

Experiences in rockburst prediction

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ISS International

Rockbursts can be traumatic phenomena



**Objectives of seismic
monitoring in rockburst prone
mines**



■ **Potential rockburst detection**

■ **Long term hazard assessment**

- ◆ **Back analysis**
- ◆ **Calibrated models**

■ **Medium term hazard assessment**

- ◆ **Spatial detection of potential instabilities**
 - **Essentially using the asperity model for spatial prediction**
- ◆ **Monthly hazard ratings**

■ **Short term hazard assessment**

- ◆ **Time history analyses for the detection of unstable processes**

Definitions



- Basic seismic source parameters
- Derived parameters

Definitions



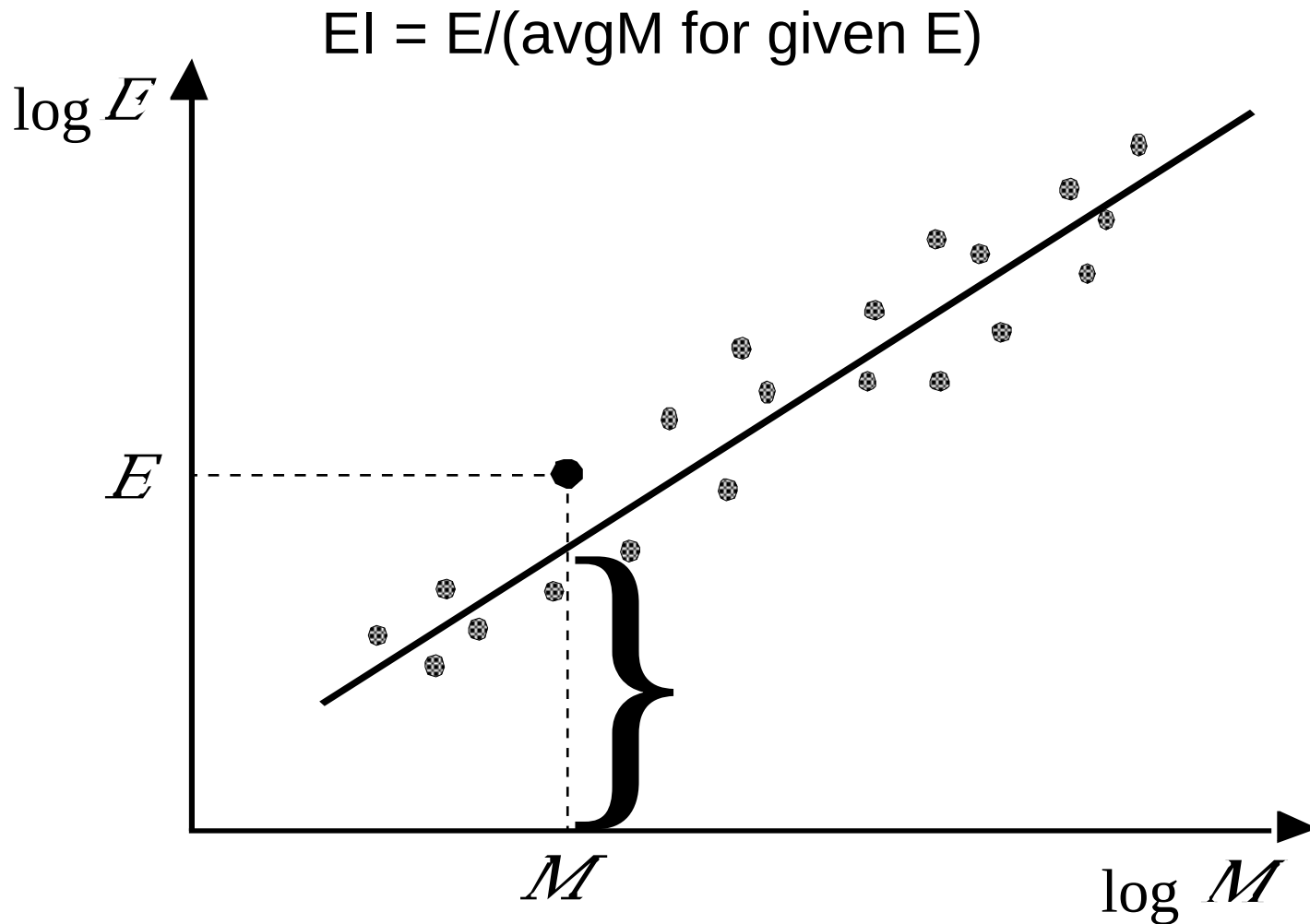
■ Basic seismic source parameters

- ◆ Time (t_0)
- ◆ Space (x, y, z)
- ◆ Seismic potency – i.e. moment/ G
- ◆ Radiated seismic energy

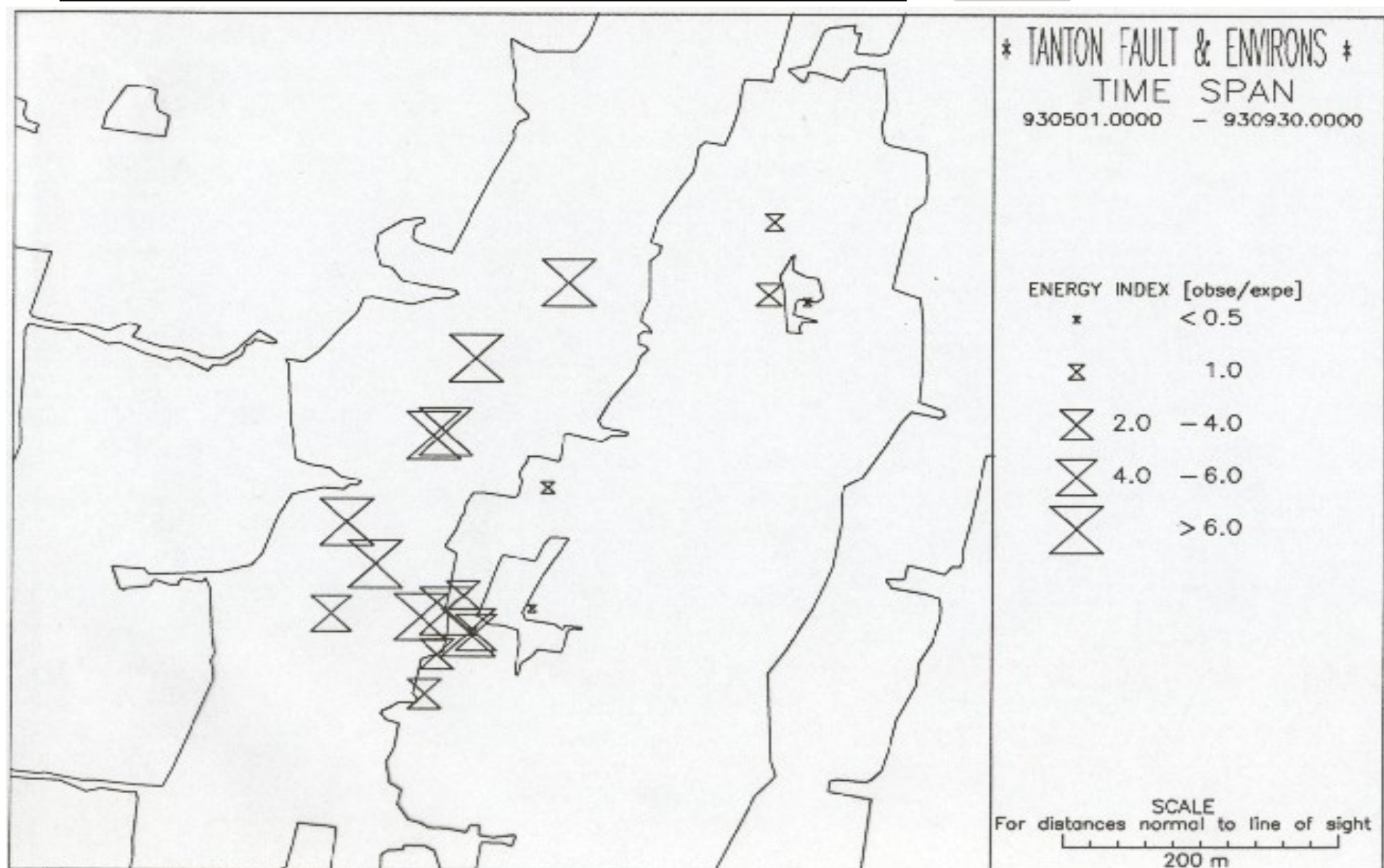
■ Derived parameters

- ◆ Energy Index
- ◆ Apparent volume
- ◆ Seismic Schmidt number

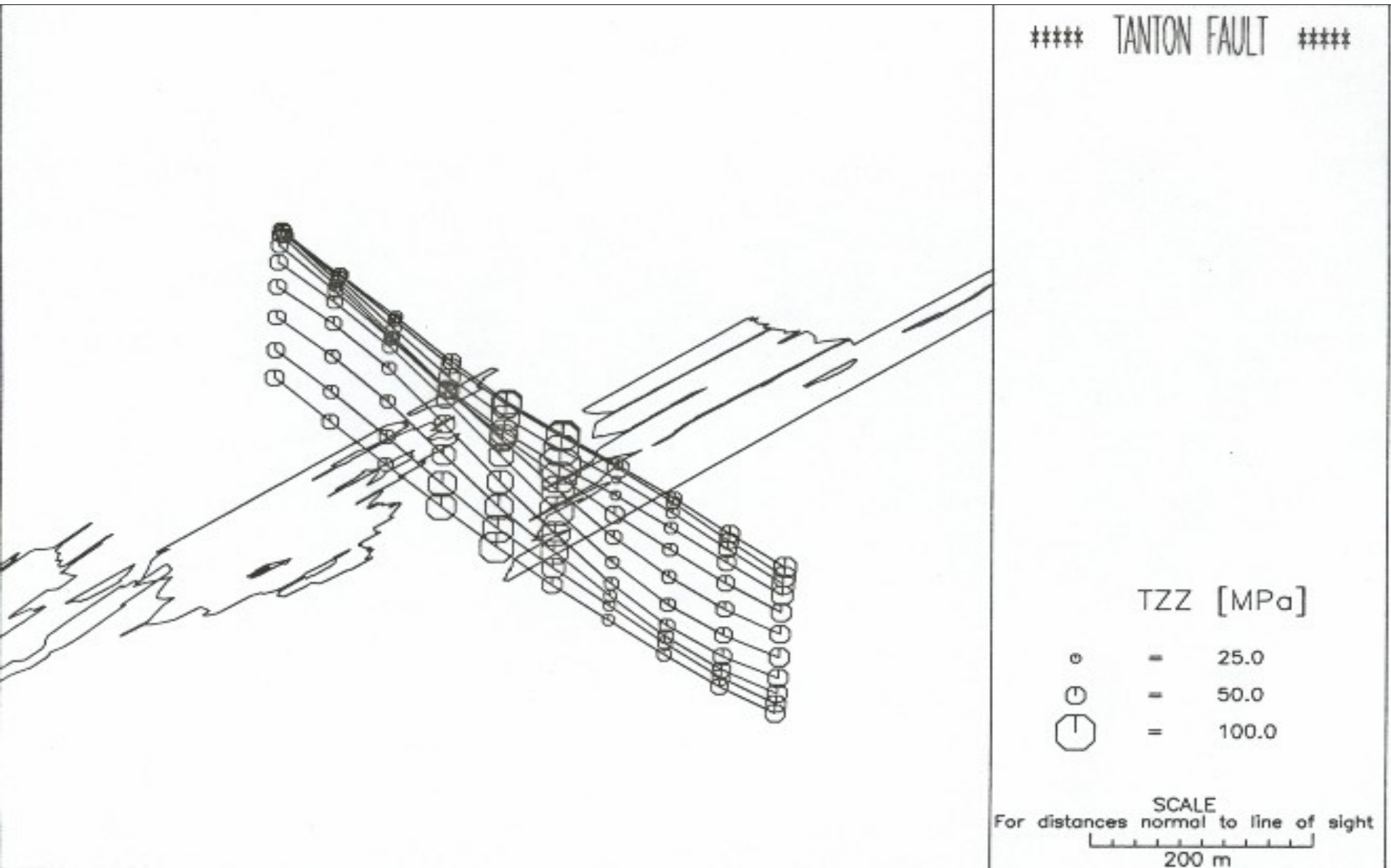
The E-M relation

$$\log(E) = c + d \cdot \log(M)$$


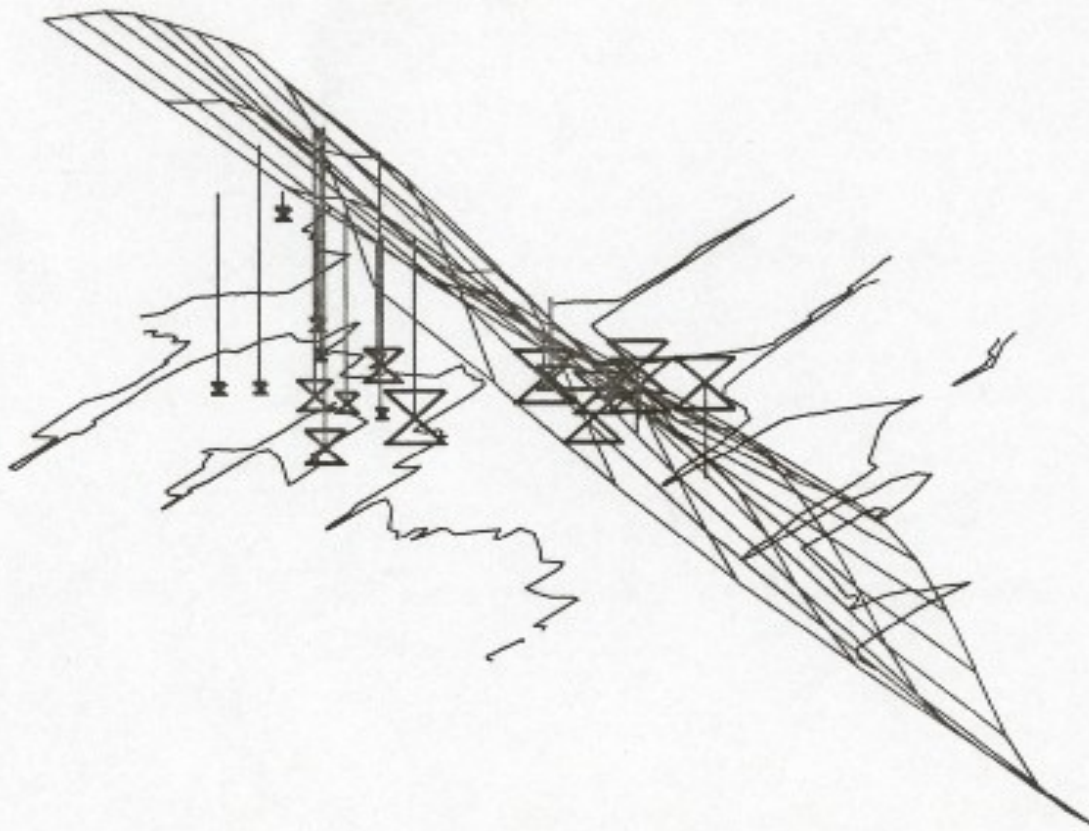
The E-M relation: EI at work



The E-M relation: EI at work



The E-M relation: EI at work



TANTON FAULT ####
TIME SPAN
930501.0000 - 930930.2400

ENERGY INDEX [obse/expe]	
✕	< 0.5
⊗	1.0
⊗	2.0 - 3.0
⊗	3.0 - 4.0
⊗	> 4.0

SCALE
For distances normal to line of sight
100 m

Apparent volume



Seismic source volume scales with
moment/stress drop

Replacing stress drop with
apparent stress:

$$\text{moment}/(E * G/M)$$

Apparent volume



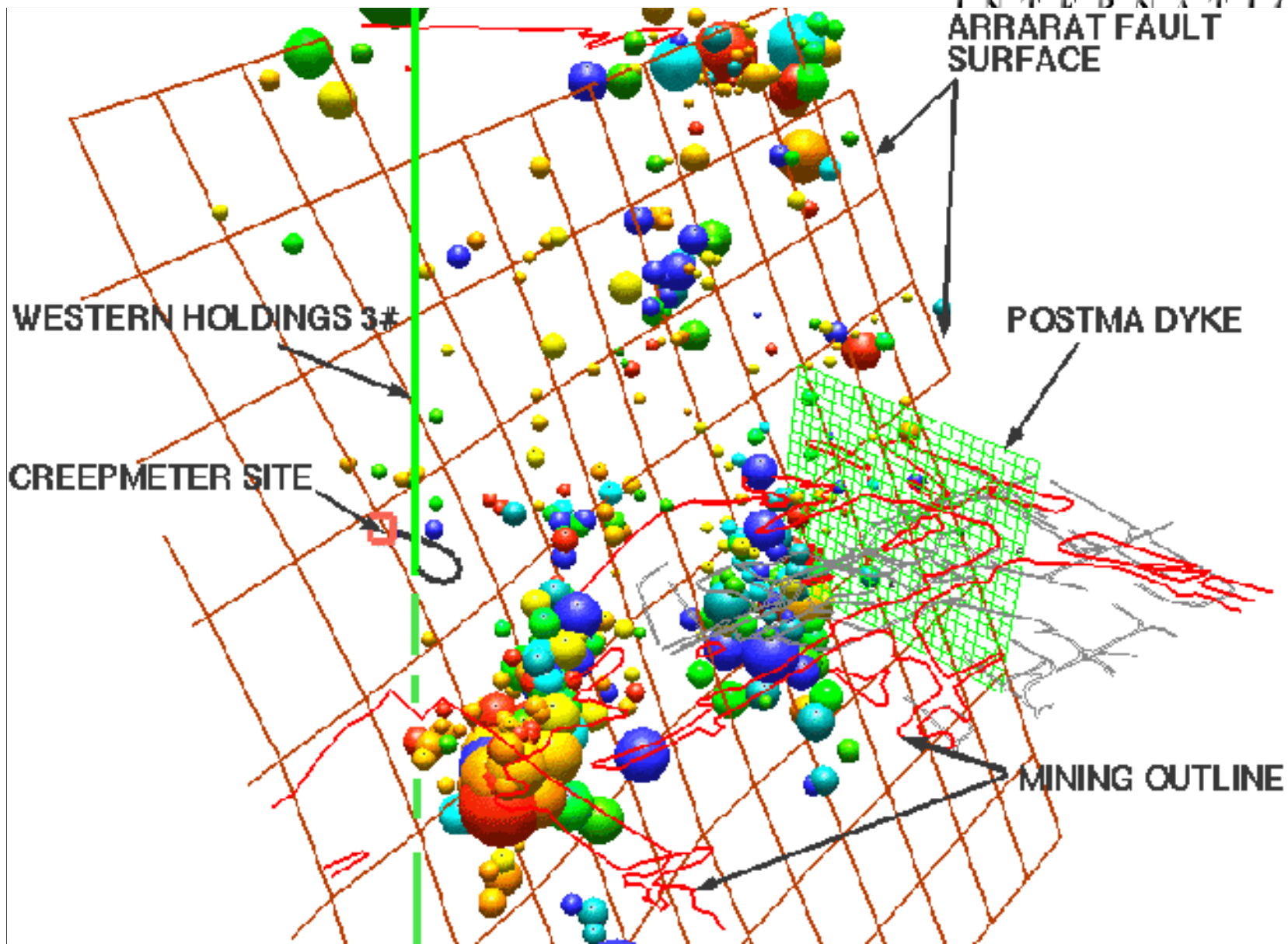
$$V_A = M(c_3 \sigma_A) = M^2 / (c_3 G E)$$

Apparent volume, [m³]

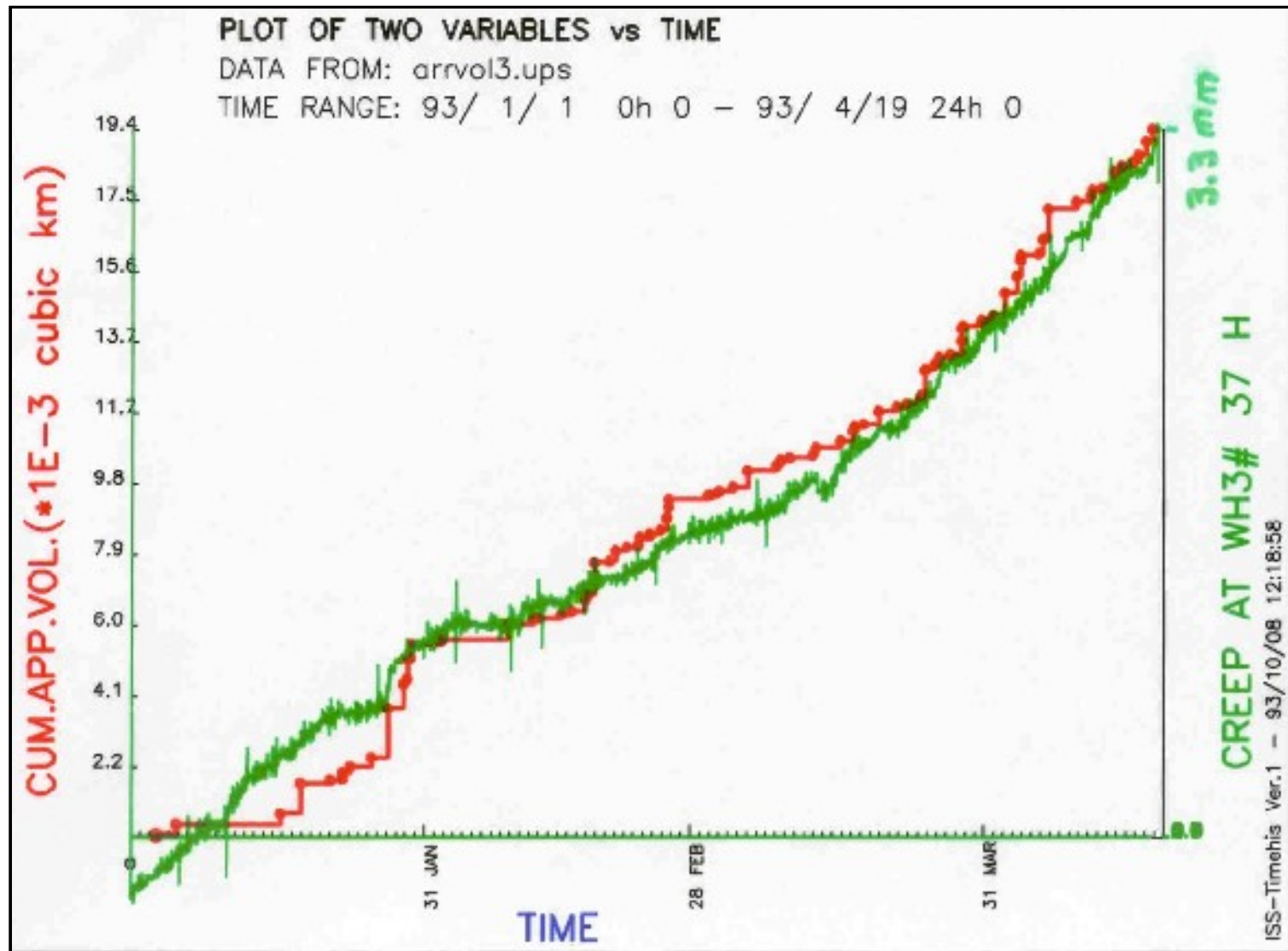
c_3 – scaling factor ~ 2 .

The apparent volume scales the volume of rock with co-seismic inelastic strain of an order of apparent stress over rigidity. The apparent volume V_A is less model dependent than the source volume V .

A creepy fault



Subsequent fault creep and ΣV_A



More derived parameters

<p>Seismic strain, $\epsilon_s (\Delta V, \Delta t) = \Sigma M / (2G\Delta V)$ and Seismic strain rate, [s⁻¹] $\dot{\epsilon}_s (\Delta V, \Delta t) = \epsilon_s / \Delta t$</p>	<p>Seismic strain measures strain due to cumulative coseismic deformations within the volume ΔV over the period Δt. Its rate is measured by $\dot{\epsilon}_s$.</p>
<p>Seismic stress, σ_s[Pa] $\sigma_s (\Delta V, \Delta t) = 2G\Sigma E / \Sigma M$</p>	<p>Seismic stress measures stress changes due to seismicity.</p>
<p>Seismic stiffness modulus, K_s [Pa] $K_s (\Delta V, \Delta t) = \sigma_s / \dot{\epsilon}_s = 4G^2 \Delta V \Sigma E / (\Sigma M)^2$</p>	<p>Seismic stiffness measures the ability of the system to resist seismic deformation with increasing stress. The stiffer systems limit both the frequency and the magnitude of intermediate and large events but have time-of-day distribution with larger statistical dispersion, thus are less time predictable.</p>
<p>Seismic viscosity, [Pa · s] $\eta_s (\Delta V, \Delta t) = \sigma_s / \dot{\epsilon}_s$</p>	<p>Seismic viscosity characterises the statistical properties of the seismic deformation process. Lower seismic viscosity implies easier flow of seismic inelastic deformation or greater stress transfer due to seismicity.</p>

More derived parameters

Seismic diffusivity, [m²/s]

$$D_s(\Delta V, \Delta t) = (\Delta V)^{\frac{2}{3}} / \tau_s,$$

or in a statistical sense

$$d_s = (\bar{X})^2 / \bar{t}.$$

Seismic diffusivity can be used to quantify the magnitude, direction, velocity and acceleration of the migration of seismic activity and associated transfer of stresses in space and time. There is an inverse relationship between the diffusivity D_s and the friction parameters.

Seismic Deborah number

$De_s(\Delta V, \Delta t) = \tau_s / \text{flowtime}$, where *flowtime* is a design parameter not necessarily equal to Δt .

Seismic Deborah number measures the ratio of elastic to viscous forces in the process of seismic deformation and has successfully been used as a criterion to delineate volumes of rockmass softened by seismic activity (soft clusters). The lower the Deborah number the less stable is the process or the structure over the design *flowtime* - what may be stable over a short period of time (large De_s) may not be stable over a longer time (lower De_s).

Seismic Schmidt number

$$Sc_{sd}(\Delta V, \Delta t) = \eta_s / (\rho D_s) \text{ or}$$

$$Sc_{sd} = \eta_s / (\rho d_s)$$

where ρ is rock density.

Seismic Schmidt number measures the degree of complexity in space and time (the degree of turbulence) of the seismic flow of rock. Note that seismic Schmidt number Sc_{sd} , encompasses all four independent parameters describing seismicity:

$$\bar{t}, \bar{X}, \sum M, \sum E$$

Seismic Schmidt number



Seismic Schmidt number measures the degree of complexity in space and time (the degree of turbulence) of the seismic flow of rock. Note that seismic Schmidt number Sc_{sd} encompasses all four independent parameters describing seismicity:

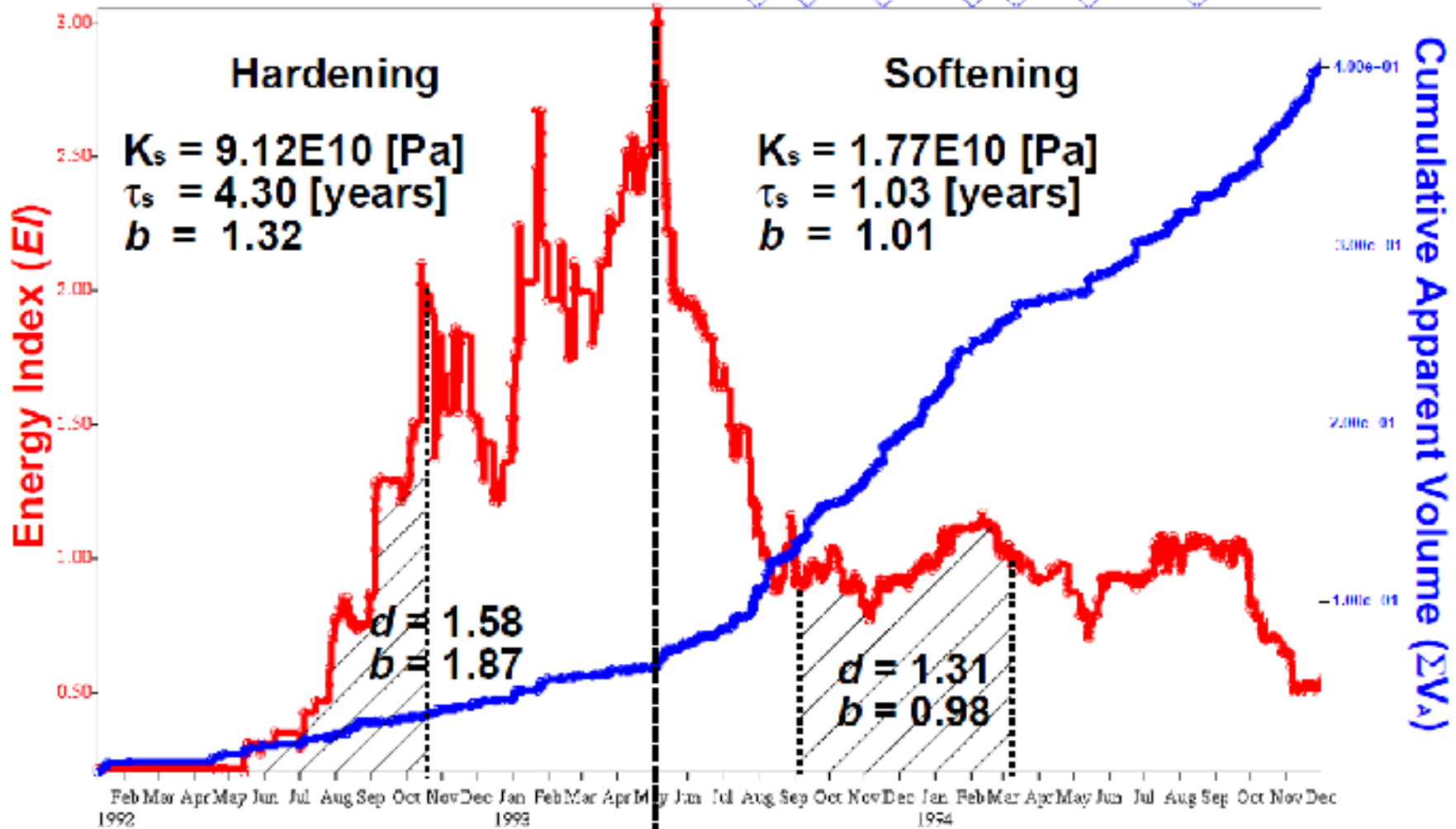
$$\bar{t}, \bar{X}, \Sigma M, \Sigma E$$

EI, Apparent Volume



INTERNATIONAL

log E = 7.32 8.23 8.23 7.58 8.11 7.92 8.15

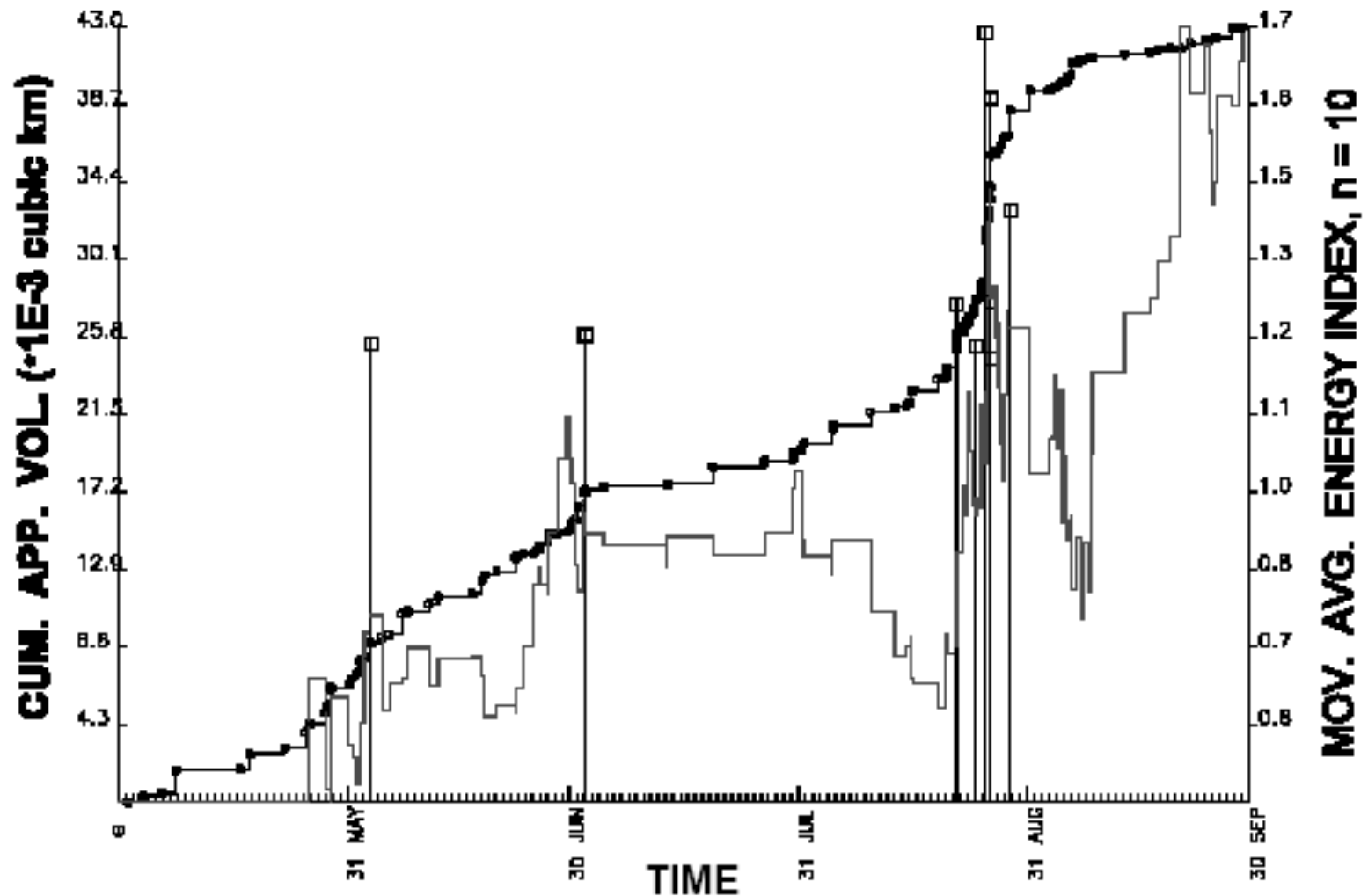


Mendecki & van Aswegen, 1997

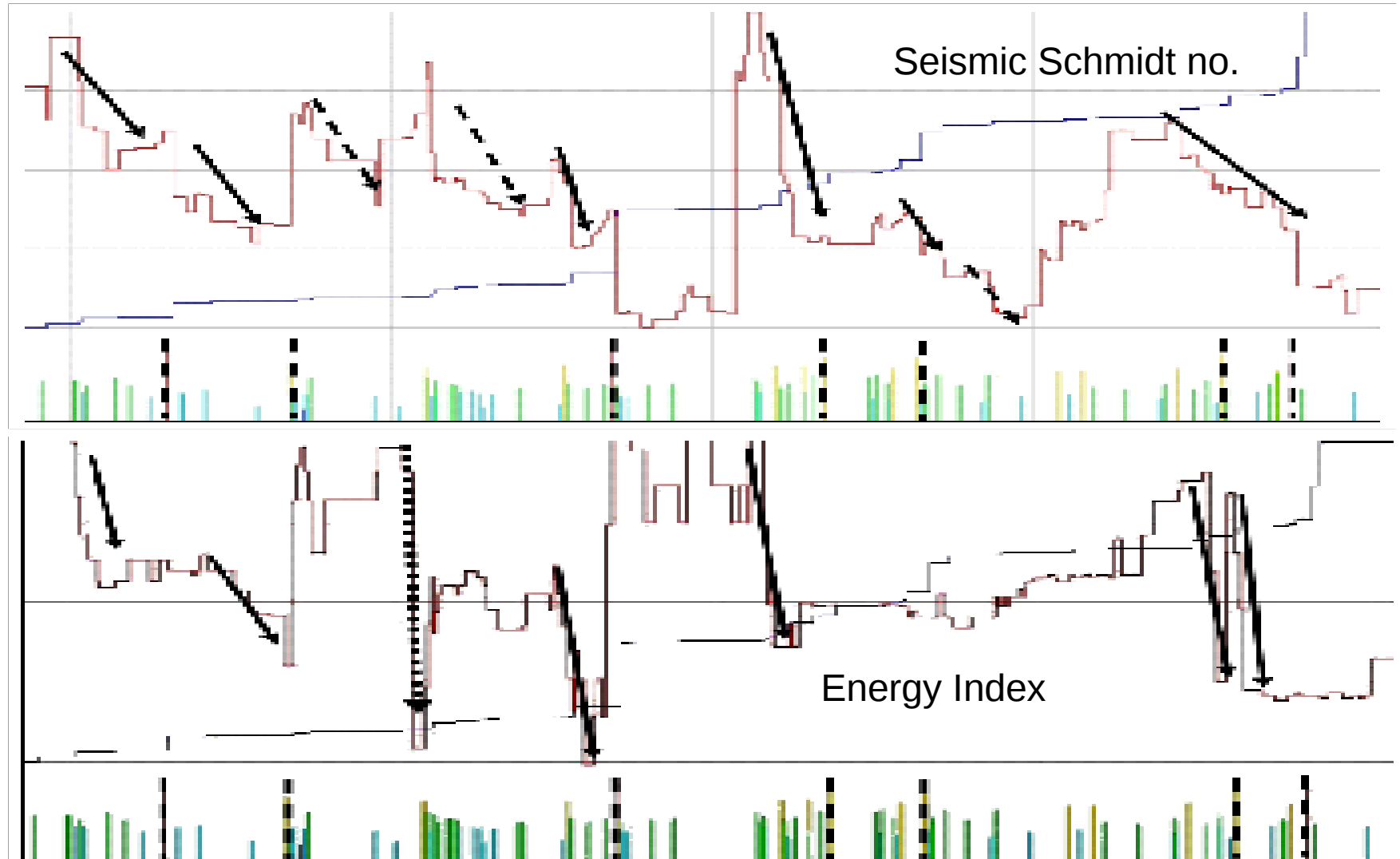
Short term seismic hazard assessment

Through time history analysis

Time history analyses



Simplistic time history analyses



Vertical broken lines depict events \geq mag 1.8

Useful parameters for short term stability assessment

Parameter	Measures
Energy Index	Stress
Apparent volume – the slope of the cumulative curve	Strain rate
Seismic Schmidt number	Turbulence of seismic deformation
Seismic activity rate	Number of seismic events per time

RRoSH - rules



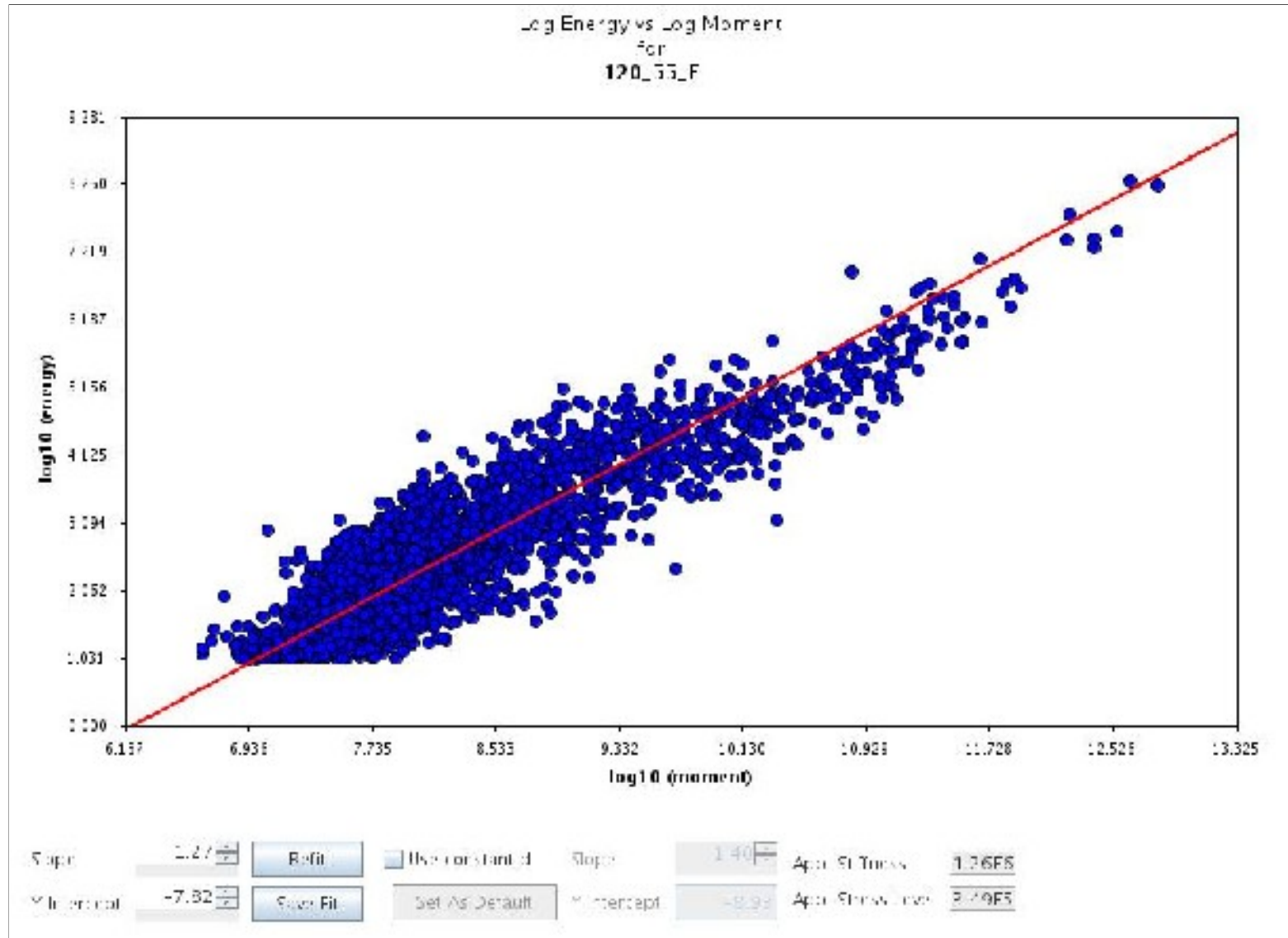
Parameter \ Rating	0	1	2
Cum. Apparent volume	No tendency towards power law behaviour	Weak tendency to power law behaviour	Strong tendency to power law behaviour
Log(Energy Index)	Absolute value of change < .25	$0.25 \geq$ absolute value of change ≤ 0.25	absolute value of change ≥ 0.5
Log(seismic Schmidt no.)	Absolute value of change < 0.5	drop in value ≥ 0.5 ≤ 1.0 ; increase in value ≥ 0.5	drop value ≥ 1.0
Activity rate	Average	Above average, < 75% of 100 day peak	> 75 % of 100 day peak

In addition, anomalous spatial patterns judged (qualitatively) and rating the increased by 1 or 2

Example: 120-55-E



The E-M relation

$$\log(E) = c + d \cdot \log(M)$$


Upper Truncated GR



CUMULATIVE FREQUENCY - UPPER TRUNCATED MAGNITUDE DISTRIBUTION ($>= m$) - $a(10^{bm} - 10^{bm_{max}})$

Filter Name: 120_05

Total Number of Events: 6922

Largest Event: $m = 2.9$ (2005-05-19 00:17:38)

First Event: 2003-09-10 13:24:54

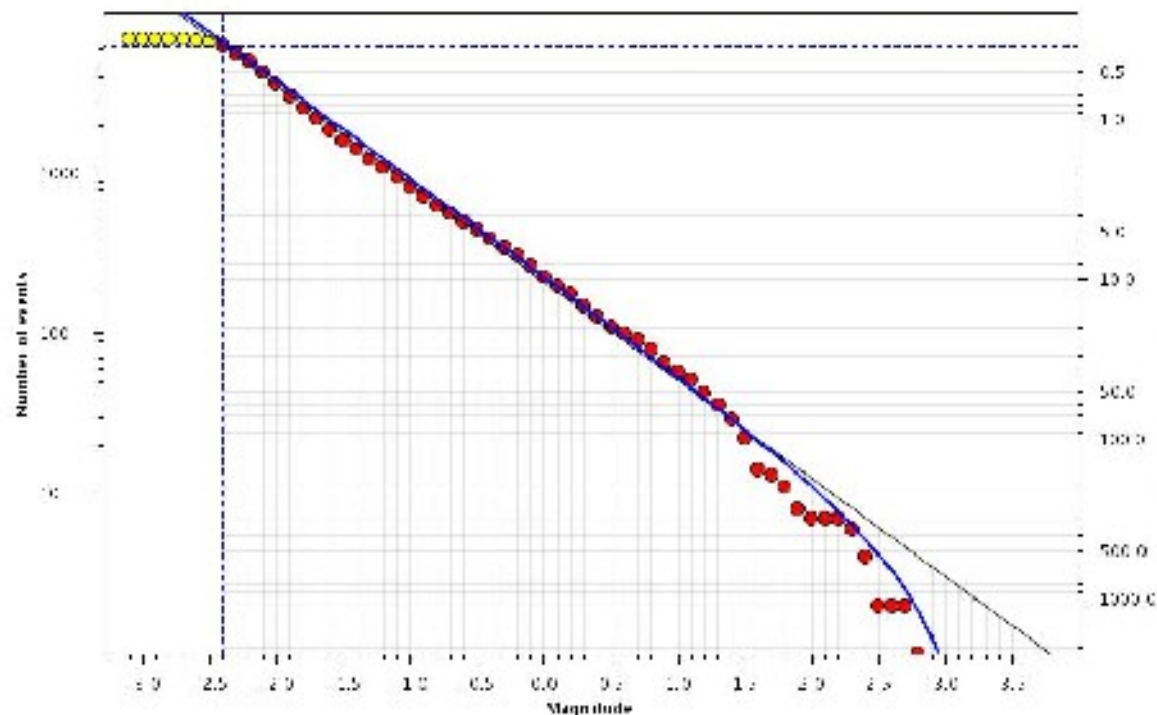
Interval Size: 0.1

Second Largest Event: $m = 2.5$ (2005-06-08 18:10:58)

Last Event: 2009-05-27 01:16:09

Time Span: 2177 days

Third Largest Event: $m = 2.5$ (2000-08-20 04:54:55)

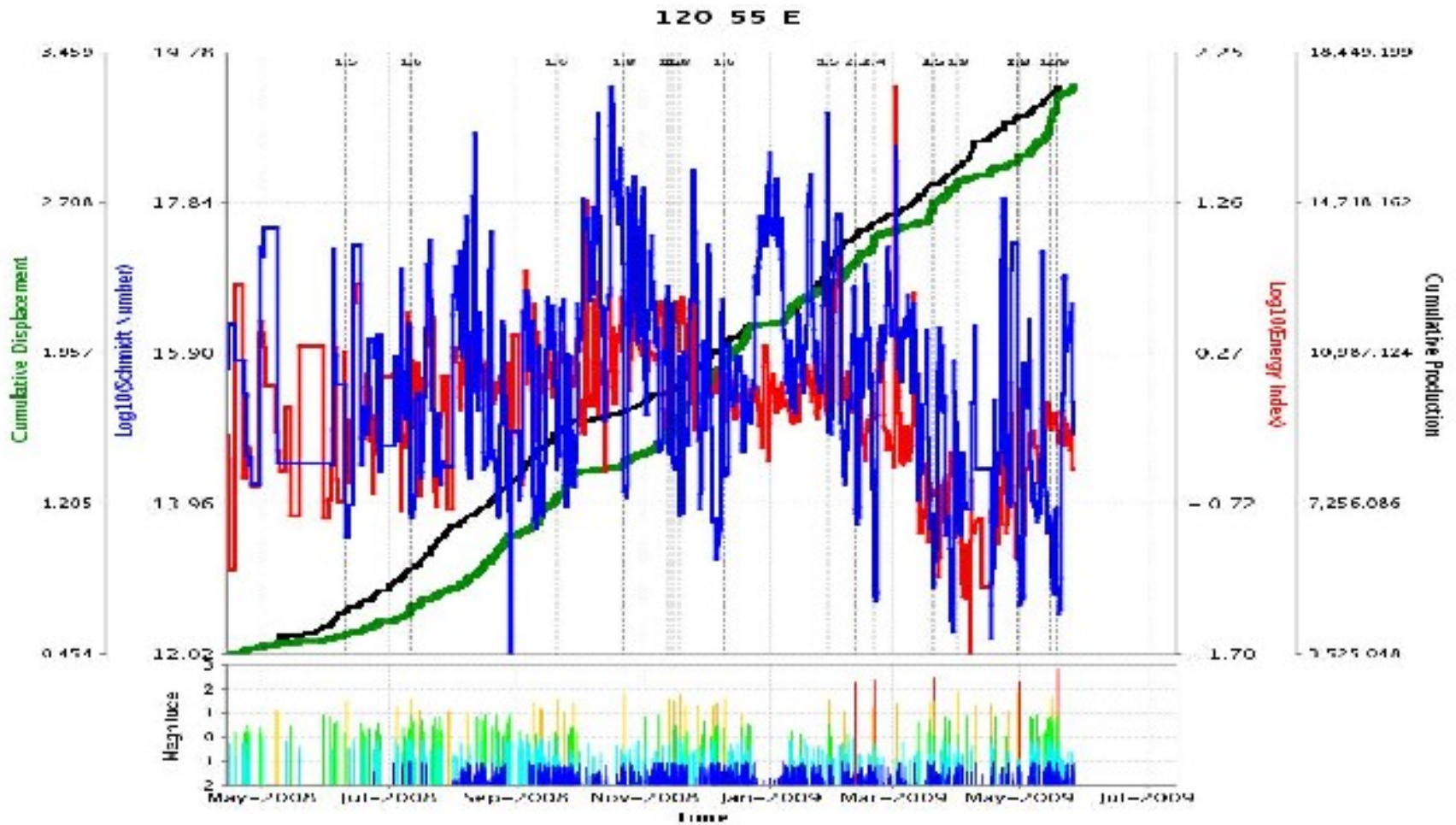


m_{min}	= 2.4
$N(>= -2.4)$	= 6250
a [year]	= 1.58
a	= 2.36
b	= 0.872 (+0.01)
m_{max}	= 2.2 (+0.00)
\bar{t} [mean]	= 30101 d
\bar{x} [mean]	= 105.40 m
$\Delta \bar{t}(-2.4, 3.2)$	= 1.47e+08 d
$\Delta \bar{x}(1.0, 3.2)$	= 1.51e+08 m
$\Delta \bar{t}(2.4, 3.2)$	= 3.11e+07 d
$\Delta \bar{x}(1.0, 3.2)$	= 6.25e+07 m
Hazard Magnitude	= 2.49
Pot. Dam. Rad.	= 1.28e+02 m
Pot. Dam. Vol.	= 8.91e+06 m ³

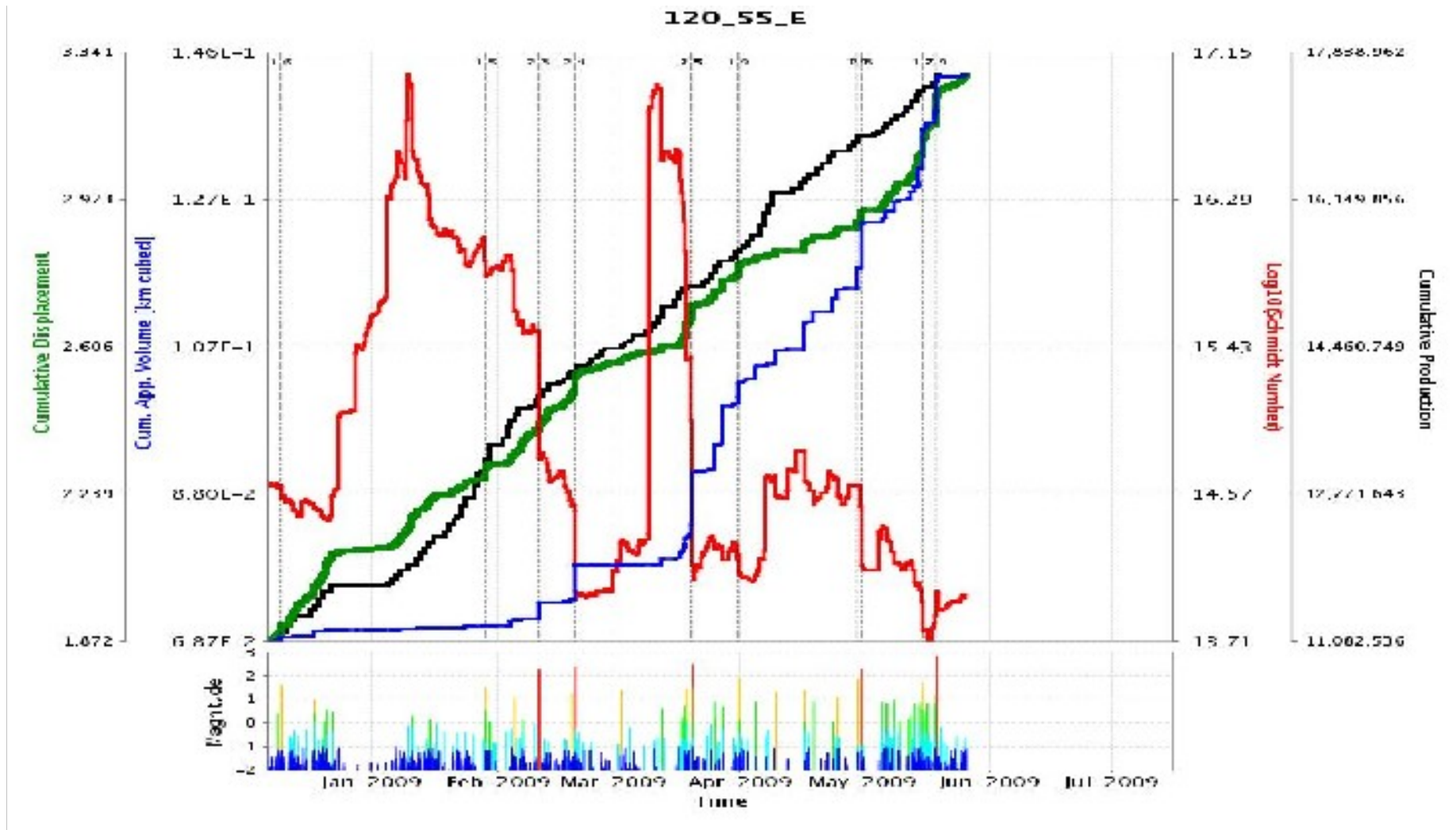
Probability Table and Recurrence Times

	$m = 1.7$	$m = 1.9$	$m = 2.1$	$m = 2.3$	$m = 2.5$	$m = 2.7$	$m = 2.9$	$m = 3.1$
mean recurrence time	110	188	316	544	926	1581	2781	4796
5% recurrence	0.123	0.073	0.054	0.037	0.027	0.019	0.013	0.009
1% recurrence	0.028	0.017	0.012	0.008	0.006	0.004	0.003	0.002
10% recurrence	0.123	0.198	0.298	0.216	0.147	0.091	0.048	0.031
1% recurrence	0.298	0.547	0.914	0.387	0.277	0.177	0.097	0.061
5% recurrence	0.042	0.070	0.103	0.028	0.019	0.013	0.009	0.007

120-55-E TH



120-55-E TH



Night shift report

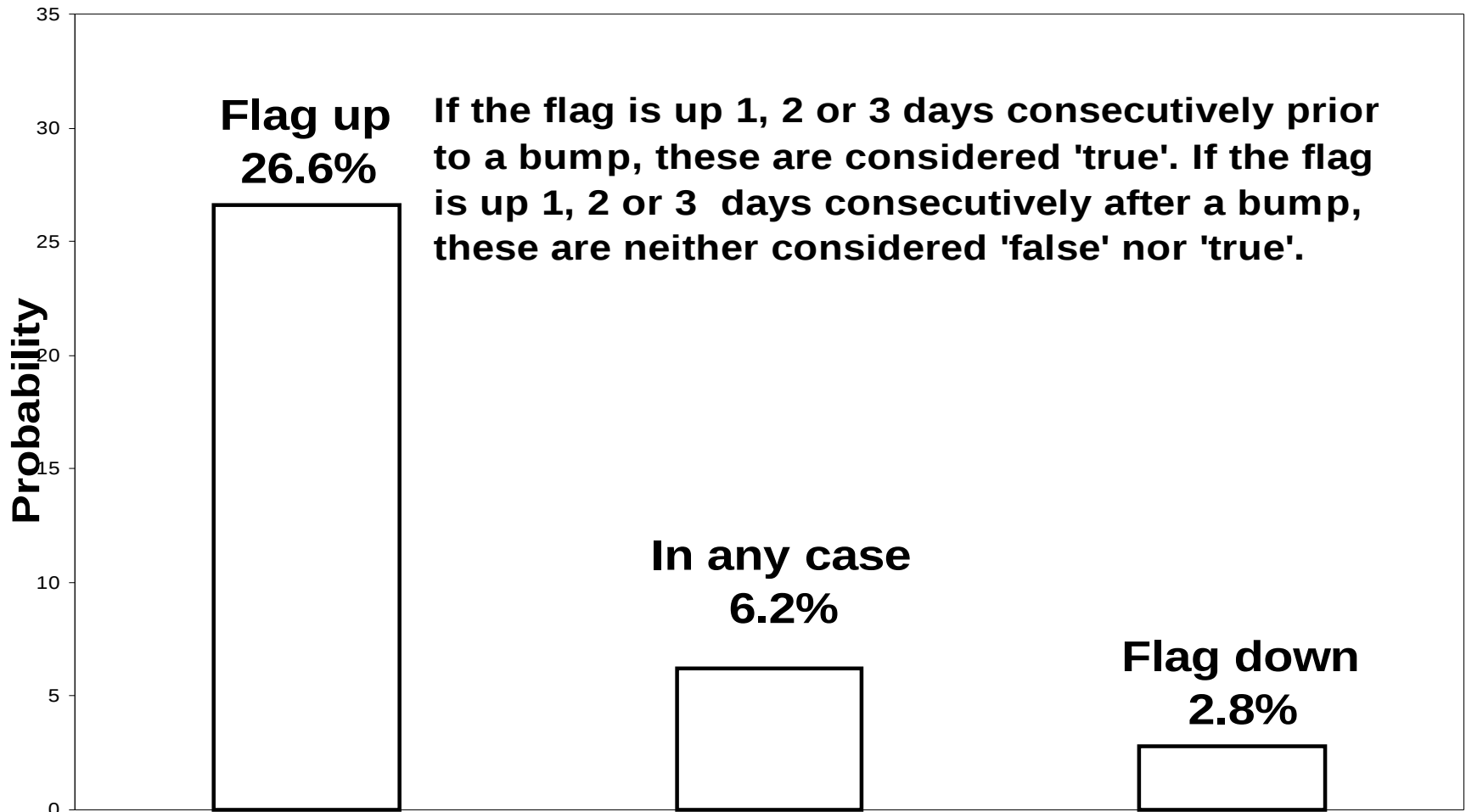
(after blasting, before night shift):

robot system (18h30 – 21h00)

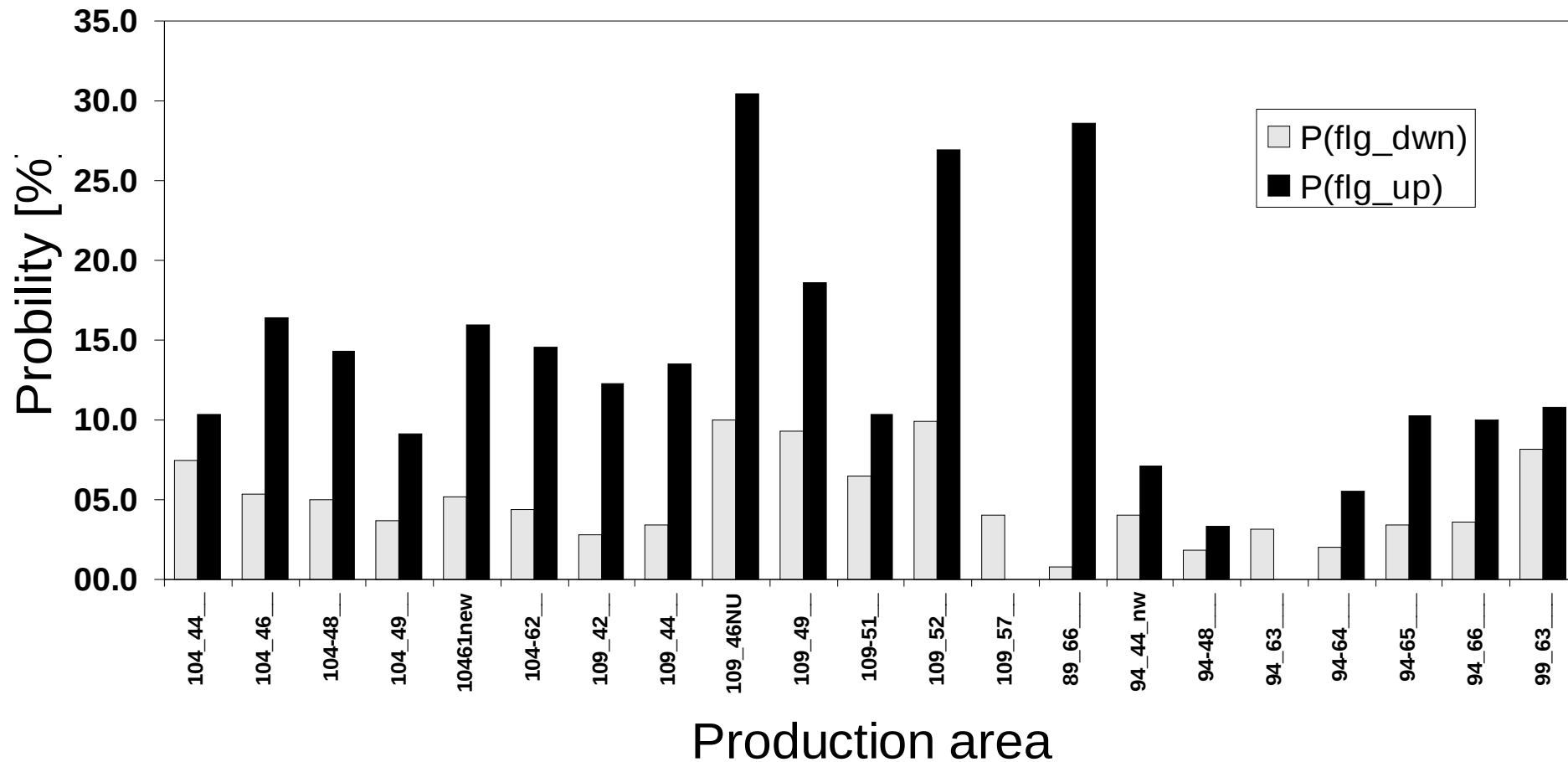


Poly	Last Event Date	Last Event Time	New Events	Va	LogICE	Schmid	Act	Large Events	Current Rating	Comments / Observations
101_102	20030226	180515	3	1.0	1.0	2.0	1.0	M0.7@26/02,18FCU.	6.00C000	. maintain RED status
101p1r	20030223	160553	0	0.0	1.0	1.0	0.0	J	2.00C000	. nn
104_106_	20030226	180145	4	0.0	1.0	2.0	1.0	M* 1@26/02,16h01 M0.9@26/02,18h01 M0.7@26/02,18h01	4.00C000	. Schmidt is falling and act. up.
00E_ger_	20030226	100515	7	1.0	1.0	2.0	2.0	M1.1@26/02,16h01.. M0.7@26/02,18h00 . M0.9@26/02,18h01...V0.7@26/02,18h01..	6.00C000	. Activity is up.
107_109_	20030226	180055	1	0.0	2.0	1.0	0.0	M* 4@26/02,18h00	2.37C000	. At last some activity
111_112_	20030223	51559	0	0.0	1.0	1.0	0.0	J	1.80C000	. nn
114_115_	20030223	182640	0	0.0	1.0	1.0	0.0	J	1.80C000	. nn
114_16W	20030226	175303	1	0.0	2.0	1.0	0.0	M0.6@26/02,17h53.	3.00C000	.
33_112st	20030219	53209	0	0.0	0.0	0.0	0.0	J	0.00C000	. Still mining here ?? last event 6 days back
118_120	20030226	174512	1	0.0	2.0	2.0	0.0	J	4.00C000	. Long term Schmid and EI cropping strongly
120A___	20030221	175722	0	0.0	0.0	0.0	0.0	J	0.00C000	. Rating cropped old event
83_03___	20030226	182416	2	1.0	1.0	1.0	1.0	M* 7@26/02,18h23.. M1.7@26/02,18h24..	4.00C000	. Activity is up
64Wp1r	20030224	200554	0	1.0	2.0	2.0	0.0	J	5.50C000	. nn
97_130__	20030226	F2004	0	0.0	1.0	1.0	0.0	J	1.60C000	.
Jeck	20030226	135710	1	0.0	2.0	1.0	1.0	J	4.00C000	.

Mine 1: $P(\text{mag} \geq 1.5, \text{rating} \geq 5)$



P[mag. ≥ 1 ON SHIFT within 2 days of rating ≥ 5 Mponeng 2003 - 2004



Hazard assessment success rate: example case 1 - 2002



Phenomenon

Probability

The occurrence of a mag. ≥ 1 event on-shift **05.28%**

The occurrence of a mag. ≥ 1 event on-shift if the RRoSH ≥ 3 during the 60 hours before, i.e. **14.18%**

FLAG UP

The occurrence of a mag. ≥ 1 event on-shift if the RRoSH < 3 during the 60 hours before, i.e. **00.23%**

FLAG DOWN

Hazard assessment success rate: example case 2 - 2002



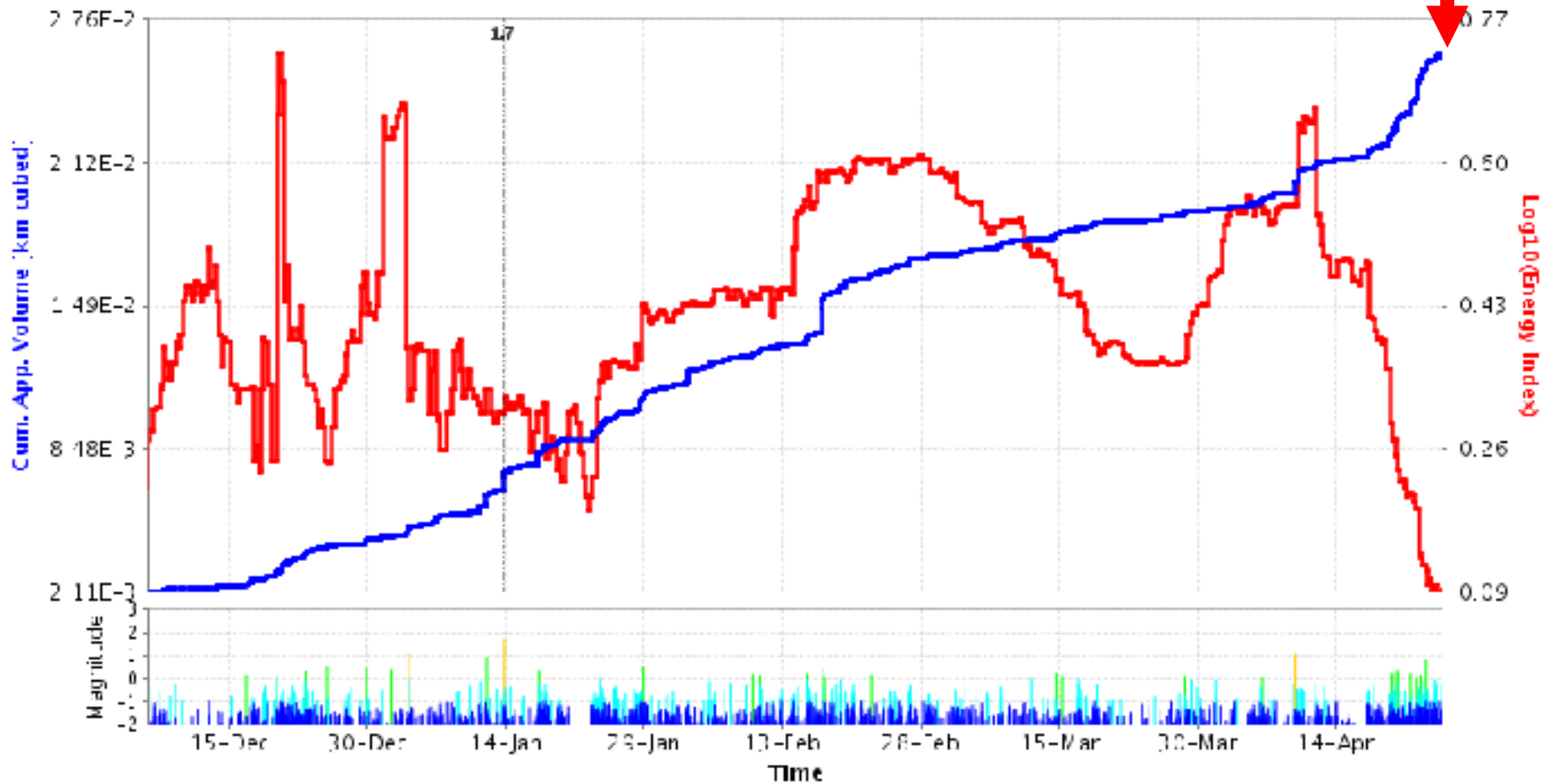
Phenomenon	Probability
The occurrence of a seismic event ≥ 1.0 on-shift	10.30
The occurrence of a seismic event ≥ 1.0 on-shift if the RRoSH ≥ 4 during the two days prior – FLAG UP	36.97
The occurrence of a seismic event ≥ 1.0 on-shift if the RRoSH < 4 during the two days prior – FLAG DOWN	02.52

**Australian mine: *m2.3*
fault-slip event**

Australian mine: EI, cum. apparent volume



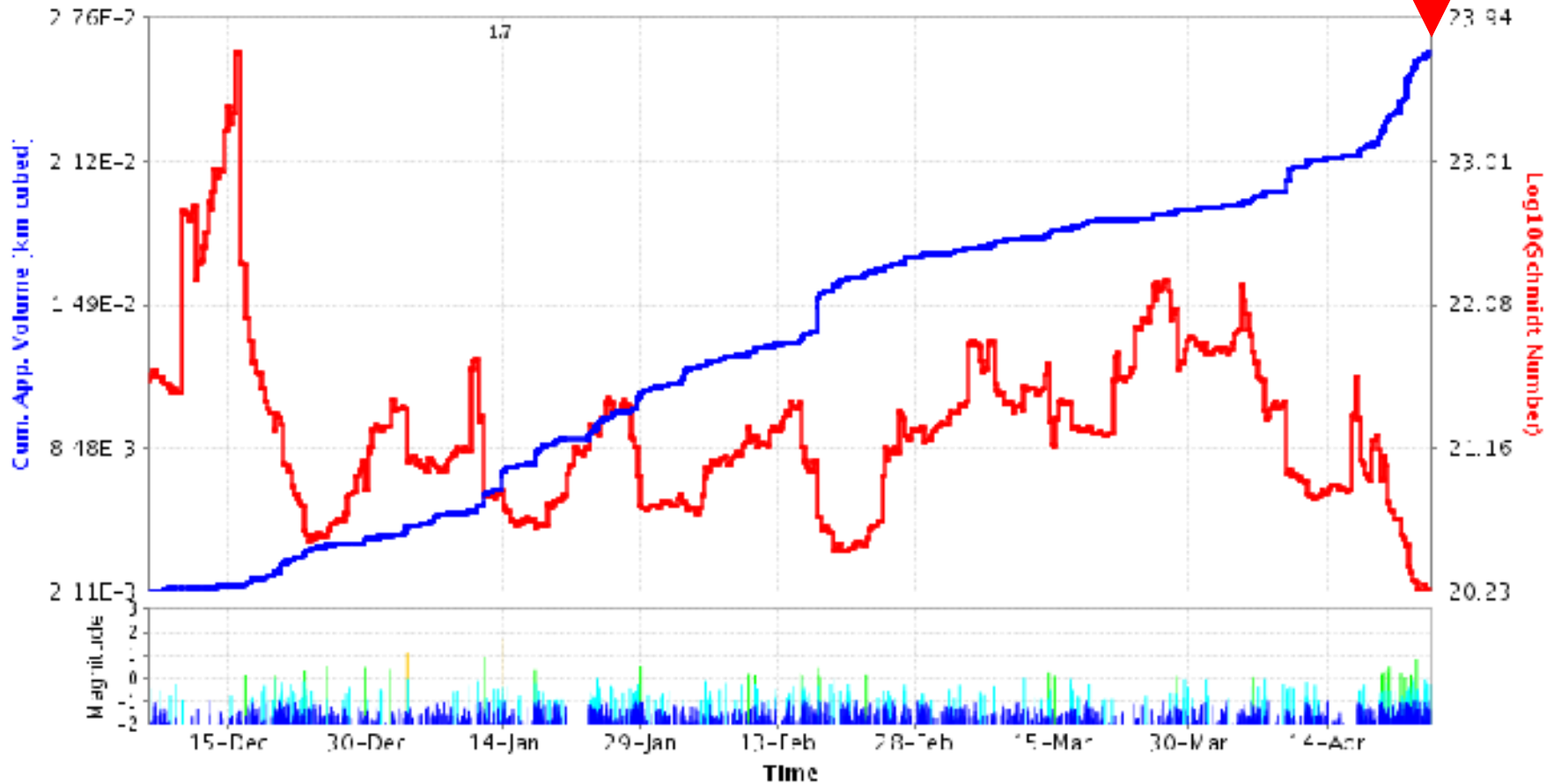
All



Australian mine: Sc, cum. apparent volume

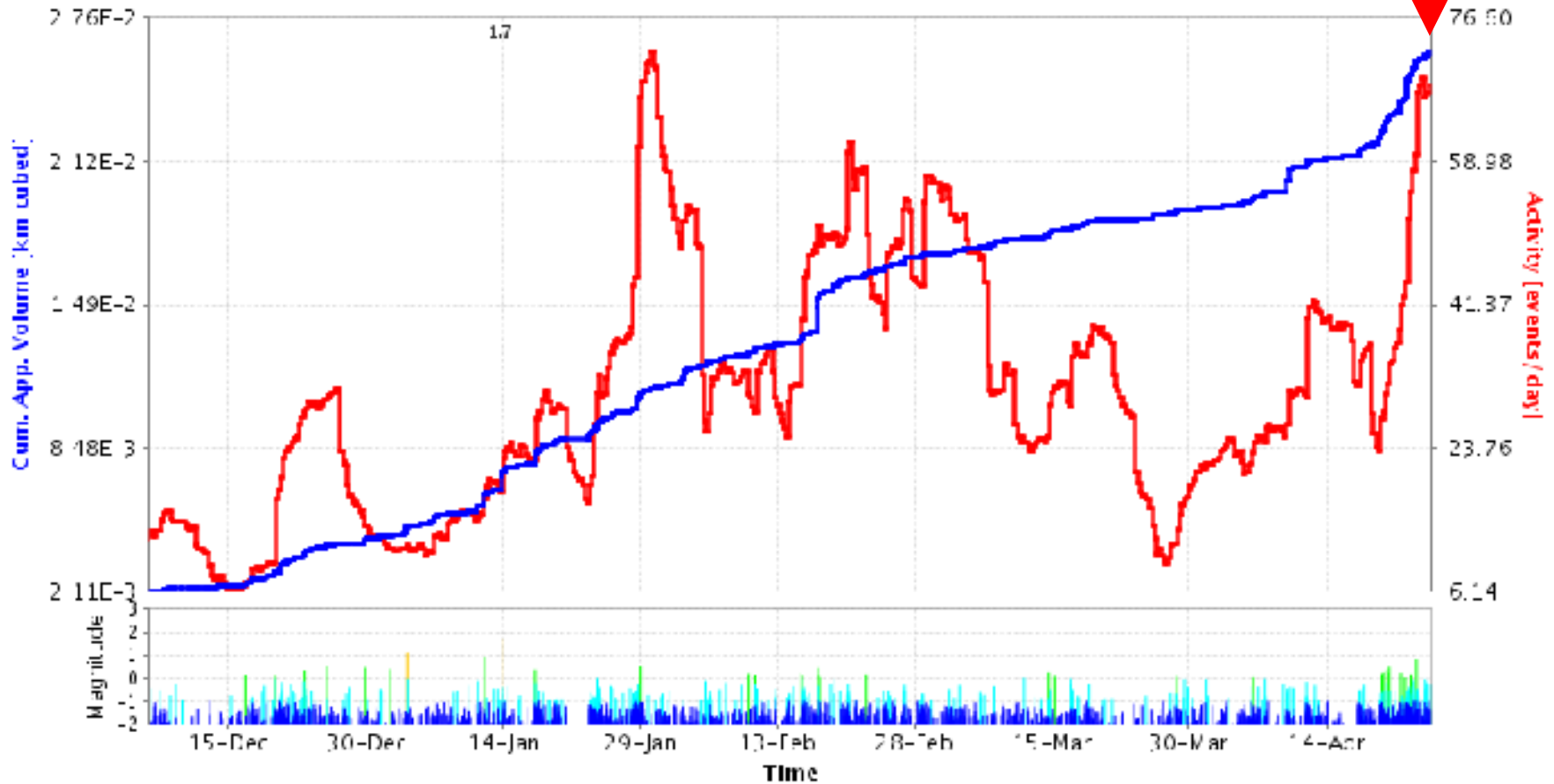


All



Australian mine: seismic activity

All



Conclusions



- We get more information from the seismic data ..
- Translate data into rock mechanics language
- Consider stability in term of stress, strain rate, turbulence, spatial patterns
- Last step is to integrate all this with on-line numerical modeling