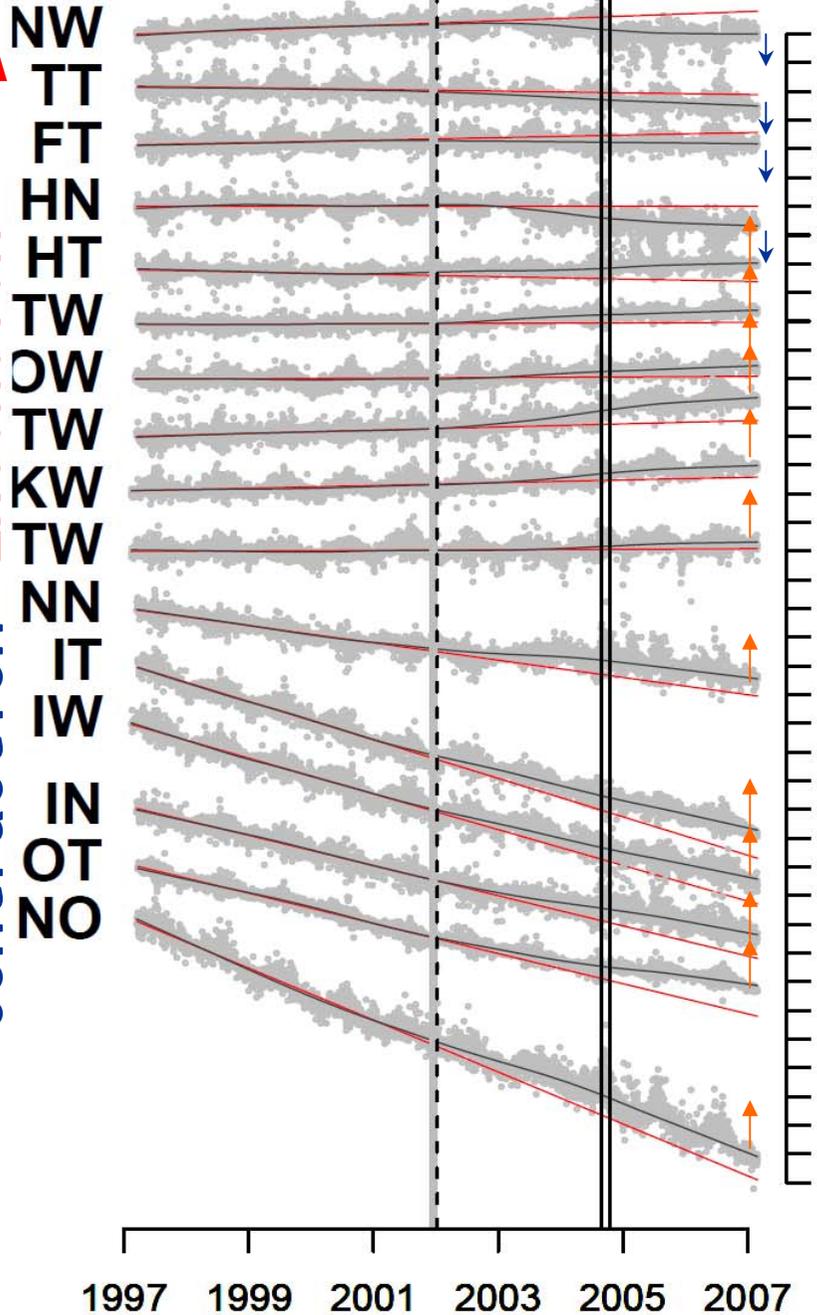




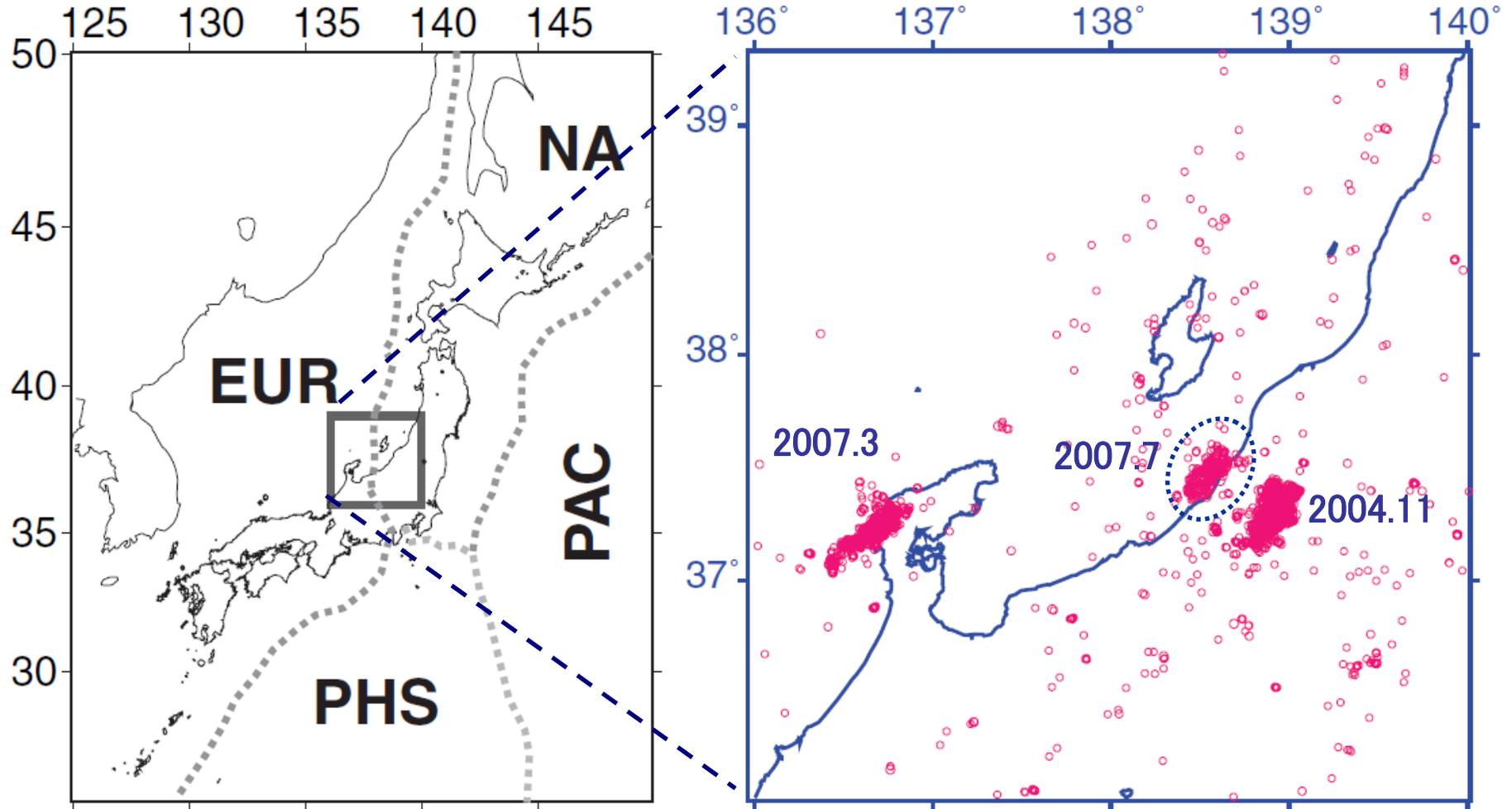
**Extensional acceleration**  
**Contractional deceleration**

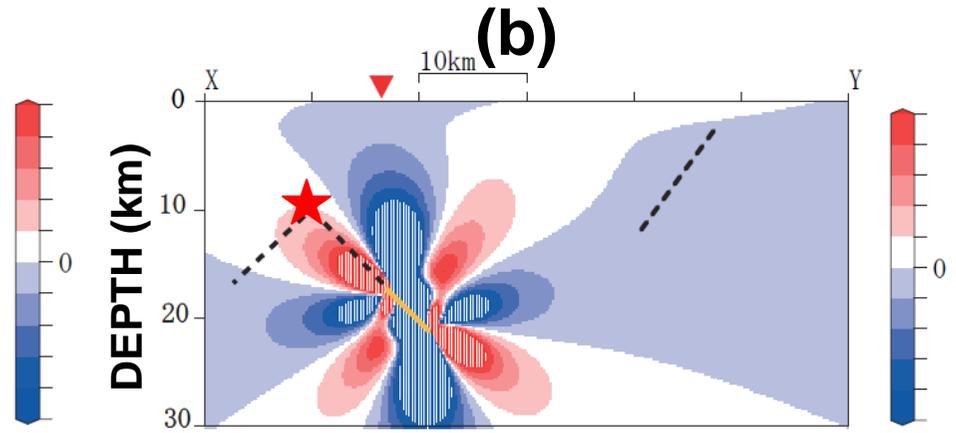
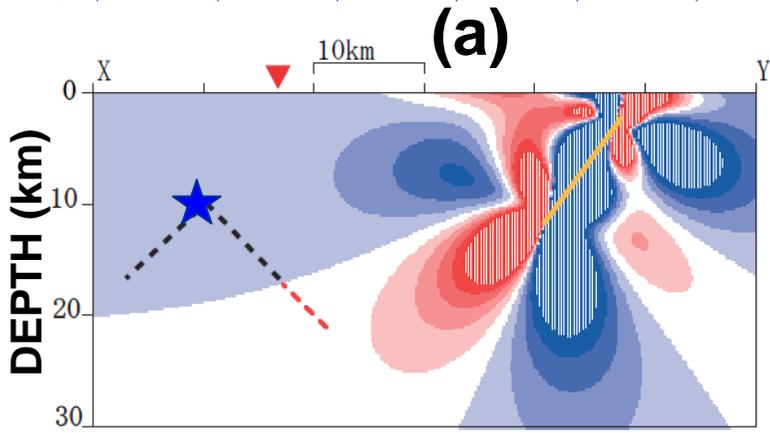
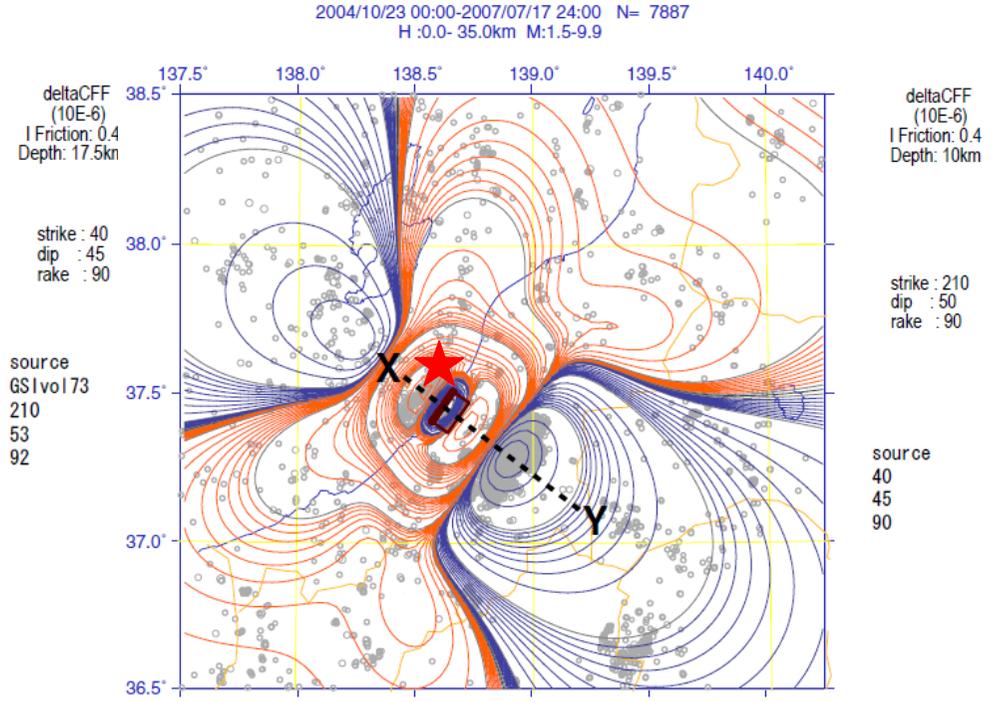
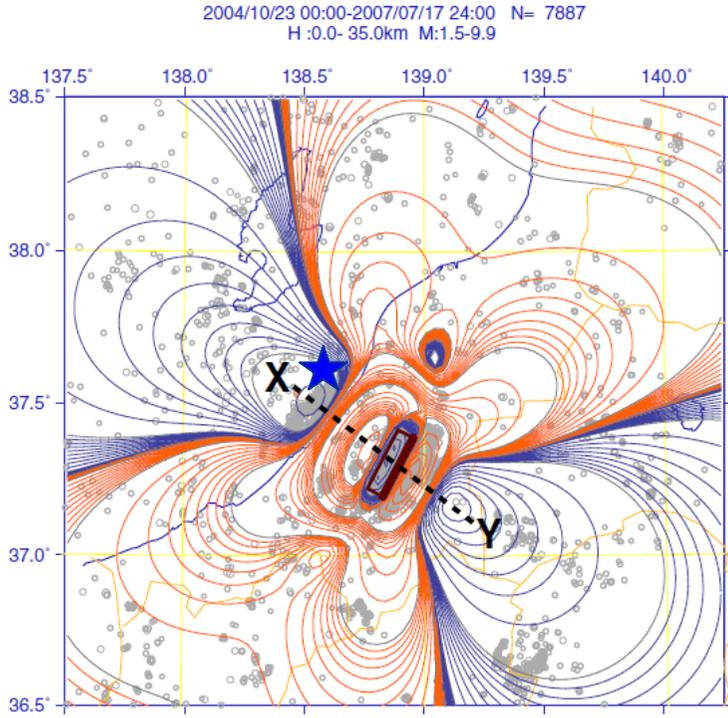
**Extensional deceleration**

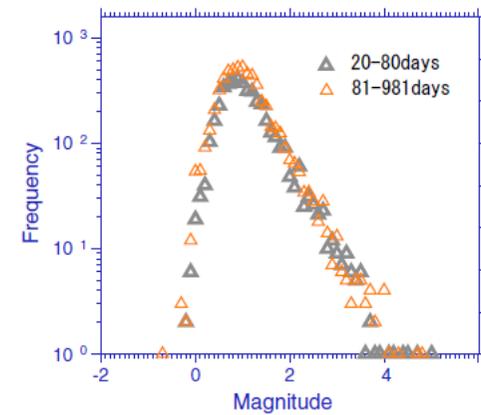
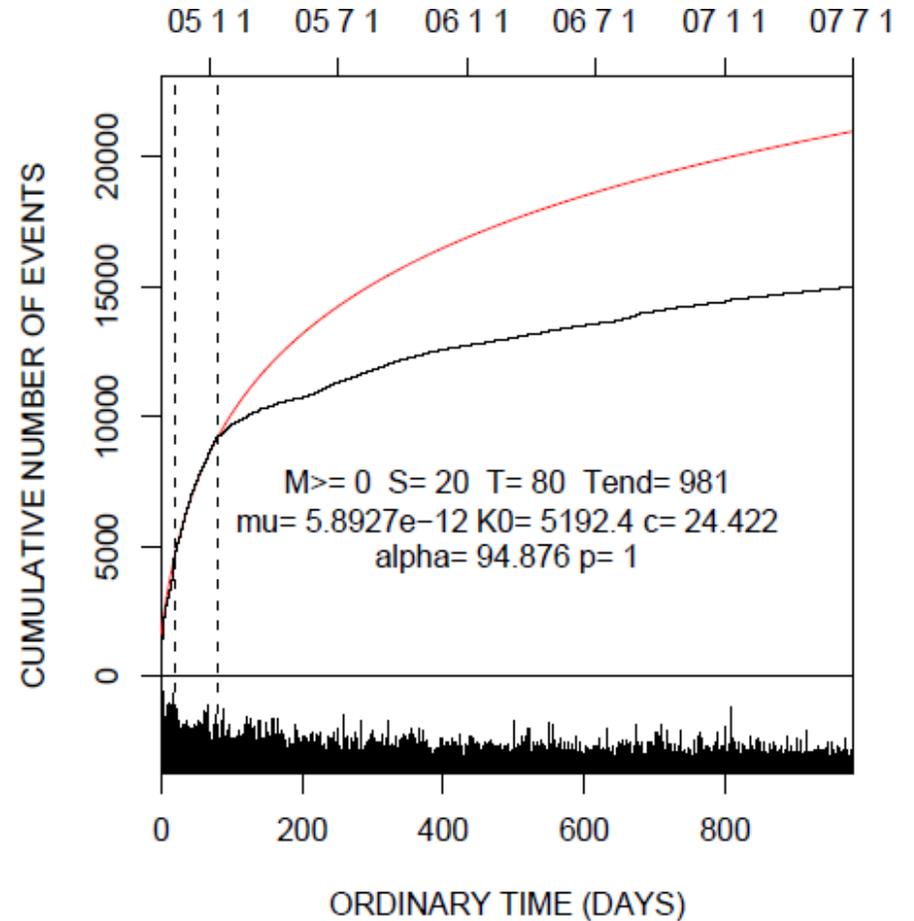
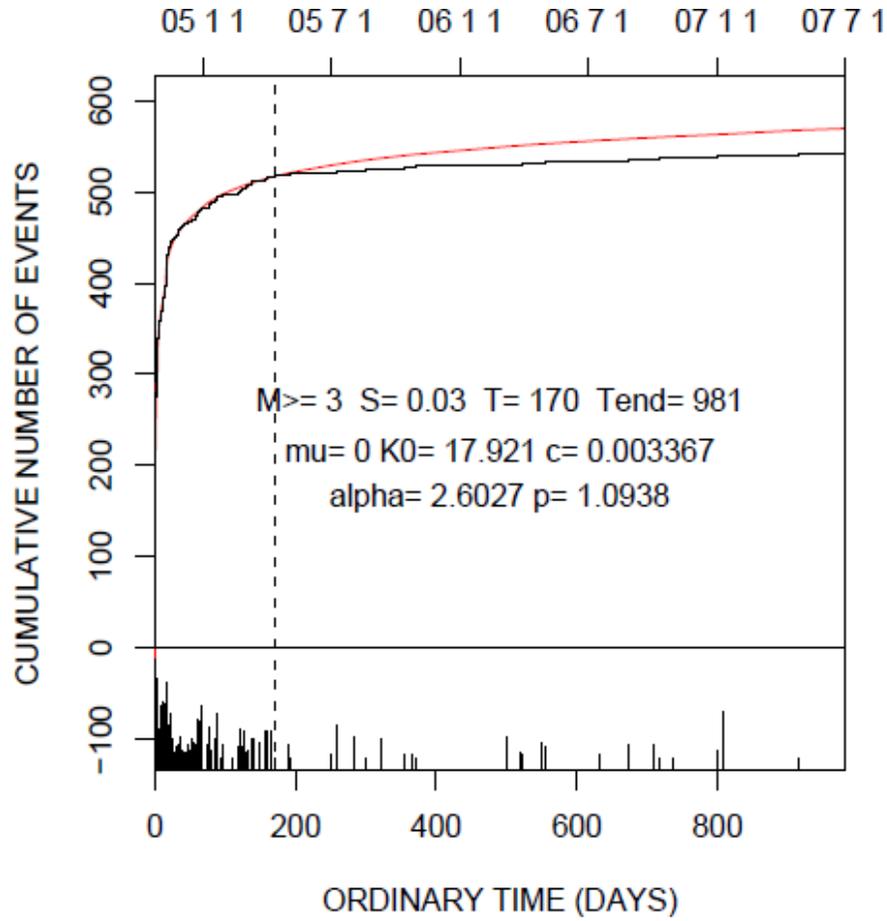
**Extension** ↑  
**Contraction** ↓

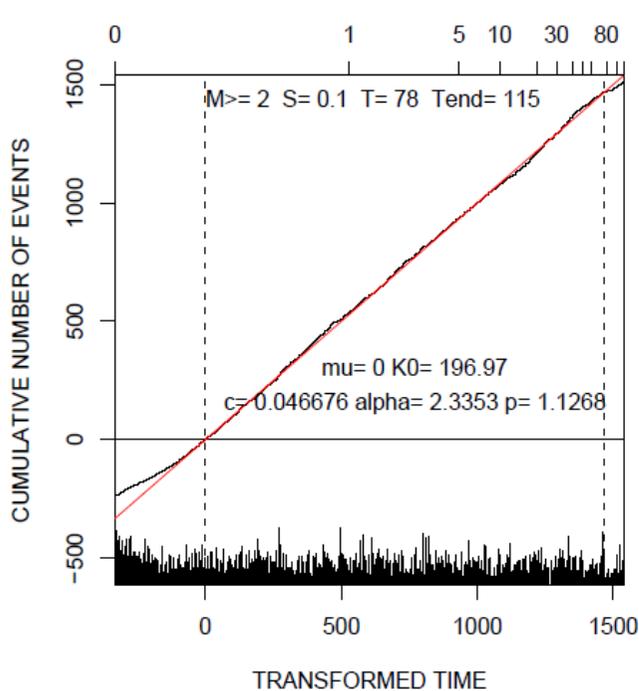
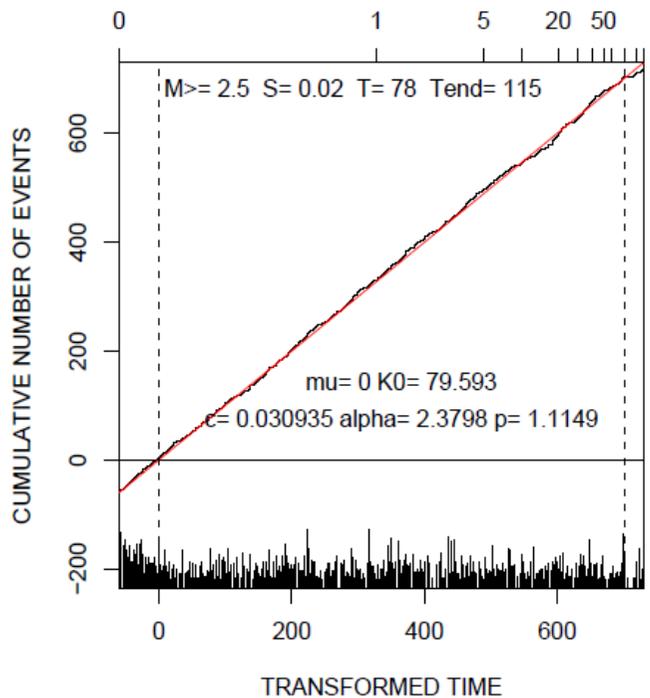
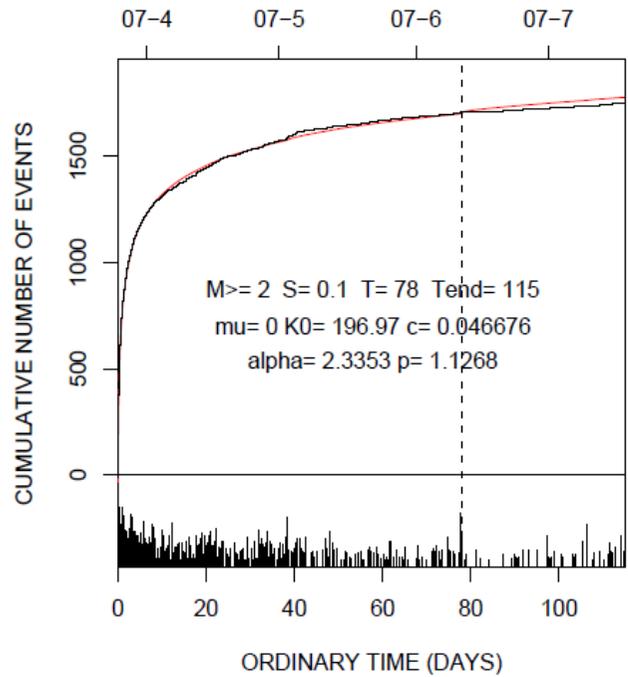
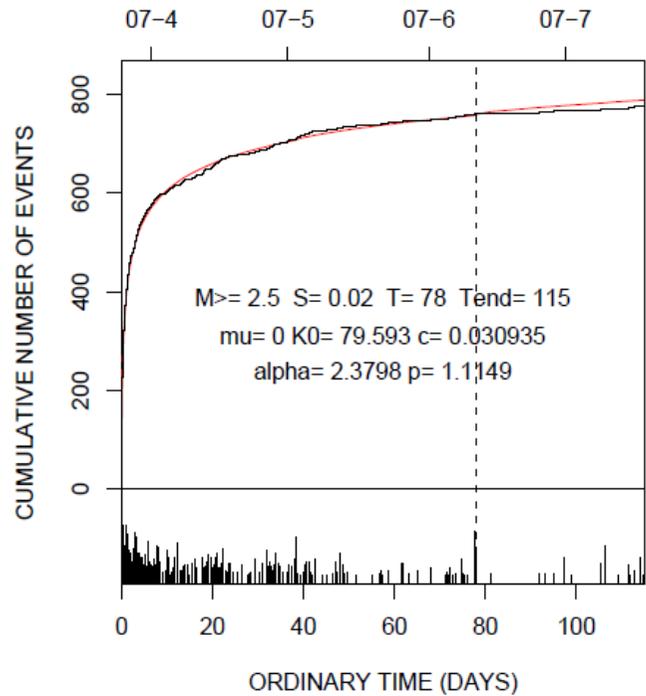


# 2007 Chuetsu-Oki M6.8



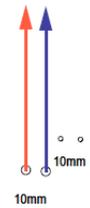
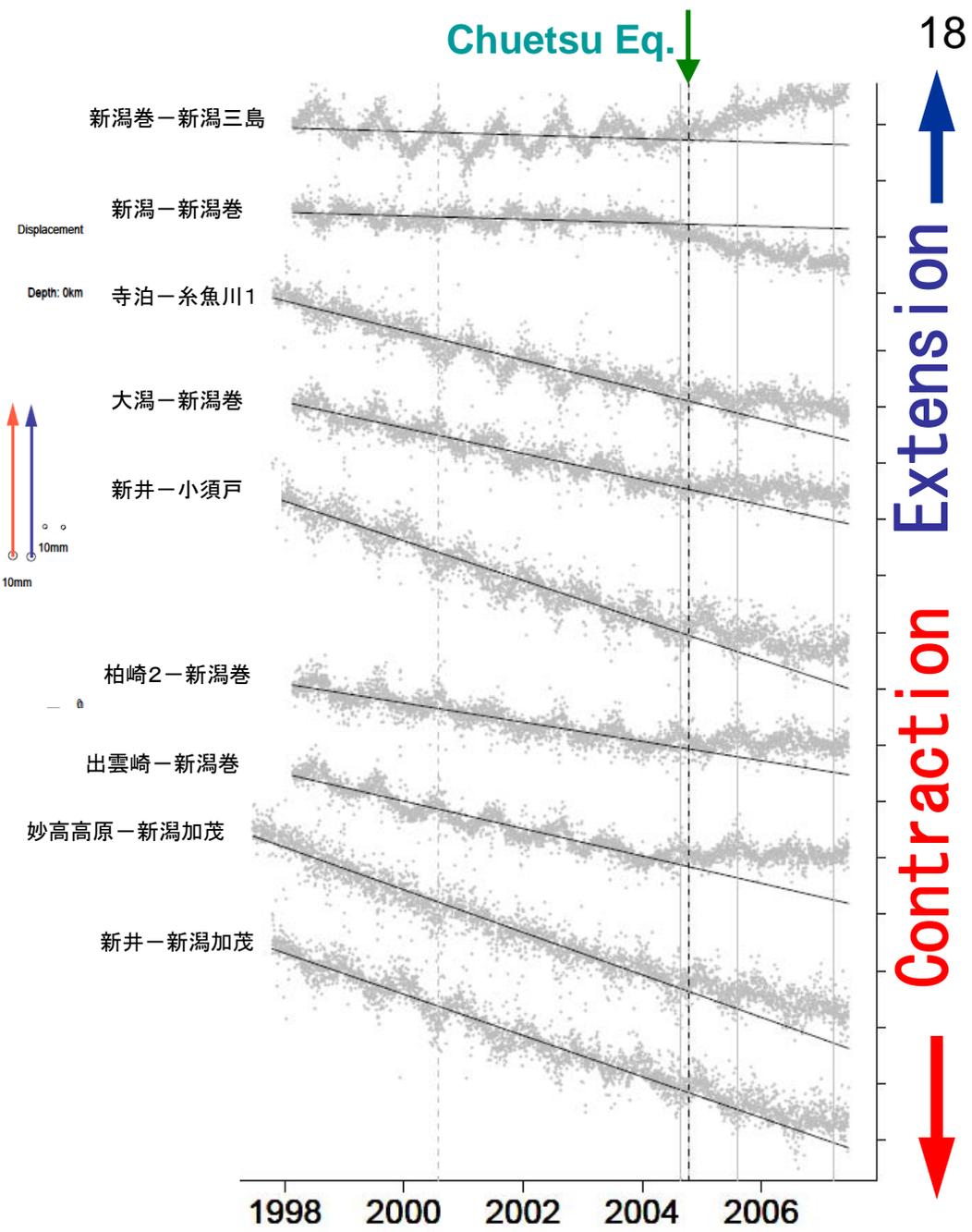
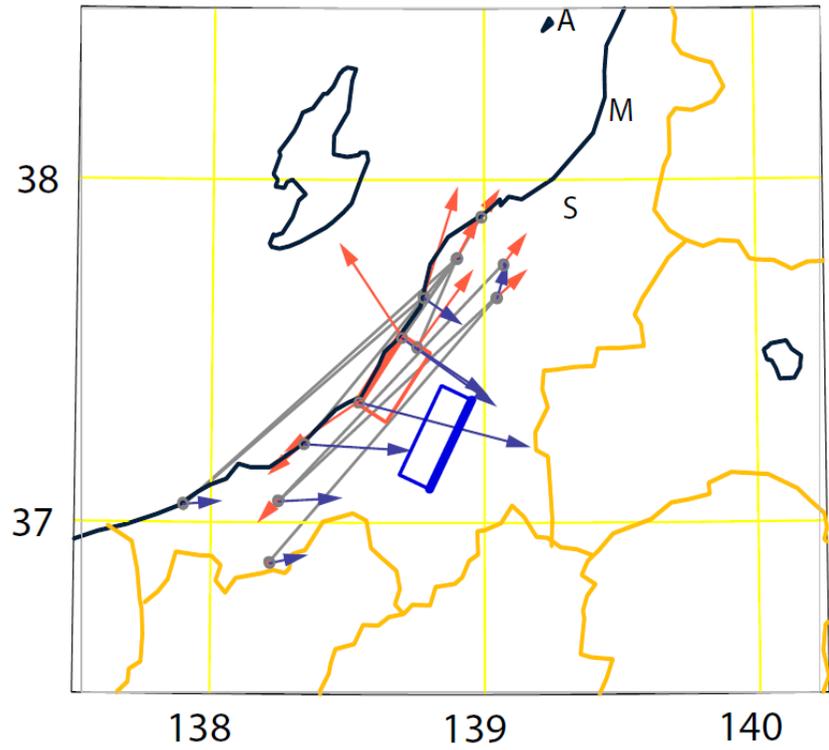






## 2007年能登半島地震の余震活動

の2007年7月20日までの、MT図と累積関数。右側図は下限マグニチュードがM2.5で、左側はM2.0の余震。赤い累積曲線は、本震後一定の時間(右側図は0.02日で左側図は0.1日)から変化点(78日)までの余震データにETASモデルをあてはめて残りの期間を予測した、理論的累積曲線。横軸は、上図が通常的时间推移で、下図がETASモデルによる変換時間。



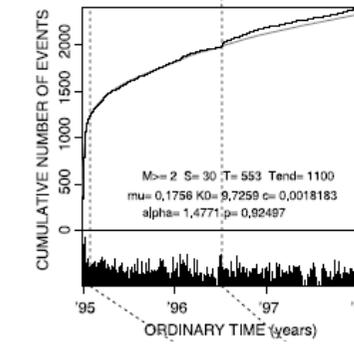
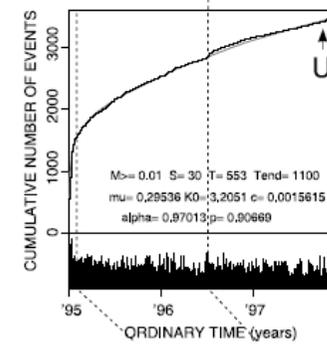
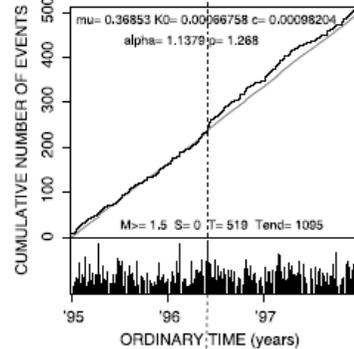
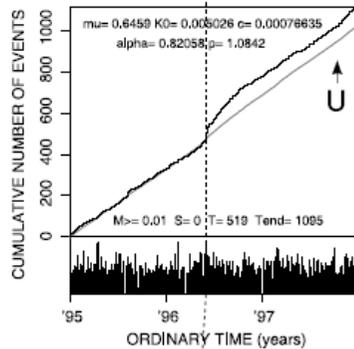
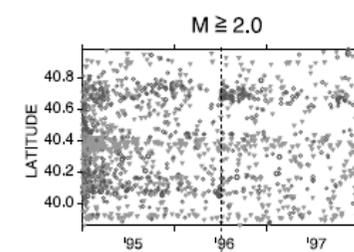
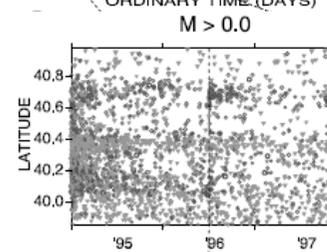
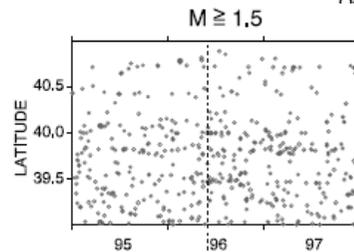
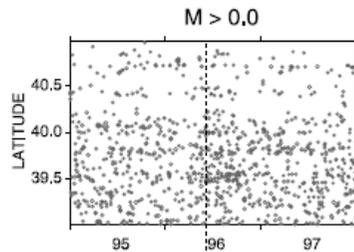
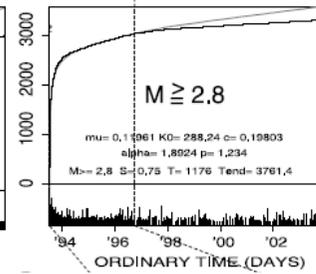
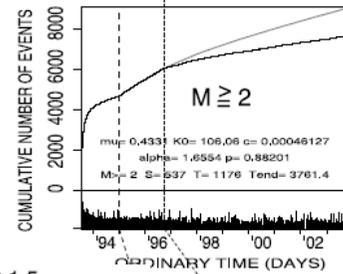
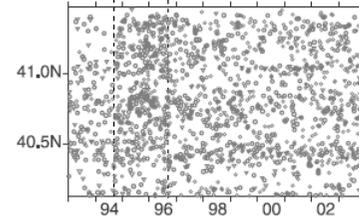
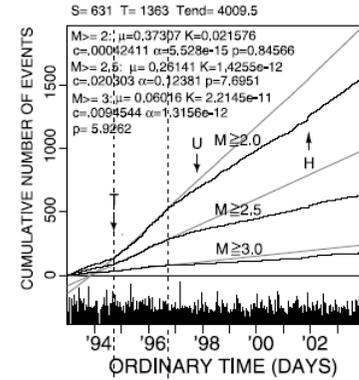
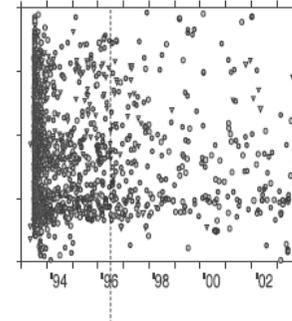
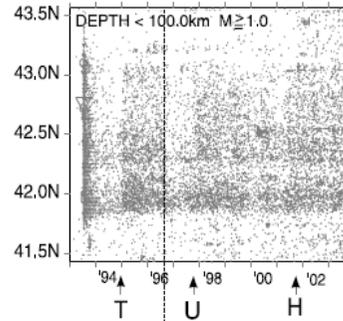
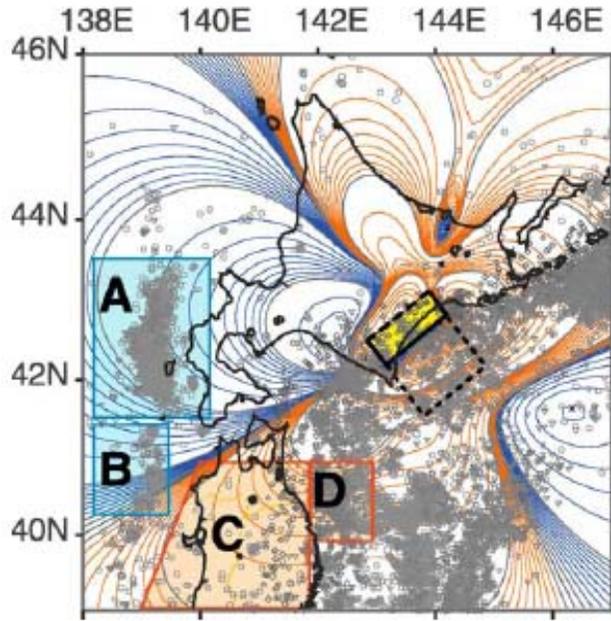
# Summary

- The ETAS model summarizes the seismicity due to triggering effect within the contiguous hierarchical complex faults.
- Deviations of the seismic activities from the predicted rate by the ETAS model are useful to detect regional stress changes.
- The respective deviations are explained by the changes in Coulomb failure stresses that are caused by seismic or aseismic slips.
- These are further supported by transient crustal movement around the source due to the aseismic slips preceding the ruptures.

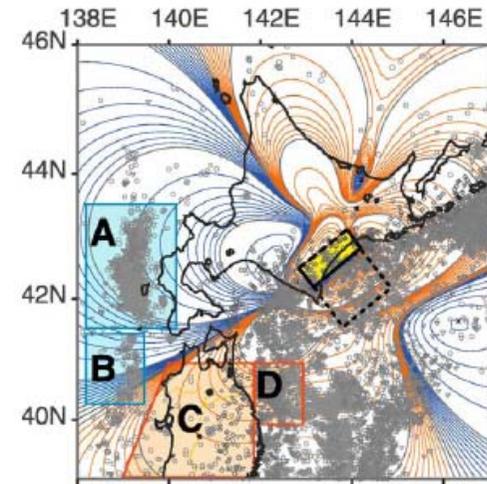
Software and manuals are available:  
Search “SASeis2006” by Google

Ogata (2005, *JGR*)

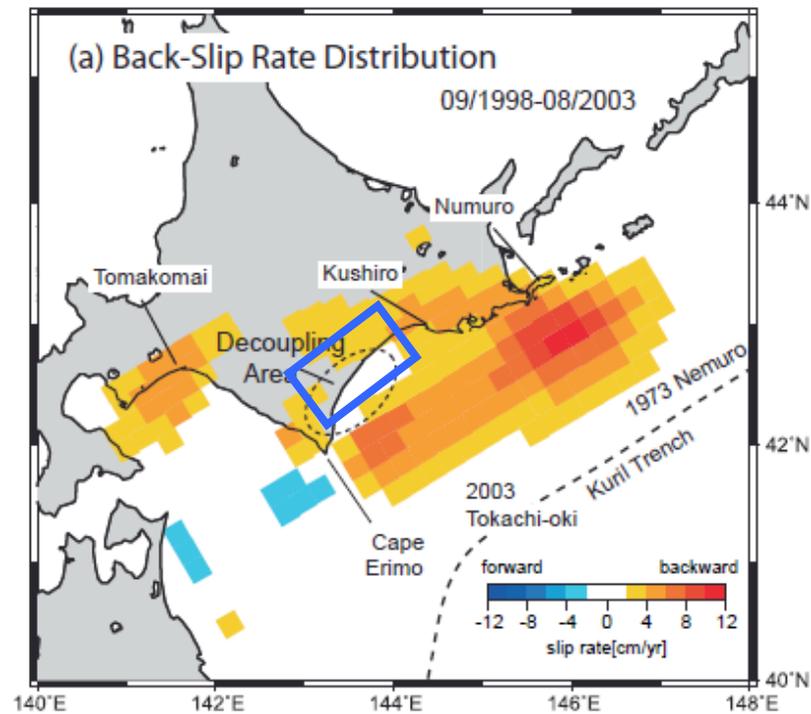
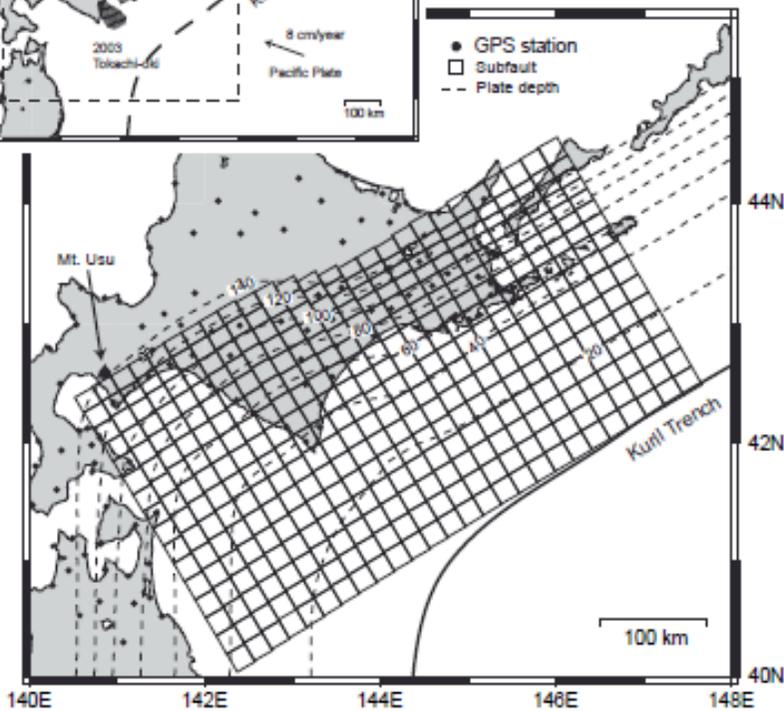
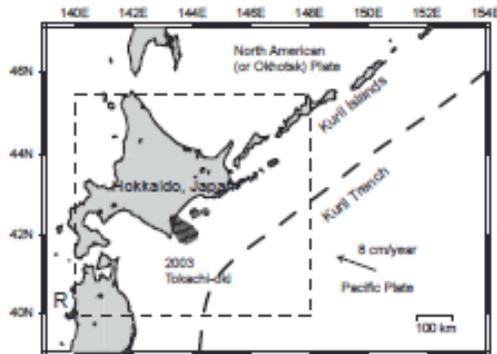
# Period of anomaly: mid1996 ~ mid2003



# Back slip rate estimate during 1998–2003AUG by inversion of GPS network

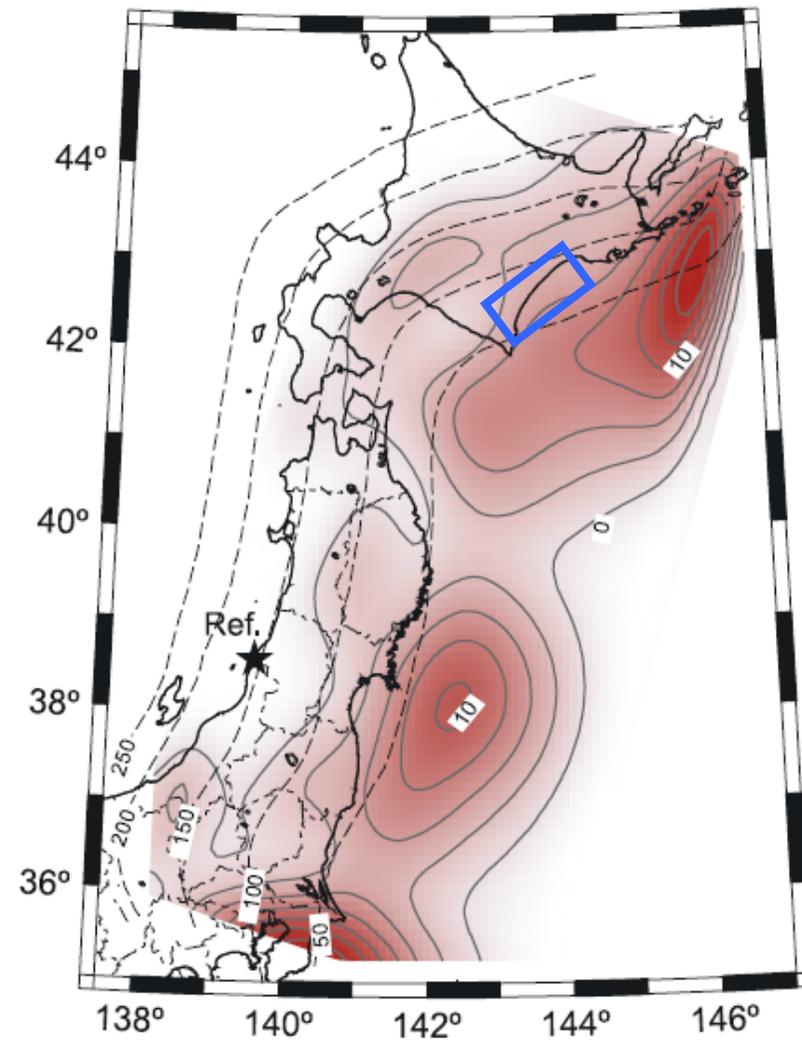
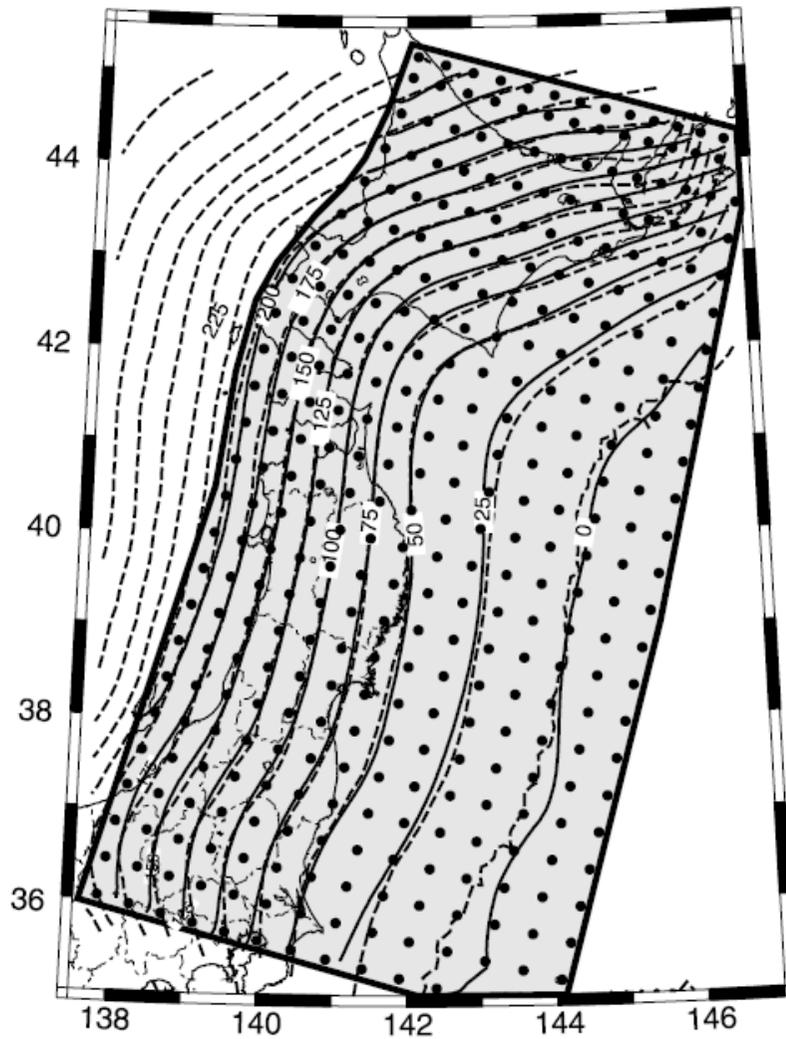


**Baba & Hori (2007)**



**Suwa et al. (2006, JGR)**

**Back slip rate 1997–2001**



Thank you very much for  
listening

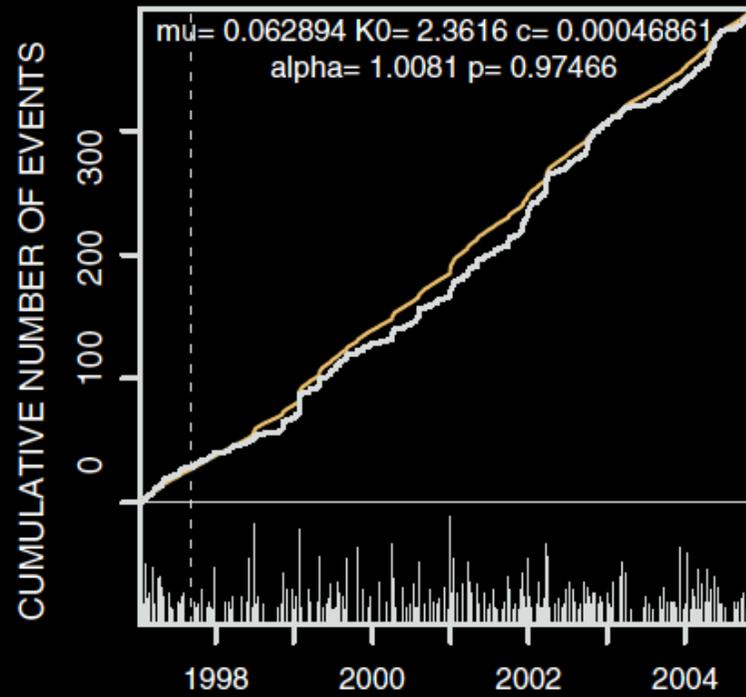


**Software and manuals are available at**

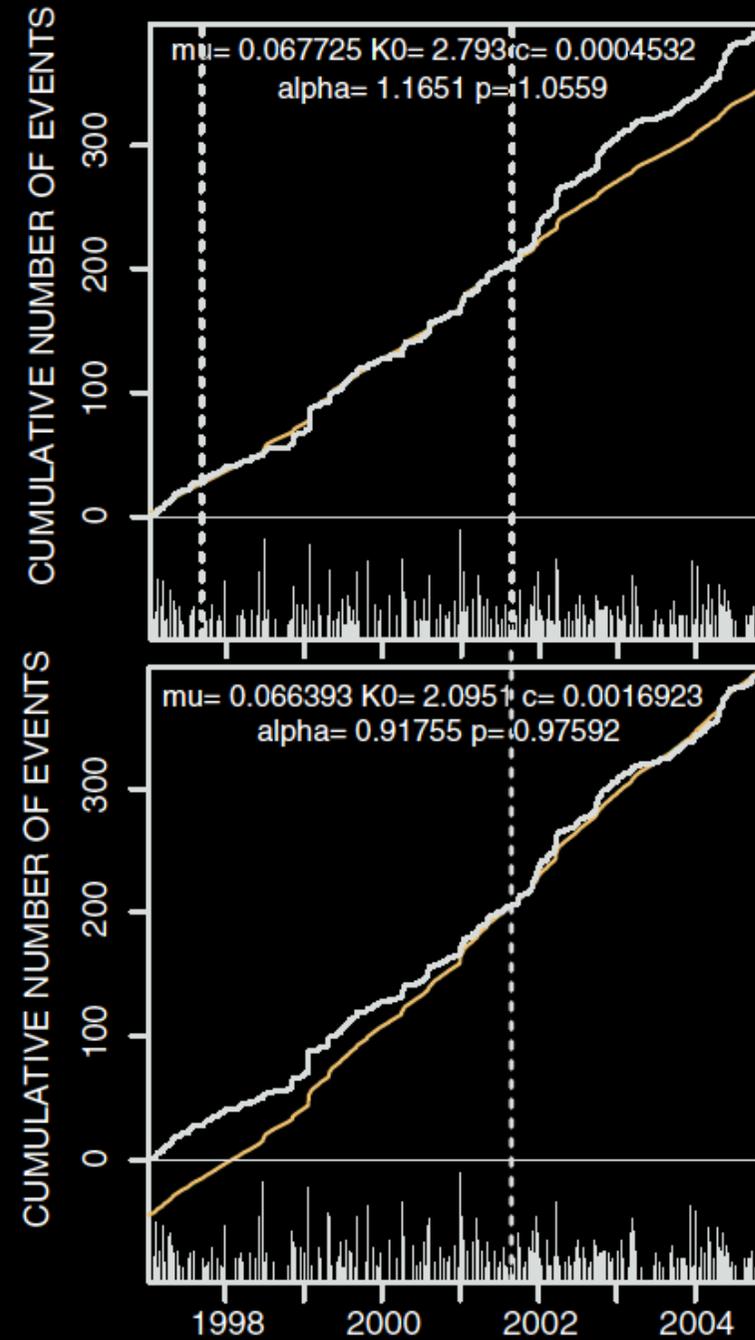
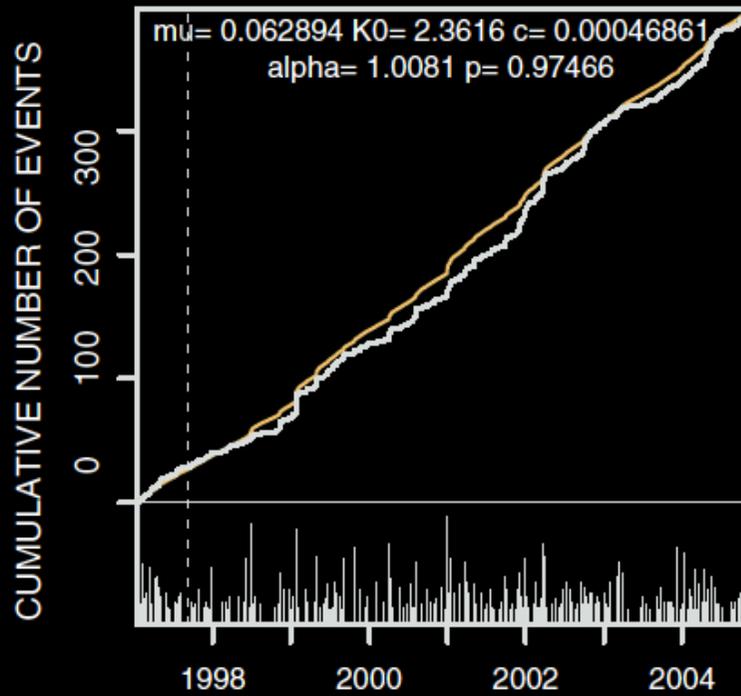
<http://www.ism.ac.jp/~ogata/Ssg/softwarese.html>

Search “SASeis2006” by Google

## No-Change



## No-Change



# Coseismic Slip during the first 60 seconds

## Yagi (2004) EPS

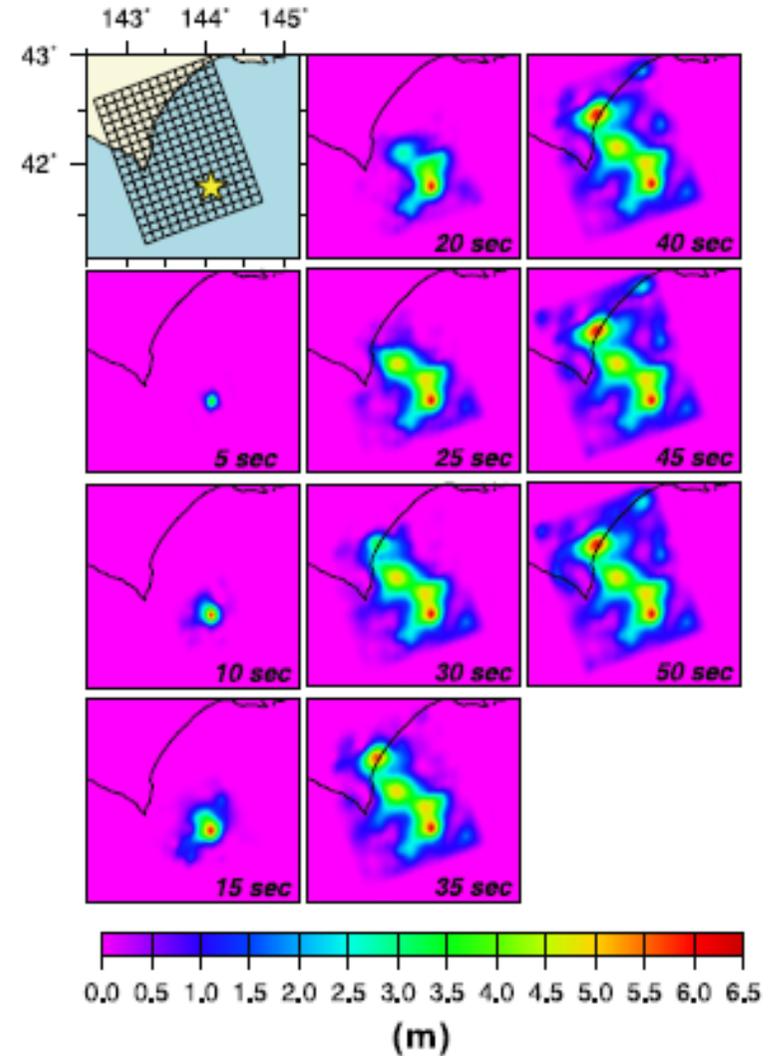
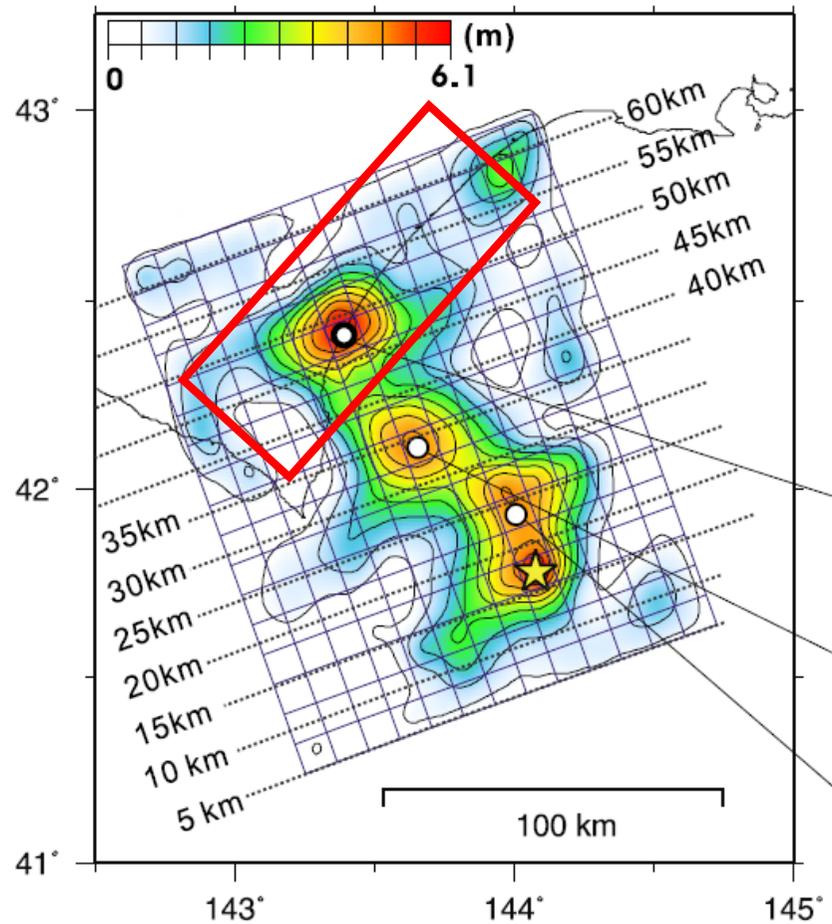
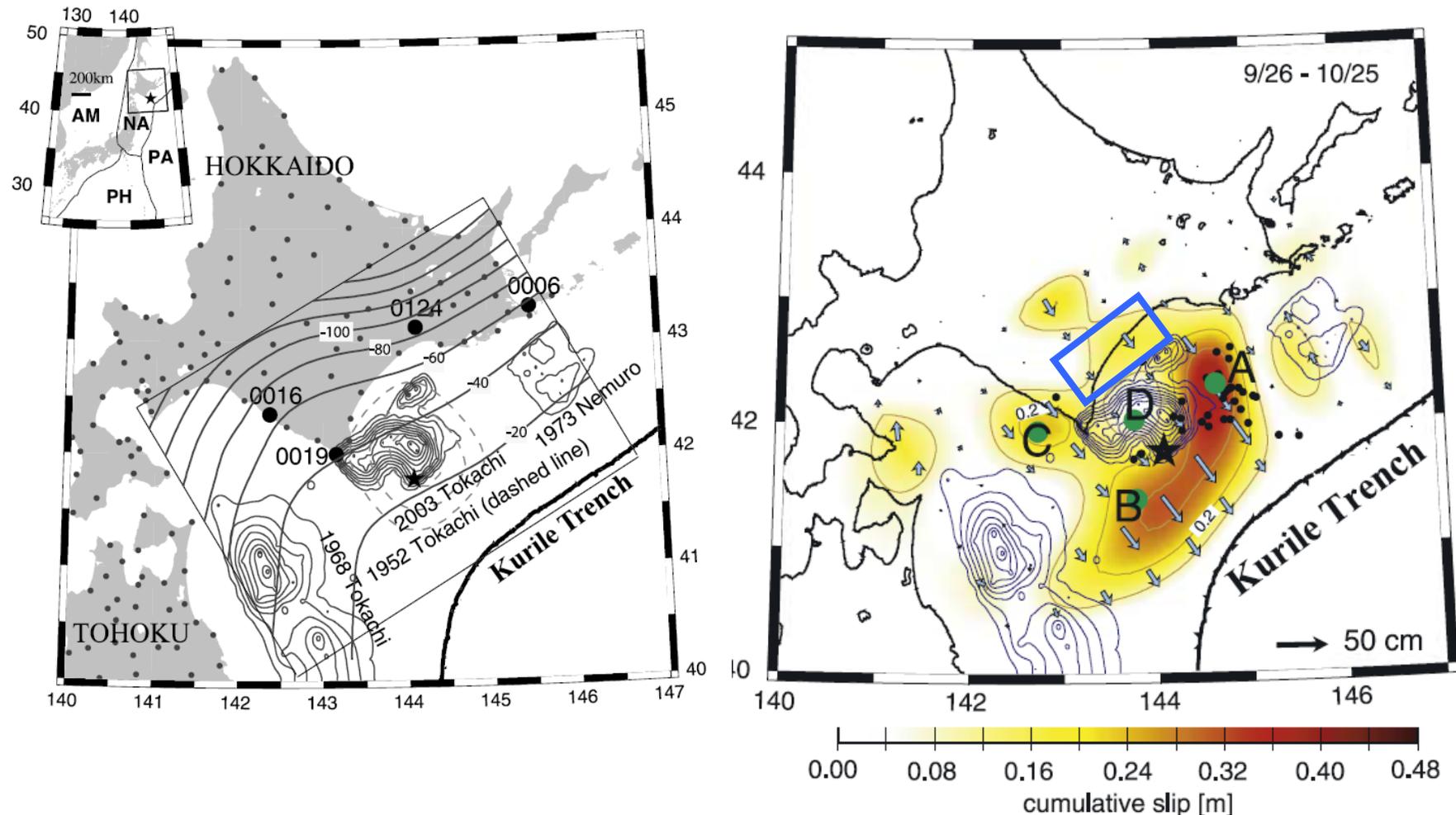


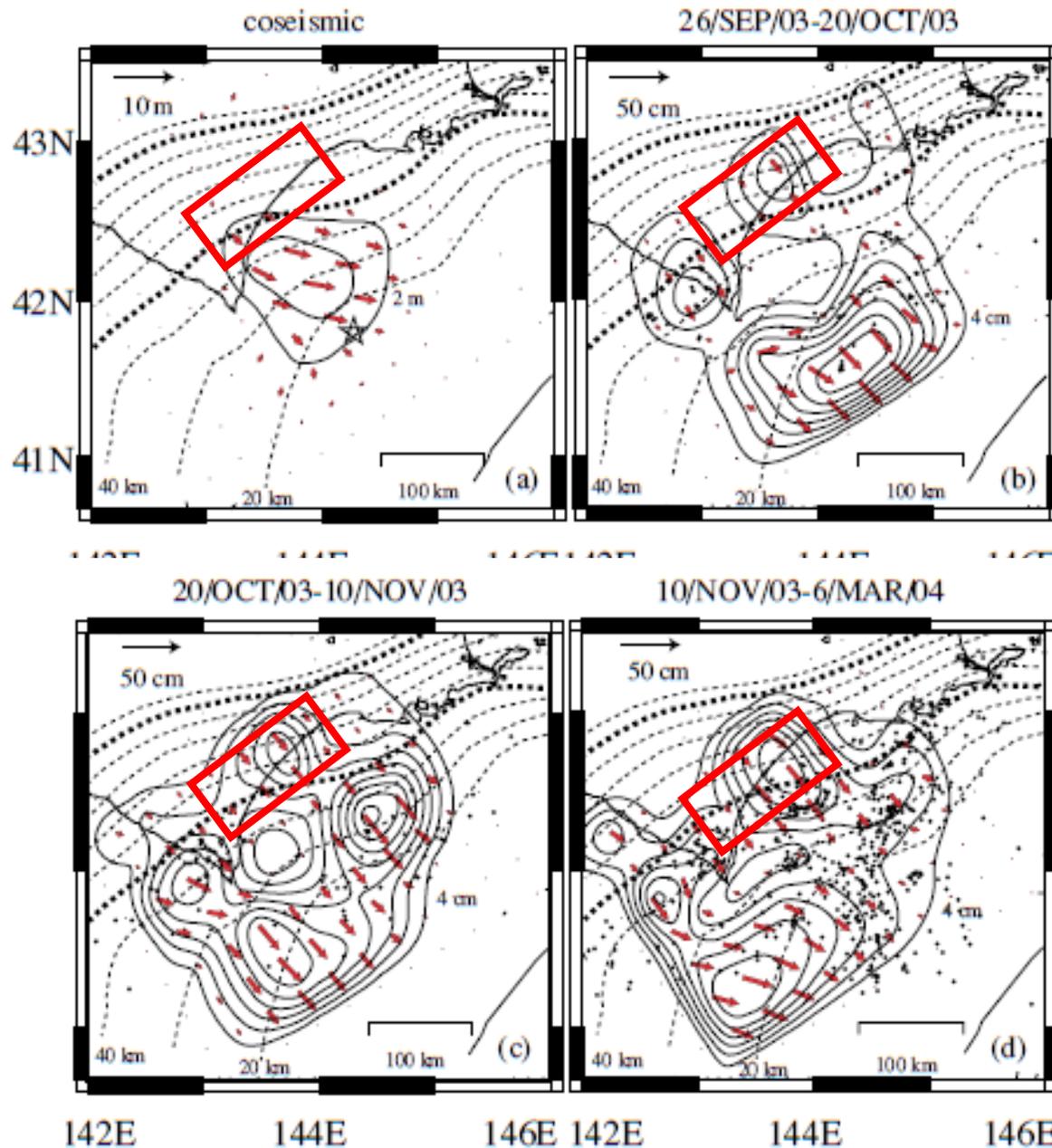
Fig. 6. Snapshots of surface projection of the dislocation at every 5 sec. The star indicates the epicenter.

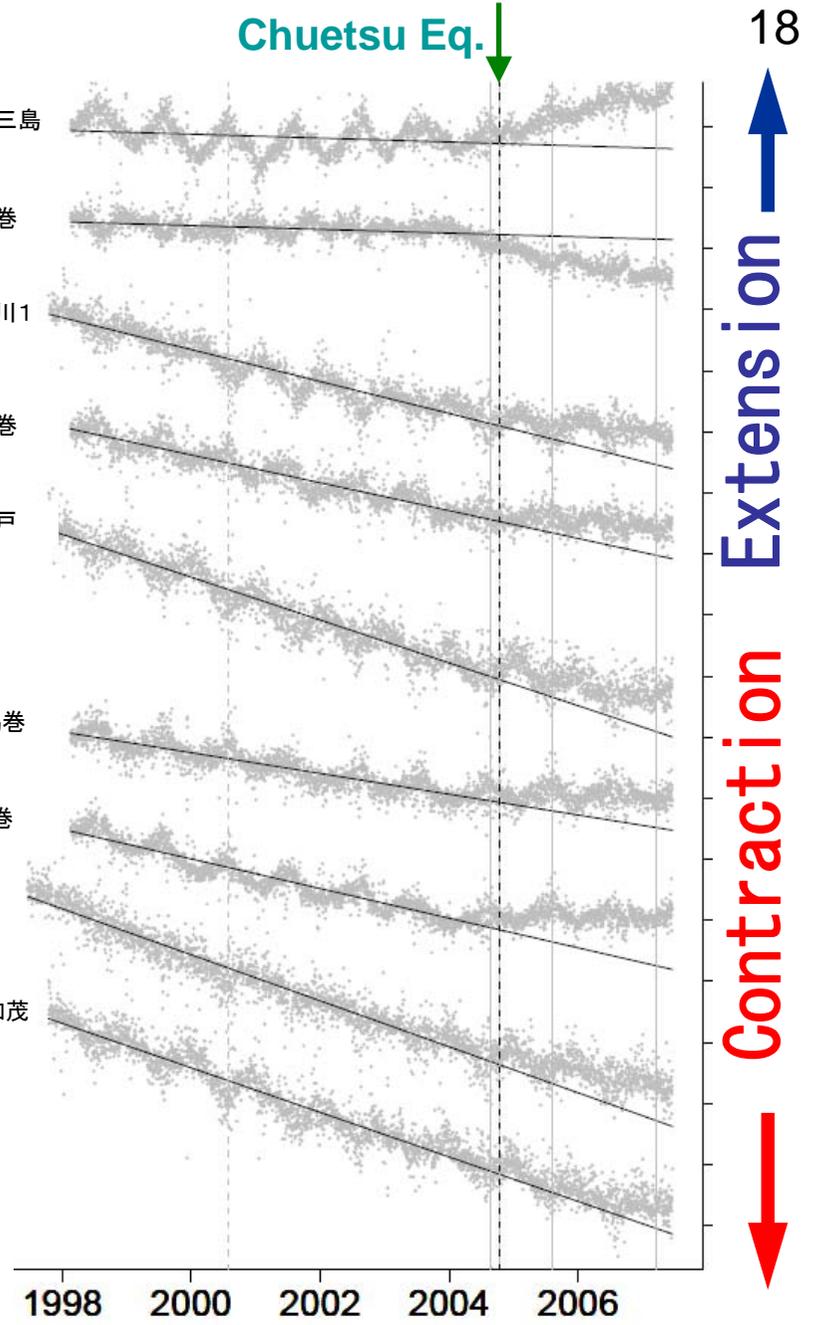
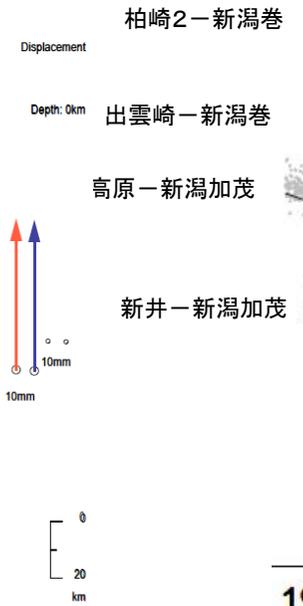
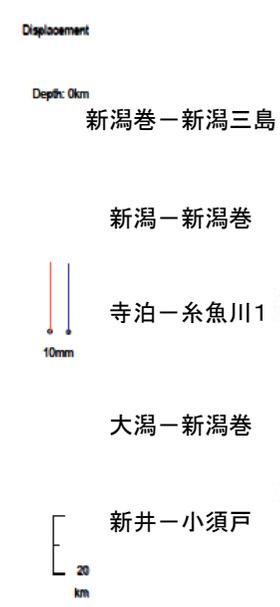
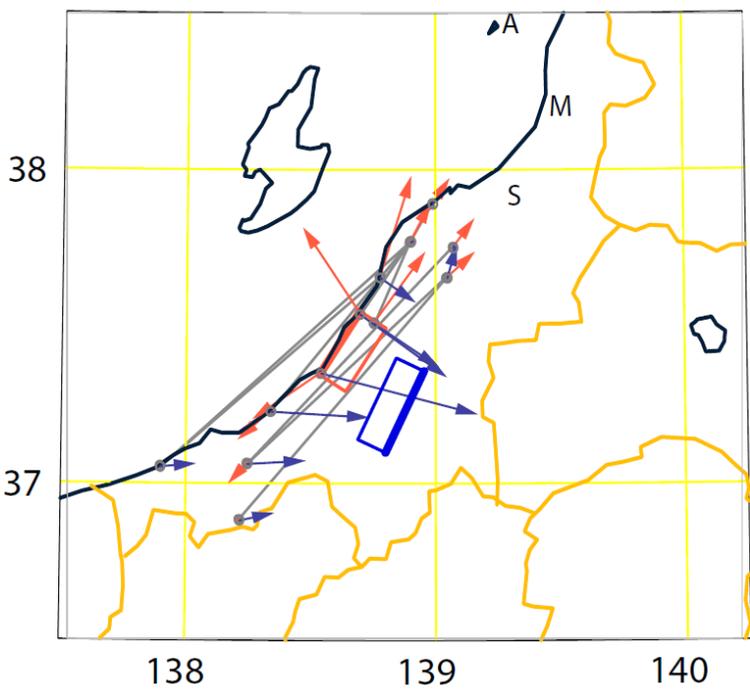
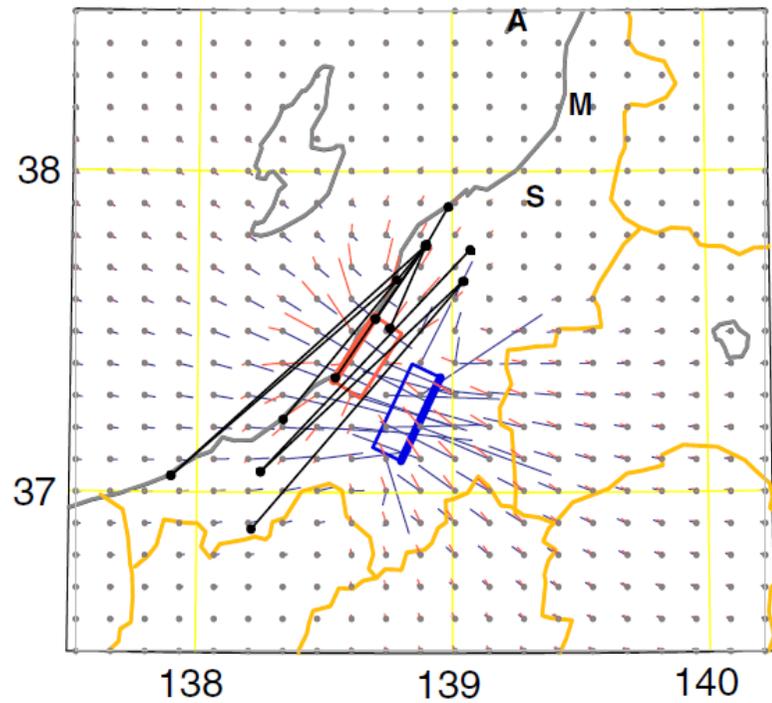
# Cumulative Slip for 30days From the mainshock

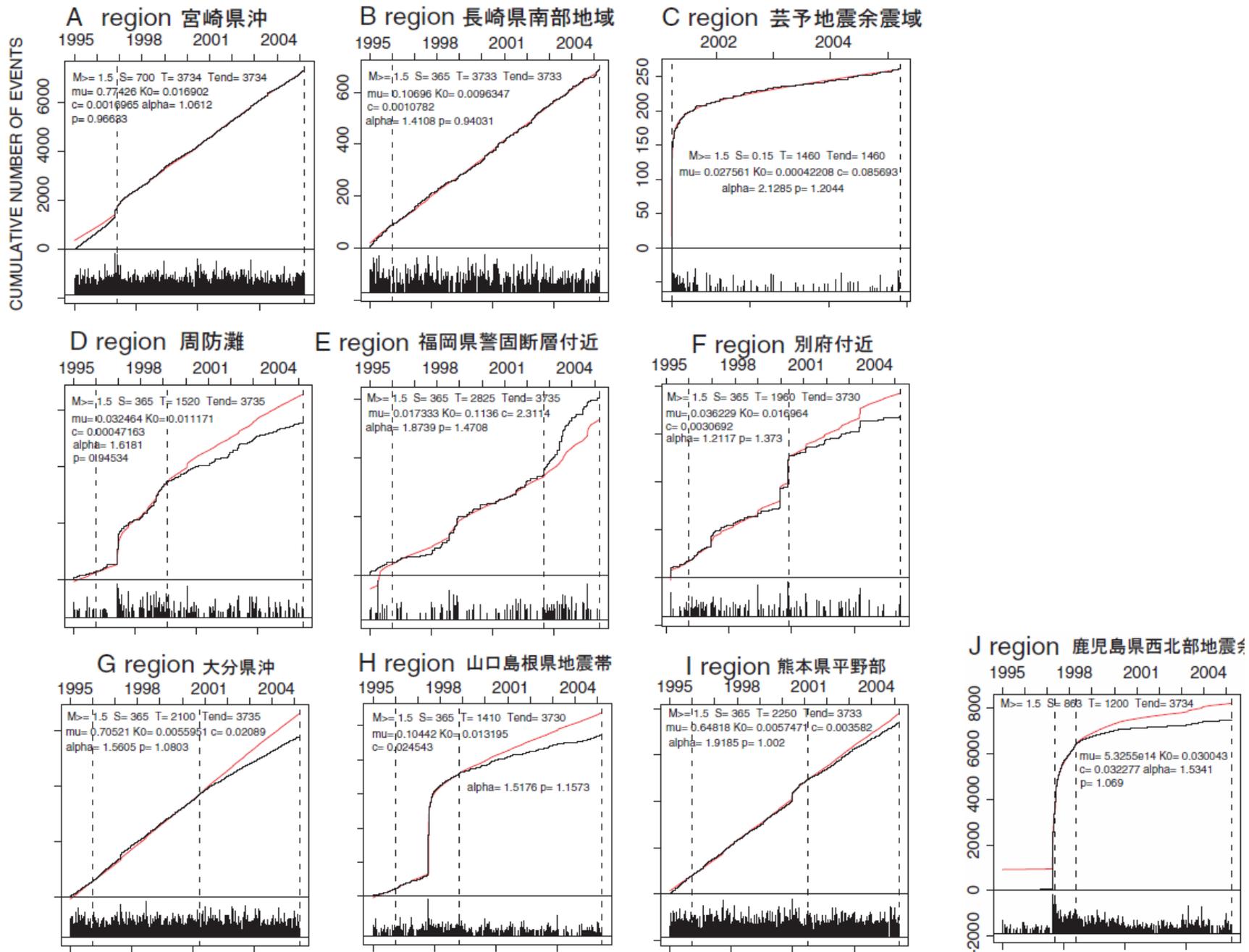
Miyazaki et al. (2004) GRL

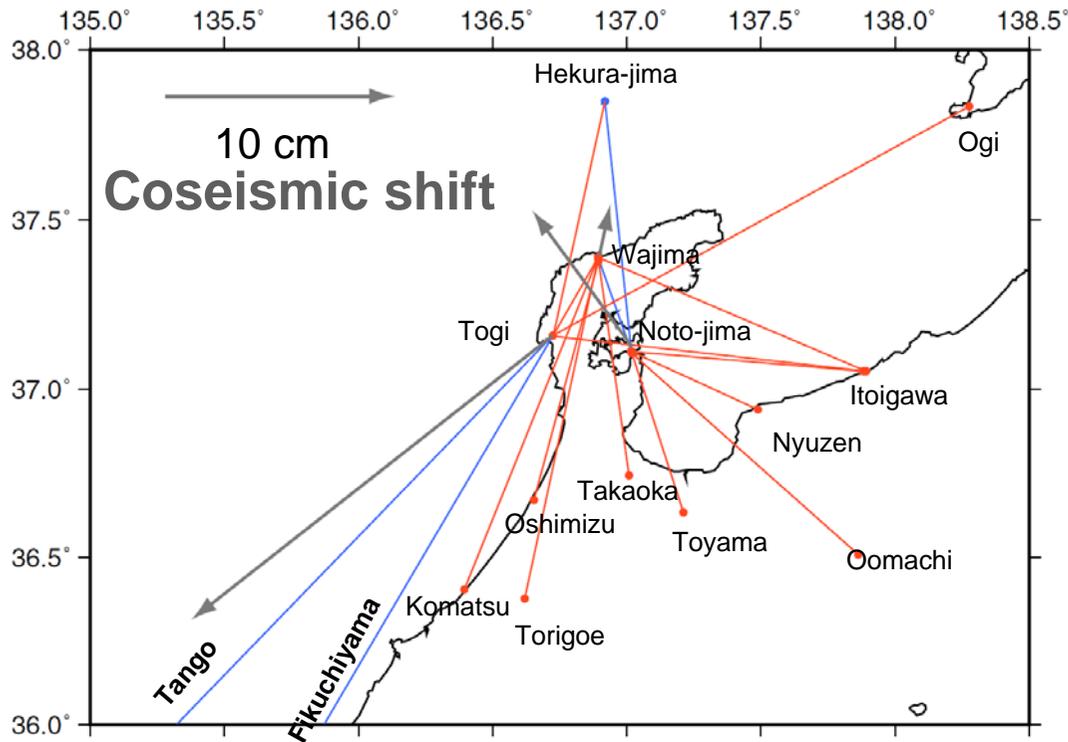


# Co- & Post-seismic slip Ozawa et al. (2004) EPS









**Extensional accerlation**  
**Contractional decerlation**

**Extensional  
decerlation**

**Contraction** ↑  
**Extension** ↓

